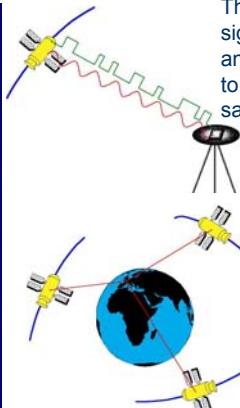
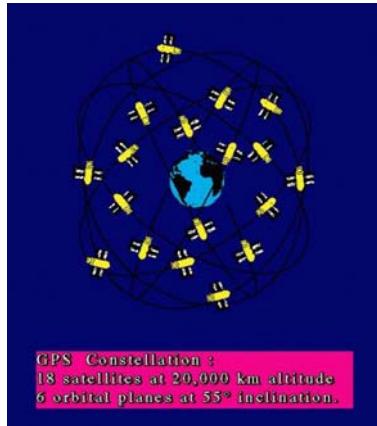


GPS (Global Positioning System)



GPS was created in the 80s' by the US Department of Defense for military purposes. The objective was to be able to get a precise position anywhere, anytime on Earth.



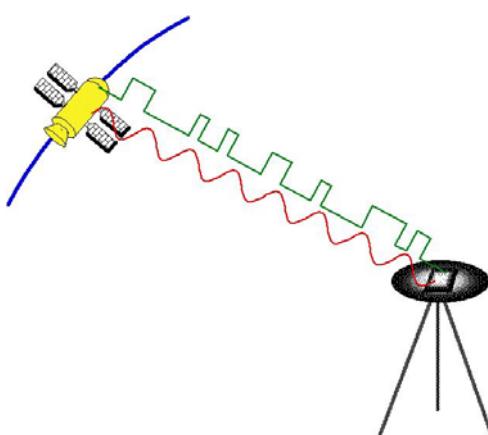
The satellites send a signal, received by a GPS antenna. Again, this allow to measure the distance satellite to antenna

With at least 3 satellites visible at the same time, we can compute instantaneously the station position. The precision can be as good as 1 millimeter

1

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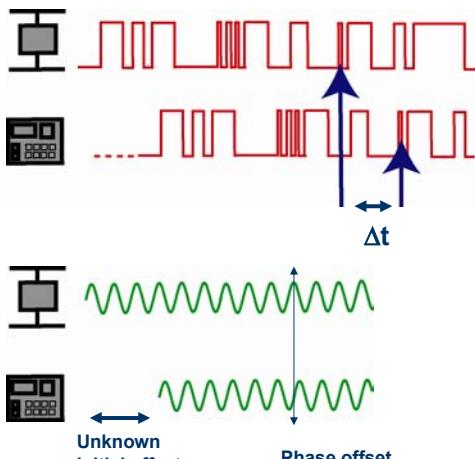
Fundamentals of GPS



2

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GPS (Global Positioning System)



pseudo-distance Measurement:

Accurate to 30 m if C/A code
(pseudo frequency of 1 MHz)

Accurate to 10 m if P code
(pseudo frequency of 10 MHz)

Easy because code never repeats
itself over a long time, i.e. no
ambiguity

Phase Measurement:

Accurate to 20 mm on L1 or L2
(1.5 GHz)

But difficult because the initial
offset is unknown.

=> Post processing of a sequence
of measurements on 1 satellite
give final station position

3

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GPS signals

- 2 carrier phases (phases porteuses)



-L1 = 1575.42 MHz, $\lambda = 19 \text{ cm}$

-L2 = 1227.60 MHz, $\lambda = 24 \text{ cm}$

- 2 carried codes (codes portés)



-P1 (C/A) ~1.023 MHz, $\lambda \sim 293 \text{ m}$

-P2 (précis) ~10.23 MHz, $\lambda \sim 29.3 \text{ m}$

C/A : Coarse acquisition = code "grossier" carried on L1

P : Precise = code "précis" carried on L2

P code is encrypted (by US army). This is known as "AS: Anti spoofing". the encrypted P-code is usually called the Y-code. Deciphering the Y-code is done through Z technology (squaring of Y-code).

4

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US army - limiting access to civilians: SA & AS

5

SA: Selective availability

- Until May 2000, US DoD artificially degraded stability of satellite clocks
- No solution for absolute and/or real time positionning. Solved by double differences (provided that GPS receivers sample the GPS signal at the same time).
- Stopped after 2000 (by Al. Gore)

AS: Anti Spoofing

- P-Code encrypted into Y-code
- done by multiplying P by an unknown code at 20KHz
- deciphered by “Z-tracking” (squaring of Y-code) receivers
- Still active today

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Offsets measurements are biased by Clocks

Satellites

- “Stable” atomic clocks (unless SA is active).
Instability = $\sim 10^{-14}$ ($\Rightarrow \sim 10^{-9}$ sec./day)
- Synchronized between all satellites
- Navigation message contains clocks corrections ==> can be modeled

Receivers

- Unstable cheap clocks: $(10^{-5} - 10^{-6})$ ($\Rightarrow \sim 1$ sec./day)

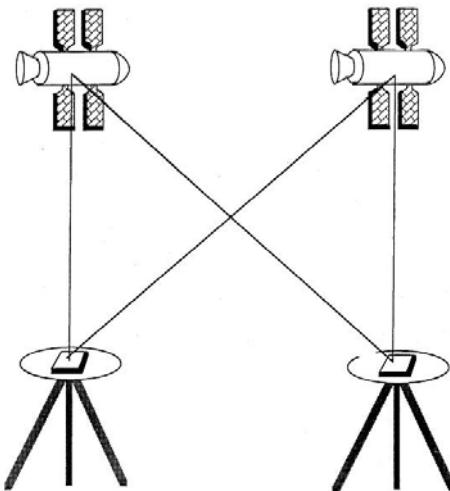
Problem : at 300 000 km/s, satellite-station distance is covered in 70 ms, and a 1m difference in position corresponds to 10^{-9} seconds...



Solution: Double difference

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Double differences



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One way phases are affected by **stations and satellites** clock uncertainties

Single differences are affected by **stations** clock uncertainties

Or

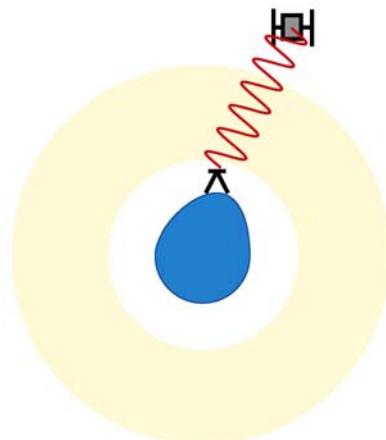
satellites clock uncertainties

Double differences Are free from all clock uncertainties **but**

=> Measurement of distances between points (= **baselines**)

=> **Relative positioning**

Other perturbation : The ionosphere



Correct measurement in an empty space

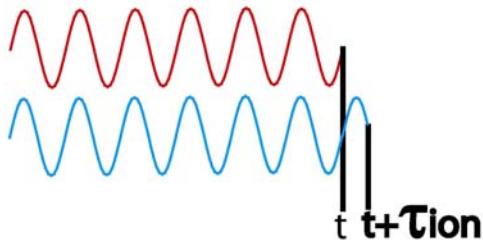
But the ionosphere perturbs propagation of electric wavelength

... and corrupts the measured distance

... and the inferred station position

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Ionosphere theory



Ionospheric delay τ_{ion} depends on :

- ionosphere contains charged particles (ions and electrons) : N_e
- Frequency of the wave going through the ionosphere : f

$$\tau_{ion} = 1.35 \cdot 10^{-7} \cdot N_e / f^2$$

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Ionosphere : solution = dual frequency

