

Subduction et Géodésie GPS

Christophe VIGNY

vigny@geologie.ens.fr

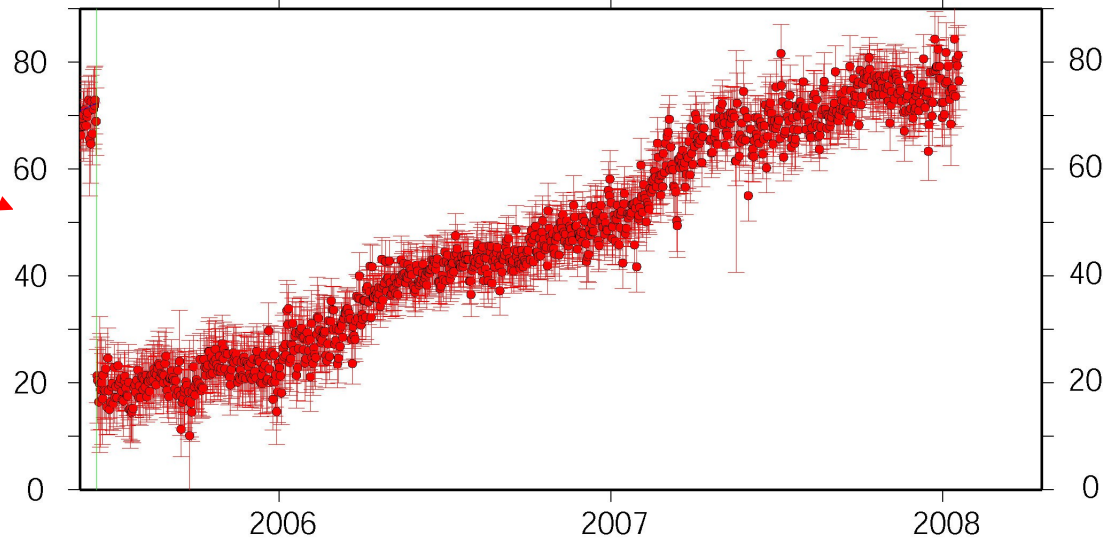
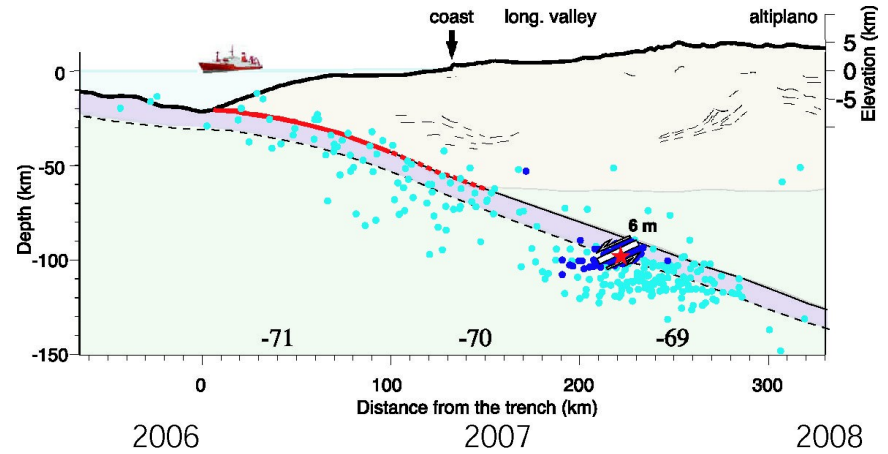
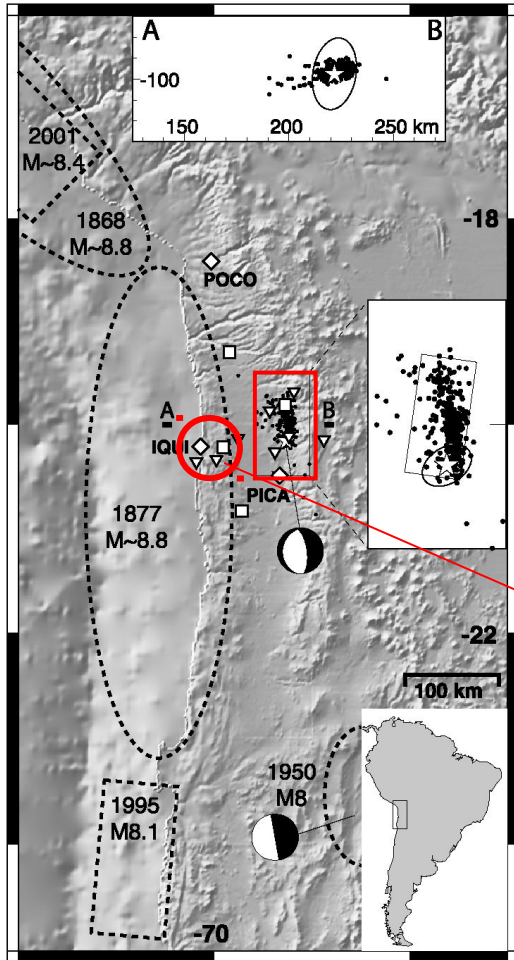
<http://www.geologie.ens.fr/~vigny>