Problem : N_e changes with time and is never known

solution : sample the ionosphere with 2 frequencies

$$\tau_{ion_1} = 1.35 \cdot 10^{-7} \cdot N_e / f_1^2$$

$$\tau_{ion_2} = 1.35 \cdot 10^{-7} \cdot N_e / f_2^2$$

$$\Rightarrow \tau_{ion_2} - \tau_{ion_1} = 1.35 \cdot 10^{-7} \cdot N_e \left(\frac{1}{f_2^2} - \frac{1}{f_1^2} \right)$$

$$\Rightarrow N_e = \left[\tau_{ion_2} - \tau_{ion_1} \right] / 1.35 \cdot 10^{-7} \left(\frac{1}{f_2^2} - \frac{1}{f_1^2} \right)$$

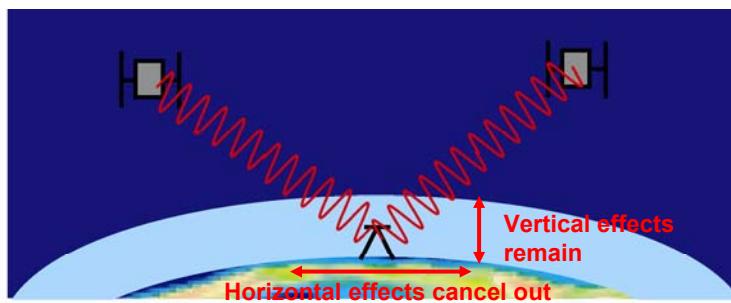
Using dual frequency GPS, allow to determine the number N_e and then to quantify the ionospheric delay on either L1 or L2.

(in fact, GPS can and is used to make ionosphere Total Electron Content (TEC) maps of the ionosphere)

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Second perturbation : The Troposphere



The troposphere (lower layer of the atmosphere) contains water. This also affects the travel time of radio waves.

But the troposphere is not dispersive (effect not inversely proportional to frequency), so the effect cannot be quantified by dual frequency system. Therefore there a position error of 1-50 cm.

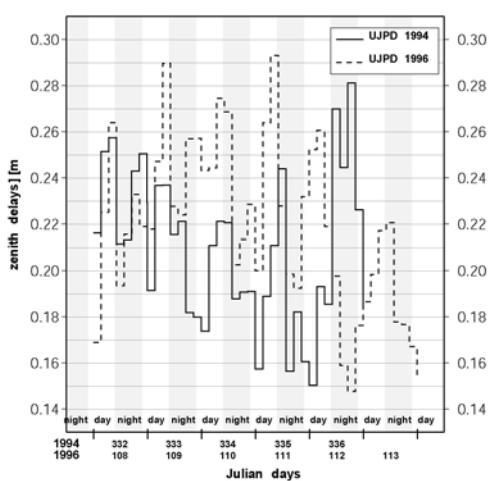
Thanks to the presence of many satellites, the effect cancel out (more or less) in average, on the horizontal position. Only remains a vertical error called **Zenith tropospheric delay**

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Troposphere zenith delay

Atmospheric Parameters at Ujung Pendang (Indonesia)



The tropospheric zenith delay can be estimated from the data themselves...

if we measure every **30s** on **5** satellites, we have **1800** measurements in **3 hours**.
We only have **3** unknowns : station **lat**, **lon**, and **altitude** !

So we can add a new one : **1 Zenith delay every 3 hours**

The curves show that the estimated Zenith delay vary from 15 cm to 30 cm with a very clear day/night cycle

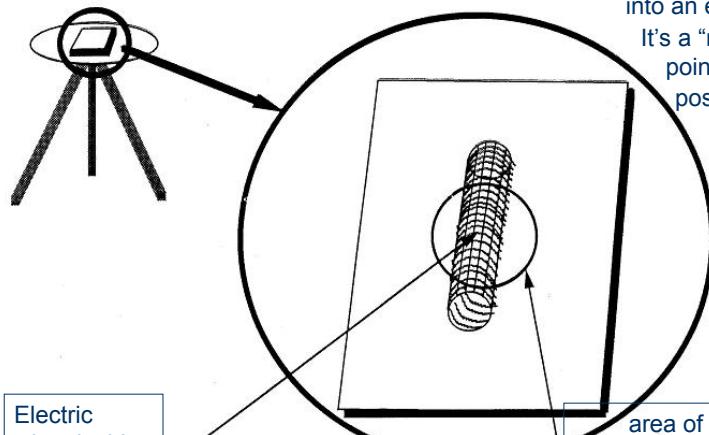
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Antenna phase center offset and variations

The Antenna phase center is the wire in which the radio wave converts into an electric signal. It's a "mathematical" point, which exact position depends on the signal alignment with the wire (azimuth and elevation)



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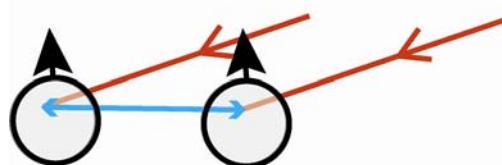
Antenna phase center offset and variations

Solution : use **identical** antennas, oriented in the **same direction**

As the signal rotates, the antenna phase centers move

But they move the same quantity in the same direction if antennas are strictly identical because the incoming signal are the same (satellite is very far away)

Therefore, the **baseline** between stations remains unchanged

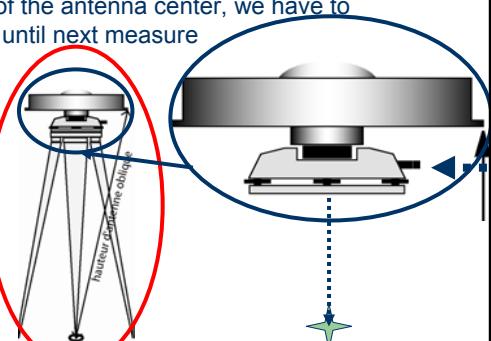
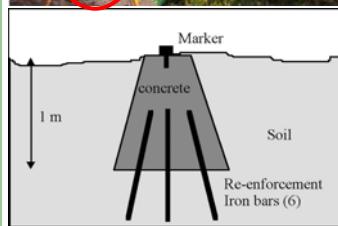


But this works for small baselines only (less than a few 100 km)

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Tripod and tribrachs source of errors

The measurement give the position of the antenna center, we have to tie it to the GPS marker which stays until next measure



The antenna has to be leveled horizontally and centered perfectly on the mark. Then :

Horiz. position of marker = horiz. position of antenna

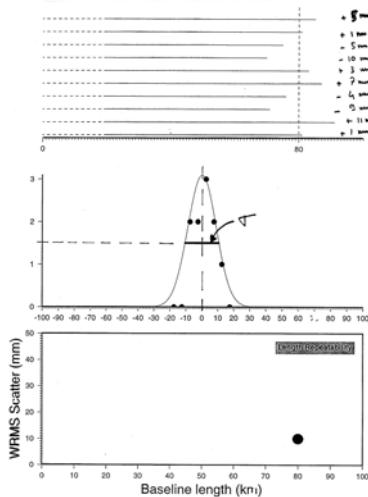
Altitude of marker = altitude of antenna – antenna height

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GPS est précis, à condition de :

1. GPS bi-fréquence pour éliminer l'ionosphère
2. Mesures longues pour « moyenner » la troposphère
3. Mesures relatives de distances entre stations
4. Utilisation d'antennes identiques, orientées parallèlement, et du centrage forcé sur les repères pour éviter les biais instrumentaux

Precision and repeatability



10 measurements of the same baseline give slightly different values :

80 km +/- 10 mm

How many measurements are between 80 and $80 + \delta$

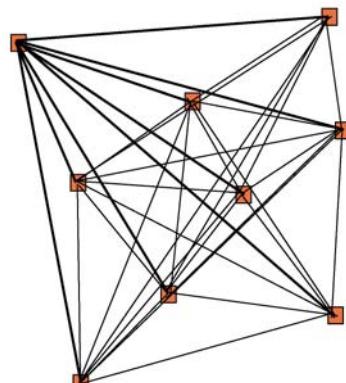
The histogram curve is a Gaussian statistic

The baseline **repeatability** is the sigma of its Gaussian scatter

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Network repeatabilities



Network of N points
(N=9)

(N-1) (=8) baselines from 1st station to all others

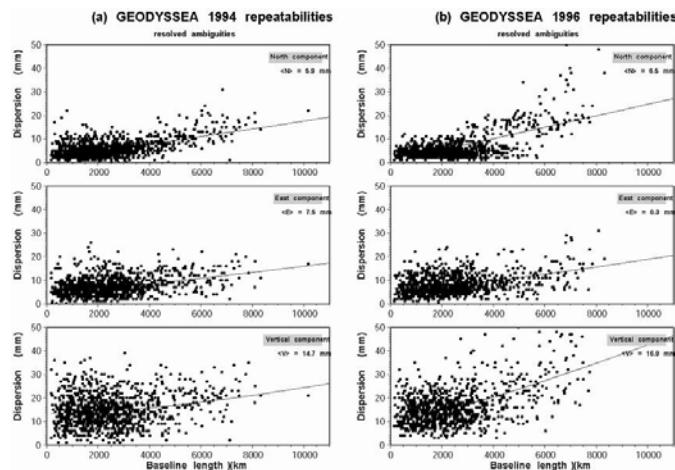
(N-2) (=7) baselines from 2nd station to all others
=> subtotal = (N-1)+(N-2)

total number of baselines
= (N-1)+(N-2)+...+1
= $N(N-1)/2$ (36 in that case)

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Typical repeatabilities (60 points => ~1800 bsl)



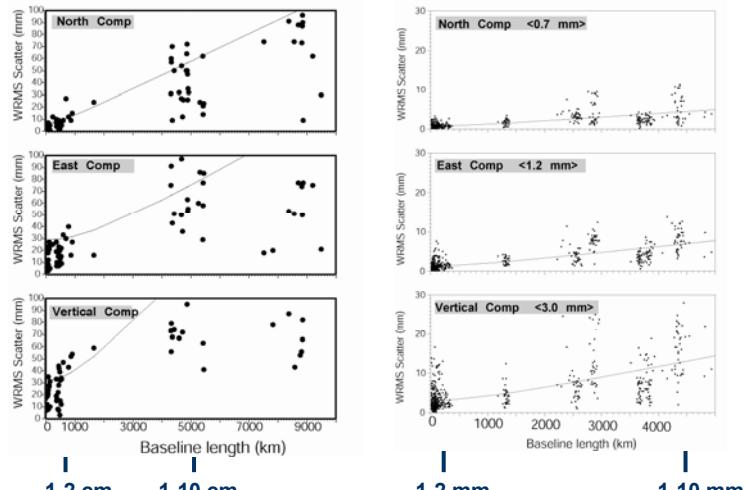
Repeatabilities are much larger than formal uncertainties !

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1 order of magnitude Improvement over a decade

Afar 1991 -bias fixed- pgga orbits Chili 2006 - biases fixed - IGS orbits

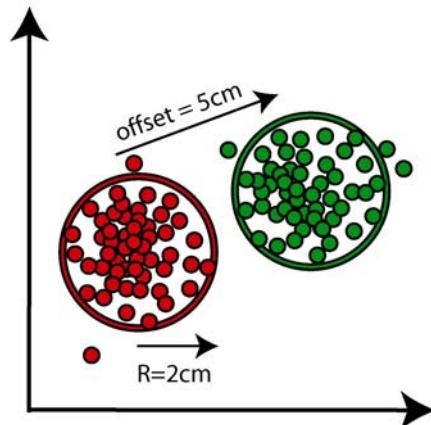


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Accuracy vs. precision (1)



Fix point :
measure 1 hour every 30 s

=> 120 positions

with dispersion $\sim \pm 2\text{ cm}$

5 hours later, measure again
1 hour at the same location

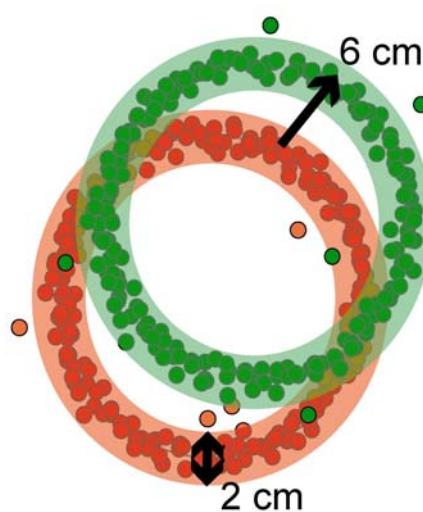
=> Same dispersion but
constant offset of 5 cm

Precision = 2 cm

Accuracy = 5 cm

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Accuracy vs. precision (2)



Measure path, 1 point every
10s

=> 1 circle with 50 points

10 circles describe runabout
with dispersion $\sim 2\text{ cm}$

Next day, measure again

=> Same figure but constant
offset of 6 cm

Precision = 2 cm

Accuracy = 6 cm

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Mapping in a reference frame (sketch)

Measure short lines with a good precision is easy.

⇒

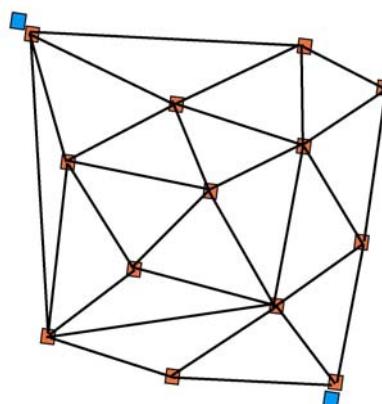
Quantify deformation within a small network is easy

BUT

Determine the motion of the whole network with respect to the rest of the world is more difficult

In order do do this...

We need to know how to map into a **reference frame**

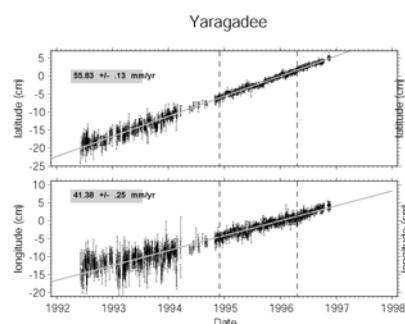


⇒ **compute Helmert transformations**

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Mapping in a reference frame (1)



Constraining campaign positions (and or velocities) to long term positions (and or velocities) works fine ...

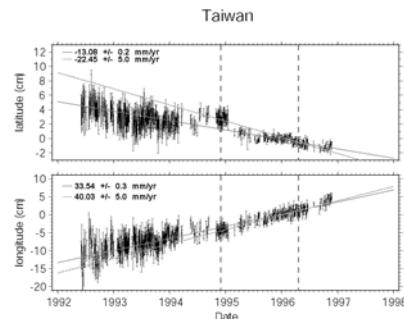
... when station displacement is constant with time

if the station motion is **linear** with time, then estimating the velocity on any time span will give the same value

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Mapping in a reference frame (2)



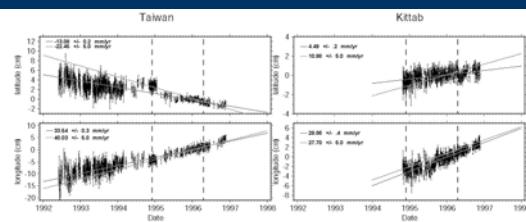
Constraining campaign positions (and or velocities) to long term positions (and or velocities) **does not work**
...
... when station displacement is **not** constant with time

if the station motion is **not linear** with time, then estimating the velocity on **different** time span will give **different** values

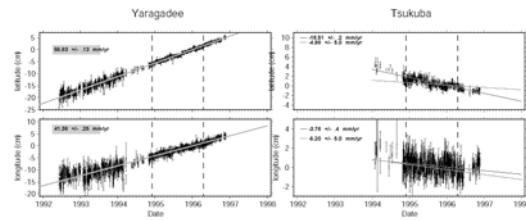
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Mapping in a reference frame (3)



some stations are better than others ...



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From position to velocity uncertainty

If one measures position P_1 at time t_1 and P_2 at time t_2 with precision ΔP_1 and ΔP_2 , what is the velocity V and its precision ΔV ?

$$V = (P_2 - P_1) / (t_2 - t_1)$$

$$\Delta V = (\Delta P_2 + \Delta P_1) / (t_2 - t_1)$$

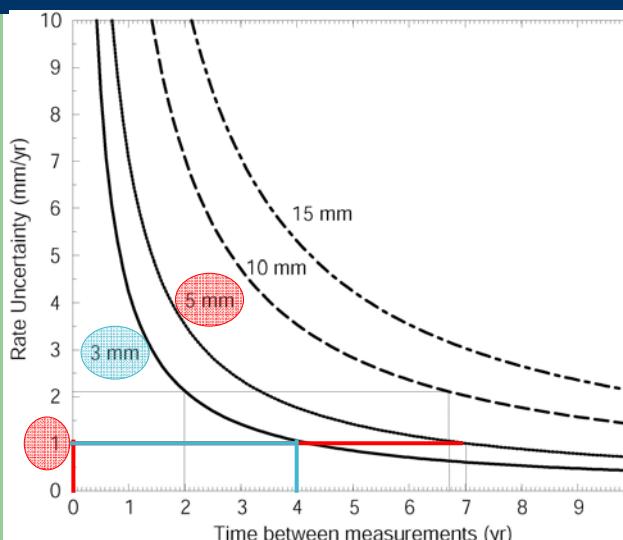
Uncertainties don't add up simply, because sigmas involve probability.

$$\Delta V = [(\Delta P_2)^2 + (\Delta P_1)^2]^{1/2} / (t_2 - t_1)$$

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Velocity uncertainties



It is only needed that we wait a given time between position measurements to detect deformation velocities

The more we wait, the more precise the velocity is

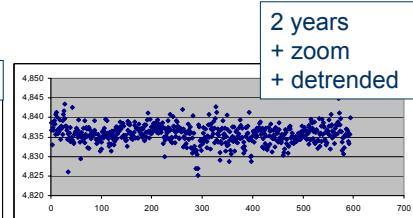
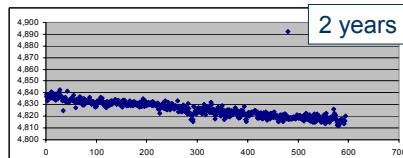
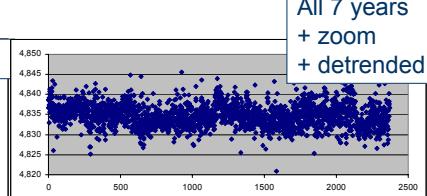
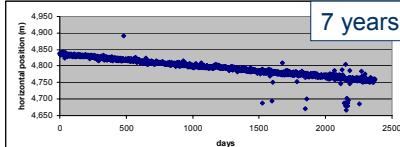
Reasonnable (but wrong) idea:

Geodesy is like good wine: measurements, even (and mostly) bad ones, gain value by aging

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In practice....: BRAZ : 2000->2007

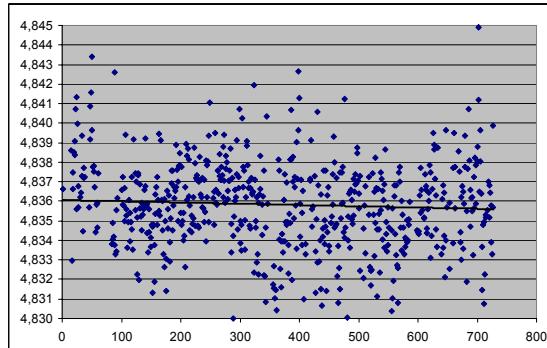


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In practice....: BRAZ : 2000->2007

24 months



month	deviation
1	18,0
2	-7,8
3	-13,2
6	-6,4
12	-1,3
18	-1,0
24	-0,2

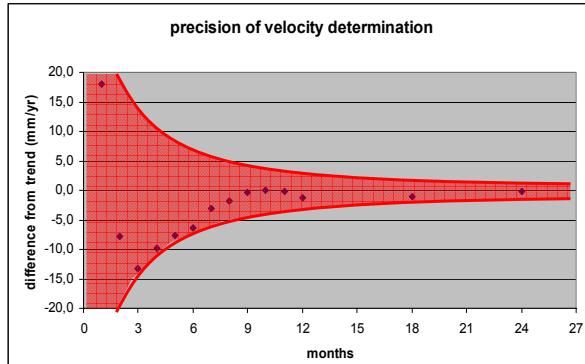
30

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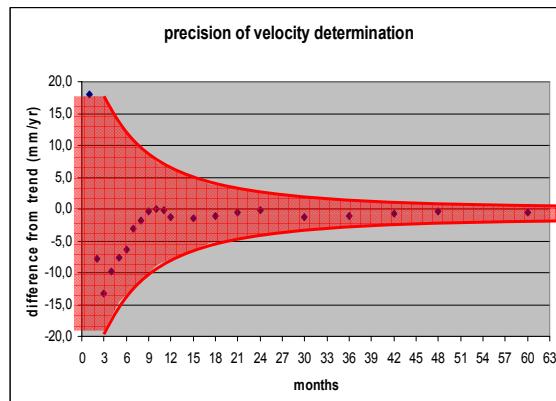
M2 STEP - Module "Outils et méthodes de la géodésie spatiale" - C. Vigny

month	deviation
1	18,0
2	-7,8
3	-13,2
4	-9,9
5	-7,7
6	-6,4
7	-3,1
8	-1,8
9	-0,3
10	0,0
11	-0,1
12	-1,3
18	-1,0
24	-0,2

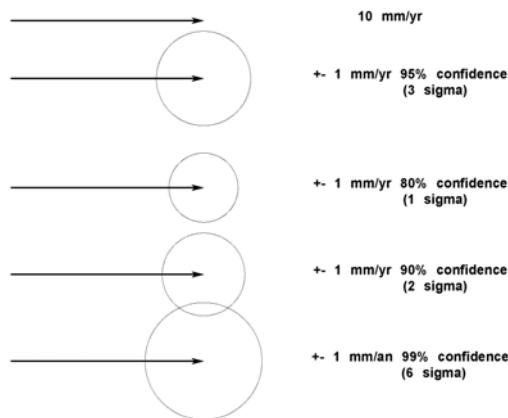


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Velocity ellipses



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Maintenant, que mesure-t-on exactement avec GPS ?

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Campagnes de répétition

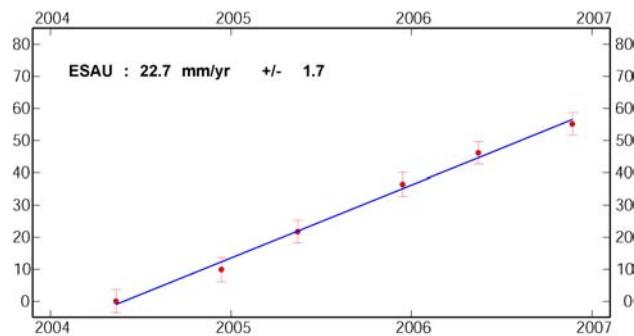


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Repeated Campaigns



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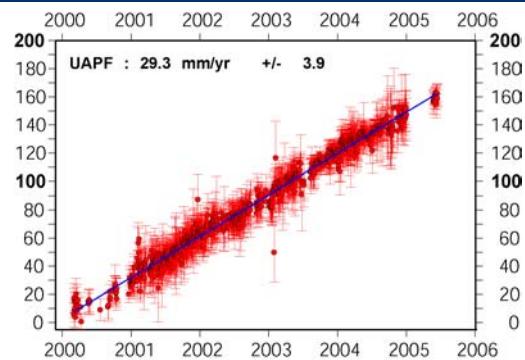
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Stations permanentes

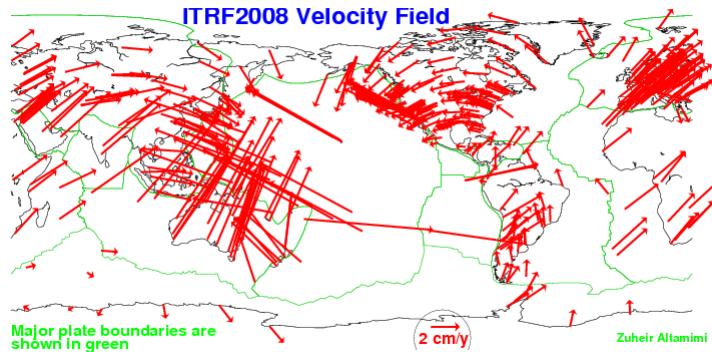




Permanent station



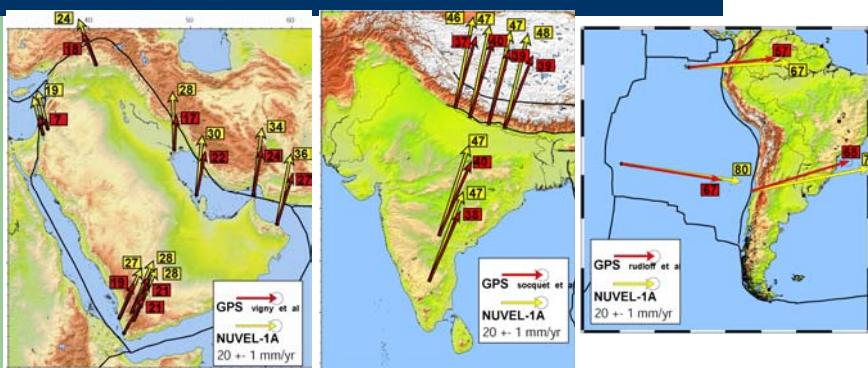
Champ de vitesse global => ITRF



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À grande échelle: les plaques



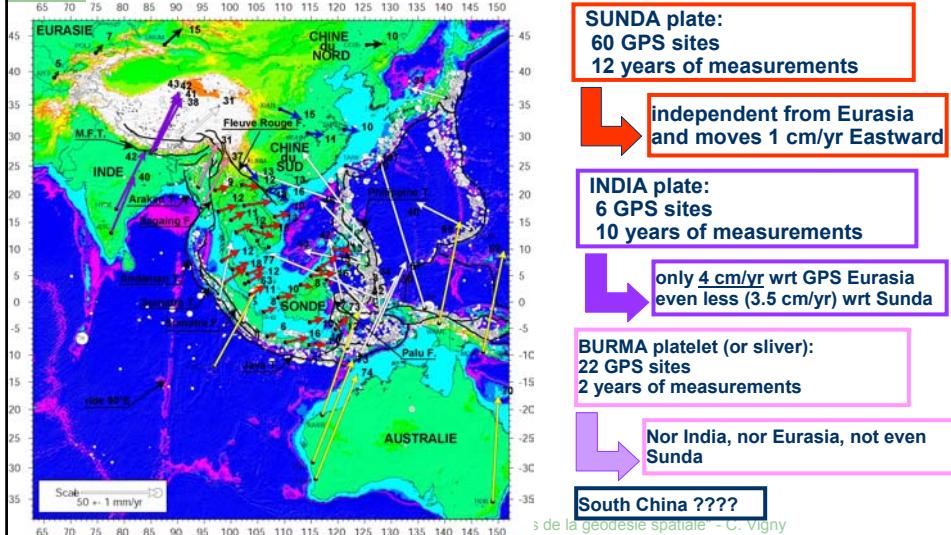
Au premier ordre, la géodésie sur une décennie « colle » bien avec la géologie sur 3 Ma!

=>Les plaques ont donc des déplacements stables sur ce type de durée mais.... GPS trouve Arabie, Inde et Nazca plus lentes actuellement
Arabie et Inde => ralentissement
Nazca => incertitude ?

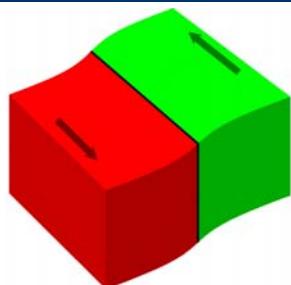
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I'Asie du Sud-Est, une zone compliquée

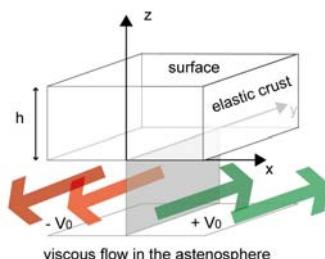


La déformation autour des frontières de plaques



Le concept: déformation élastique à cause du frottement

⇒ La solution analytique:
une courbe en arctangente

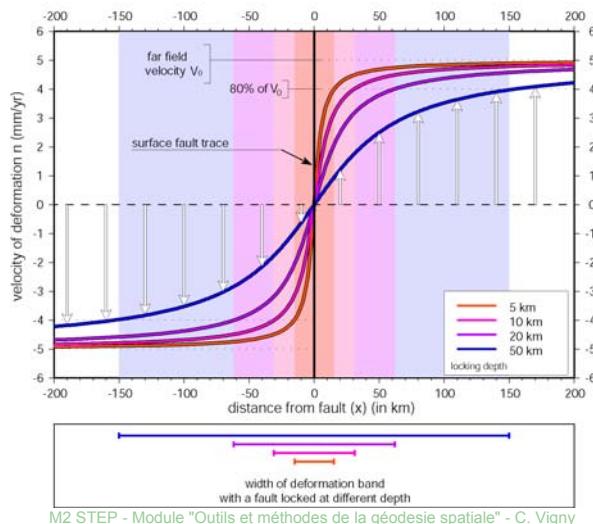


Le modèle: le flux mantellique localise la déformation sous une plaque élastique

$$U_y = 2 \cdot V_0 / \Pi \arctan(x/h)$$

Arctang profiles

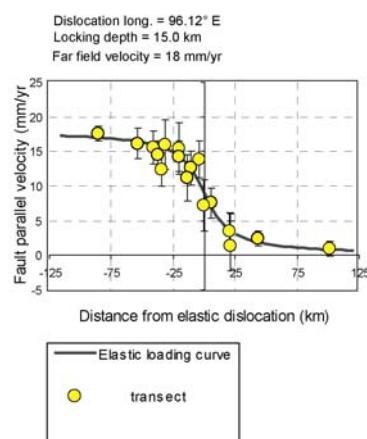
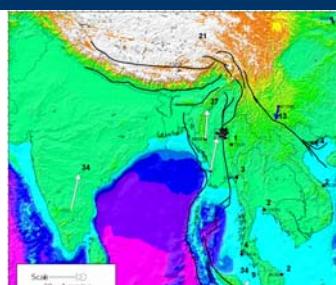
$$U_y = 2 \cdot V_0 / \Pi \arctan(x/h)$$



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La faille de Sagaing en Birmanie



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Remarques importantes (1):

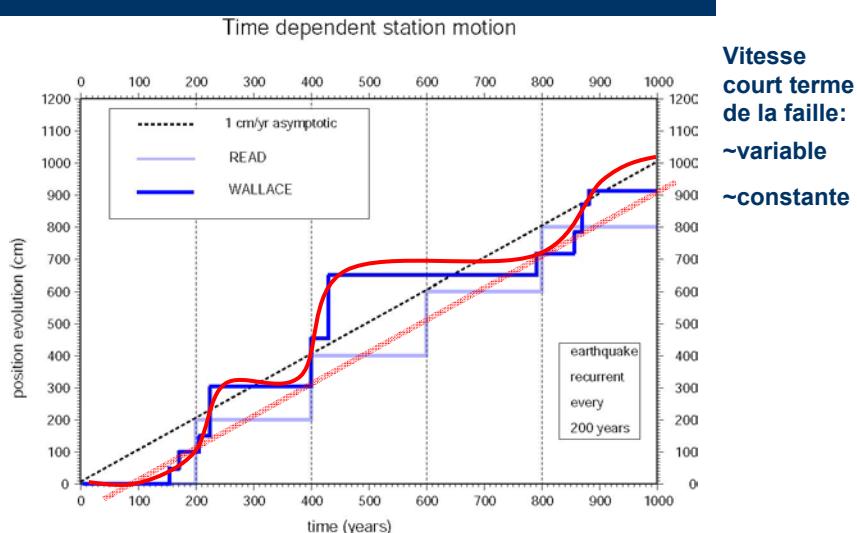
- Quand les stations sont bien à l'intérieur des plaques (loin des marges qui se déforment) GPS mesure la **tectonique des plaques**
- Quand les stations sont dans les marges qui se déforment, **GPS ne mesure pas la vitesse des failles** (qui est en général zéro si la faille est bloquée), mais la **déformation élastique qui s'accumule** en réponse au blocage. C'est au travers d'un modèle (Okada par exemple) que l'on retrouve La vitesse limite (càd la tectonique) et l'épaisseur élastique, mais toujours pas la vitesse de la faille, concept qui ne veut pas dire grand-chose.

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Le cycle sismique:

Shimazaki et Nakata, 1980
Wallace, 1987



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Remarques importantes (2):

- La tectonique active basée sur le décalage de marqueurs de part et d'autre d'une faille ne mesure donc pas forcément non plus la « vitesse de la faille » (et donc la tectonique des plaques).

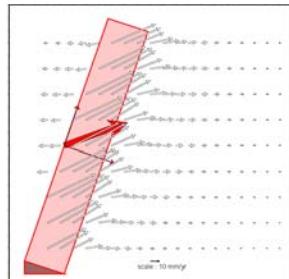
Cela dépend du nombre de cycles sismiques moyennés dans la mesure

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Déformation intersismique aux frontières de plaques

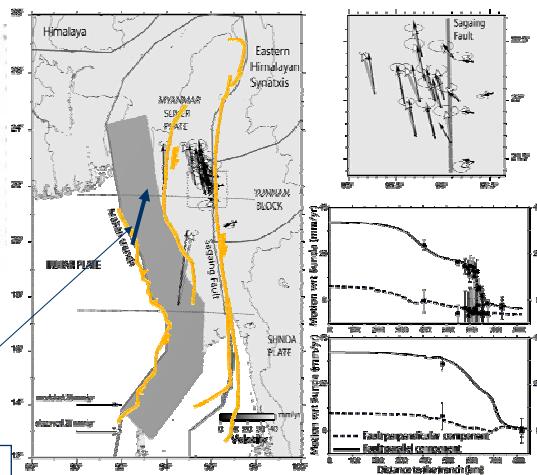
Okada back-slip model



Best fit dislocation:

2 – 2.5 cm/yr

N 30°-35°



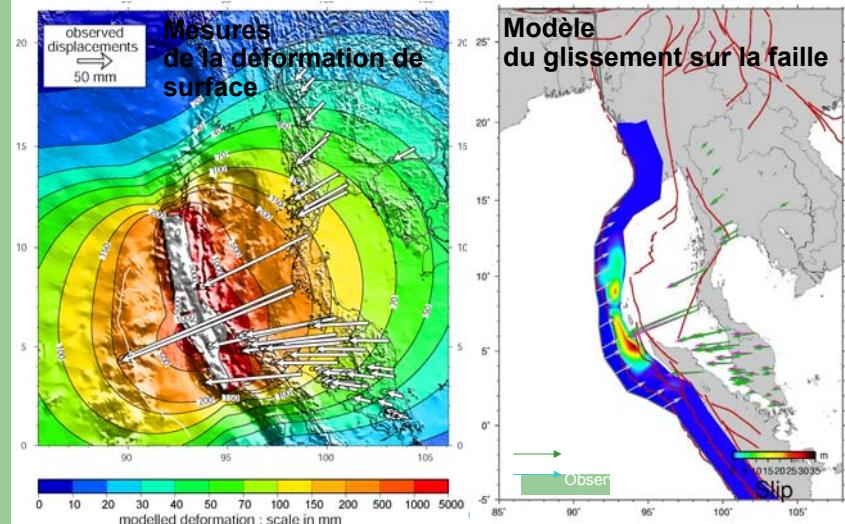
⇒ Potentiel pour un séisme

50 Mw ~8.5 tous les siècles

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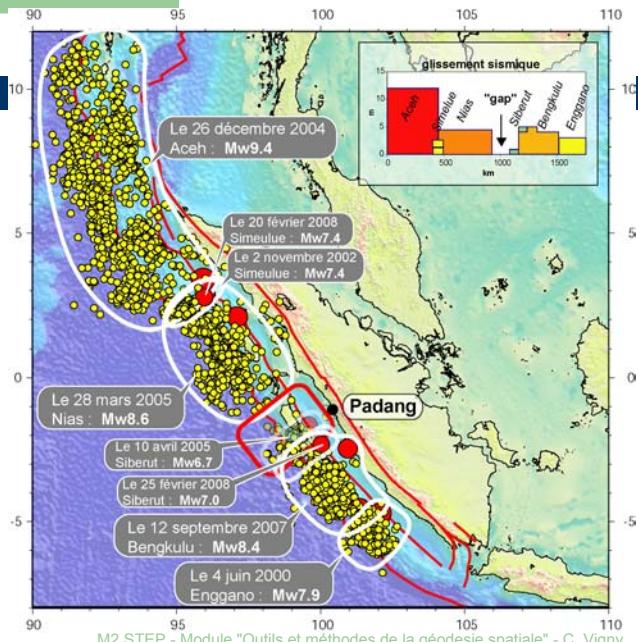
Mesure d'un séisme: Sumatra 2004 – Mw 9.2

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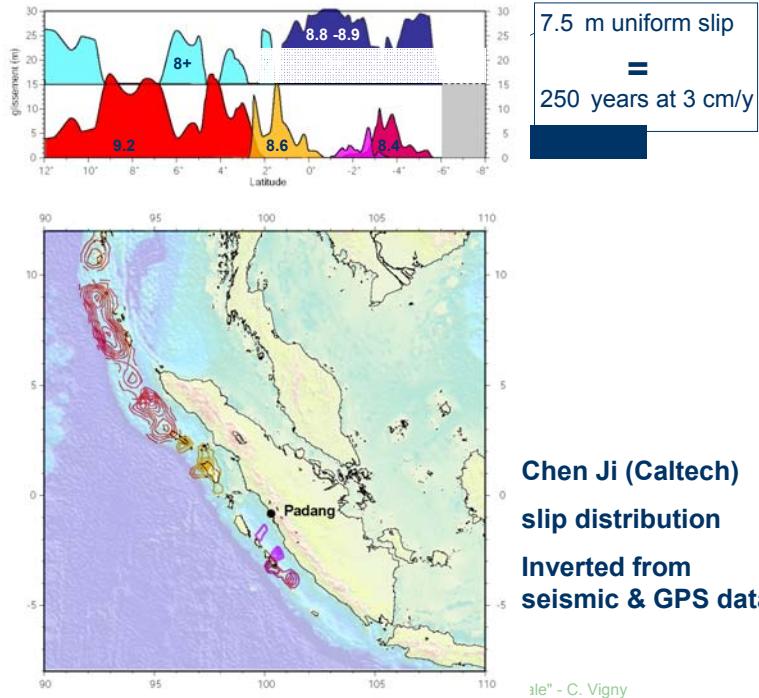


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Les Séismes de Sumatra

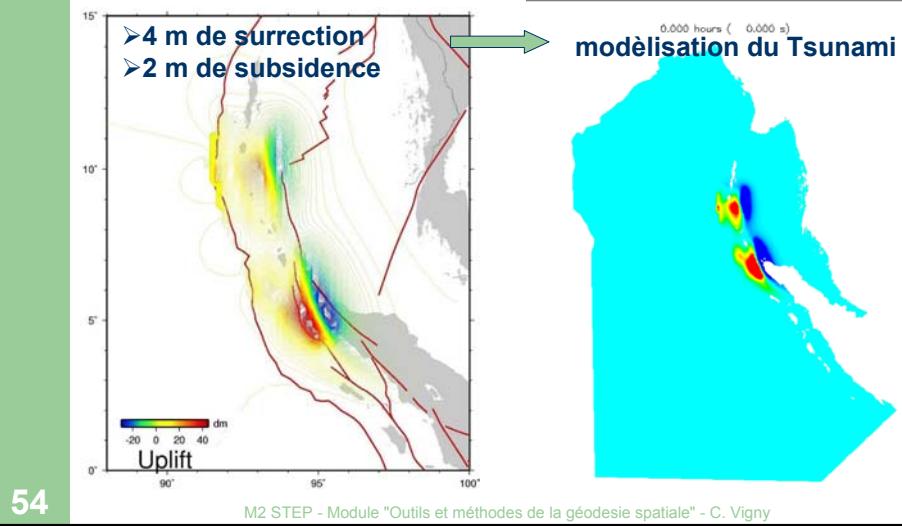


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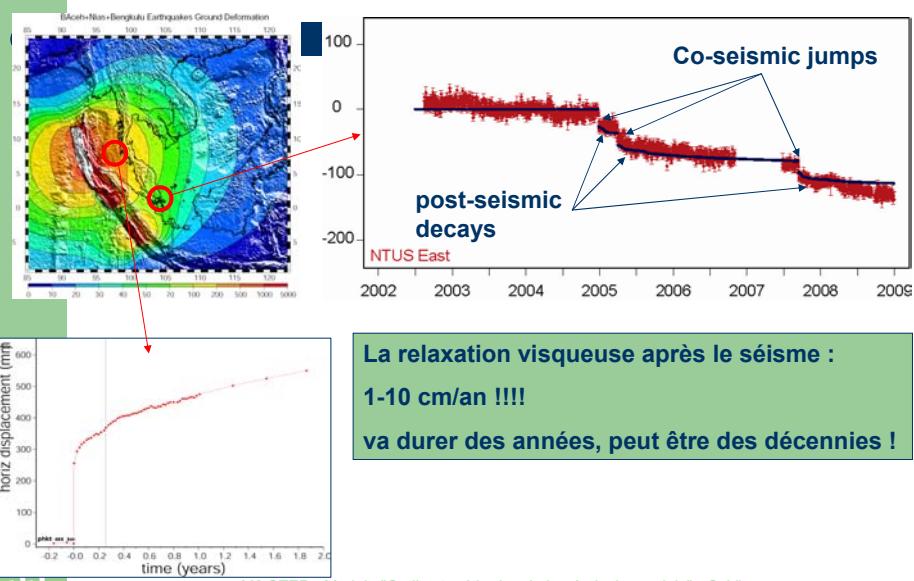
"...ale" - C. Vigny

Les mouvements verticaux prédis par le modèle



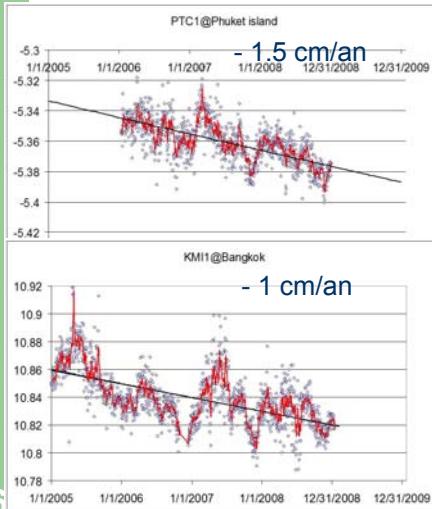
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Déformations Post-sismique

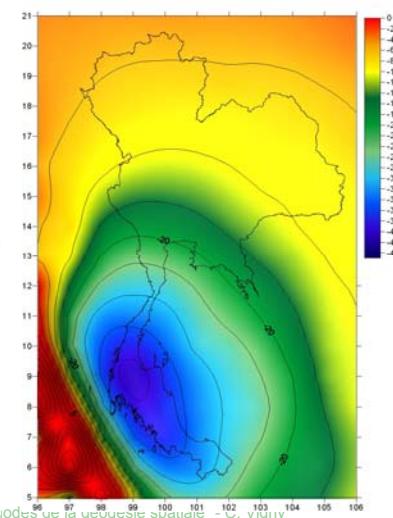


Thailande: Subsidence déclenchée par le séisme de Sumatra

Les Mesures

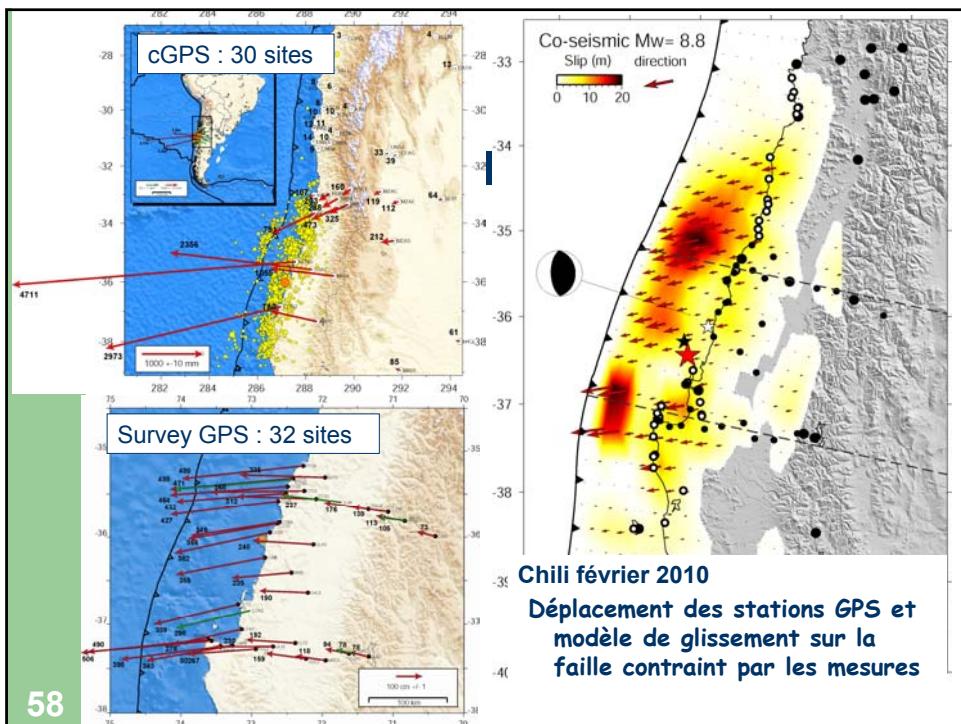


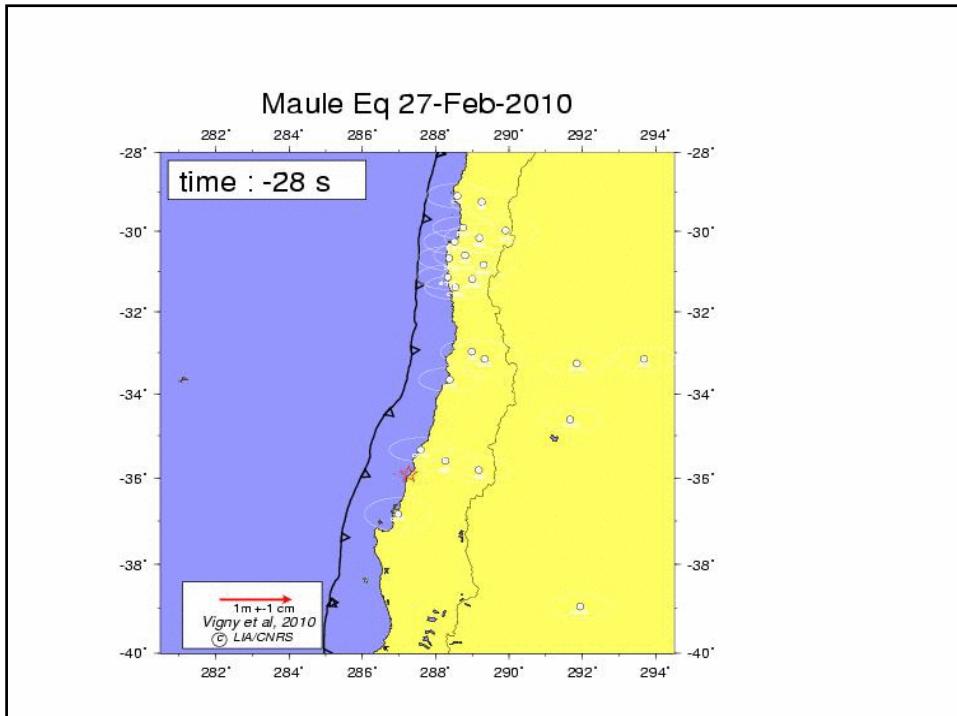
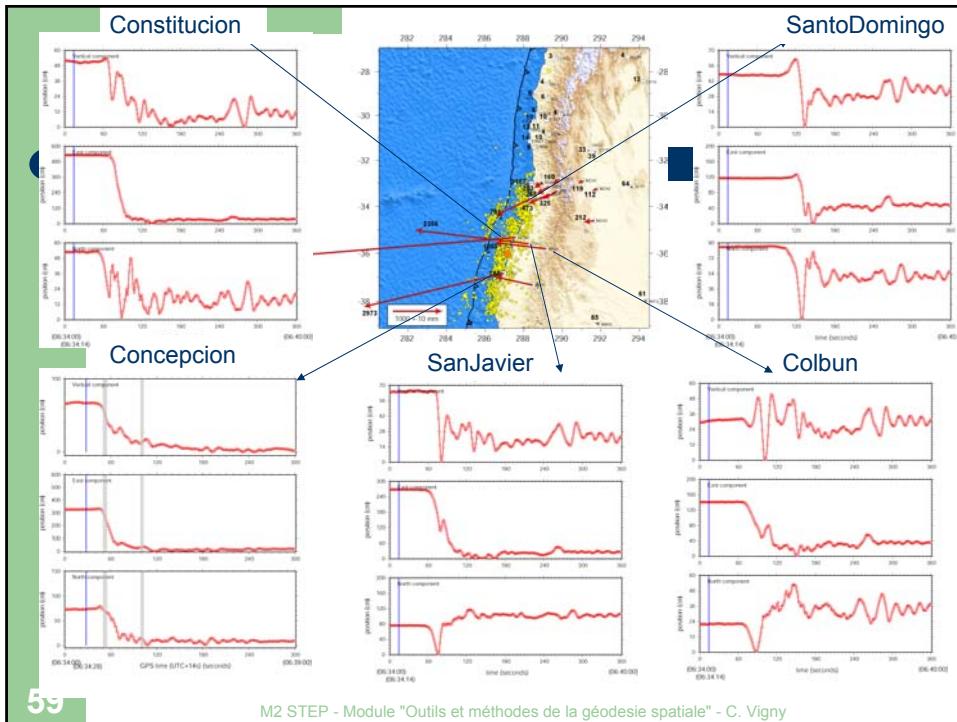
Le Modèle

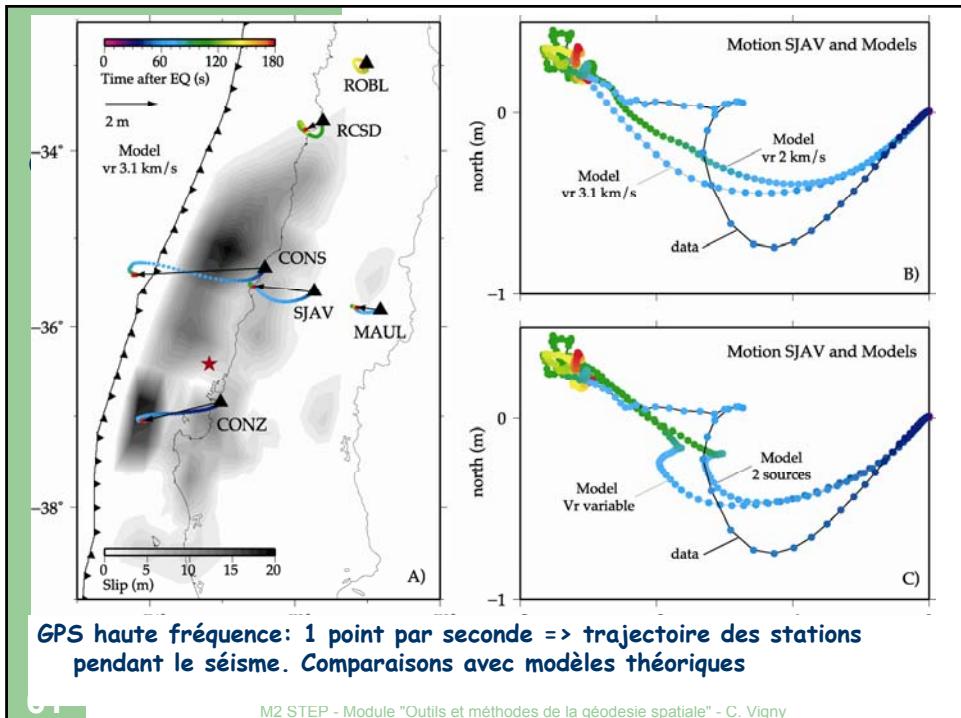




Applications de la géodésie spatiale - C. Vigny







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