UAPF after Tarapaca Eq. Mw7.7 13-june-2005 (slab pull)

Peyrat et al., *GRL*, 2006

L22308

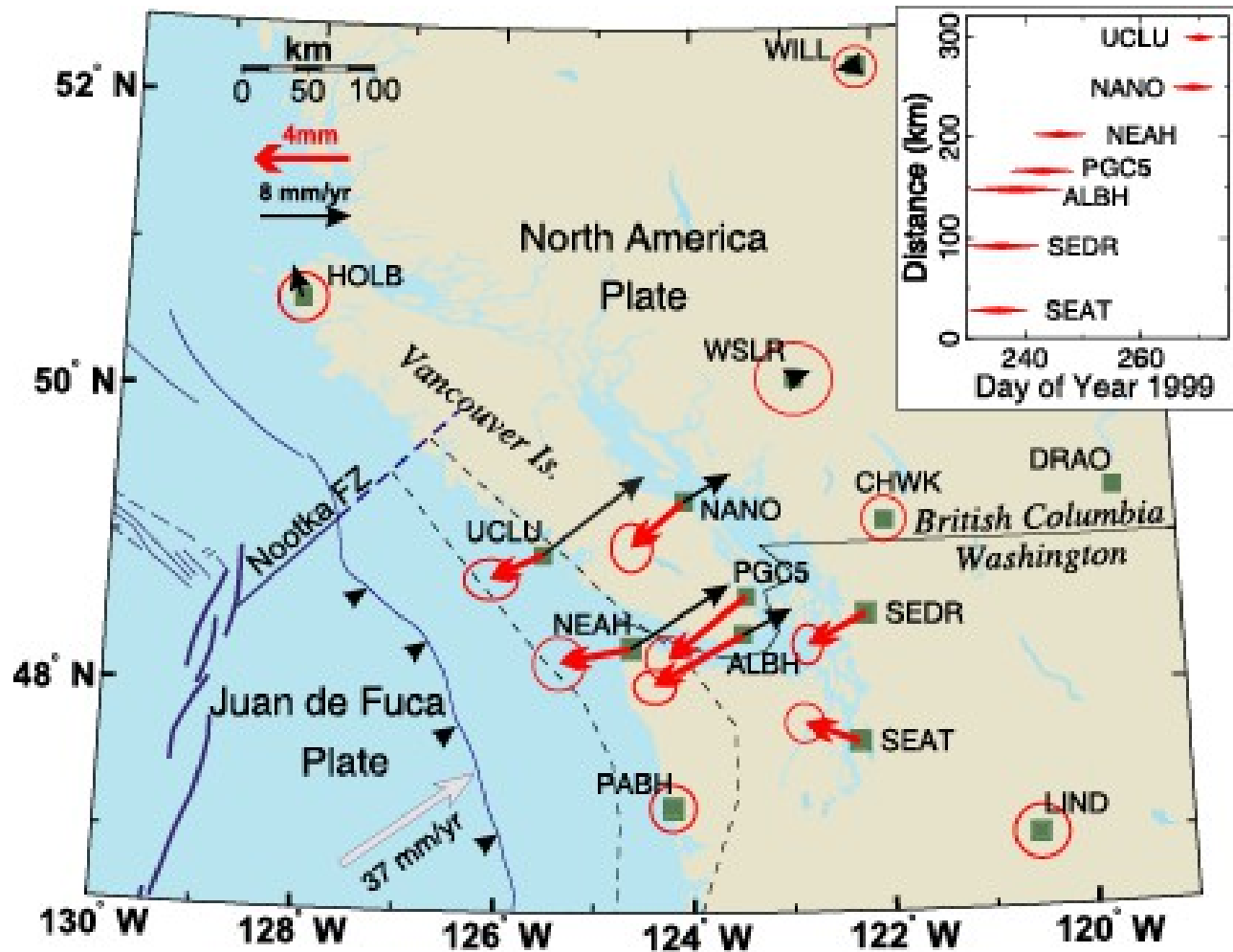
PEYRA T ET AL.: 2005



Short term transients

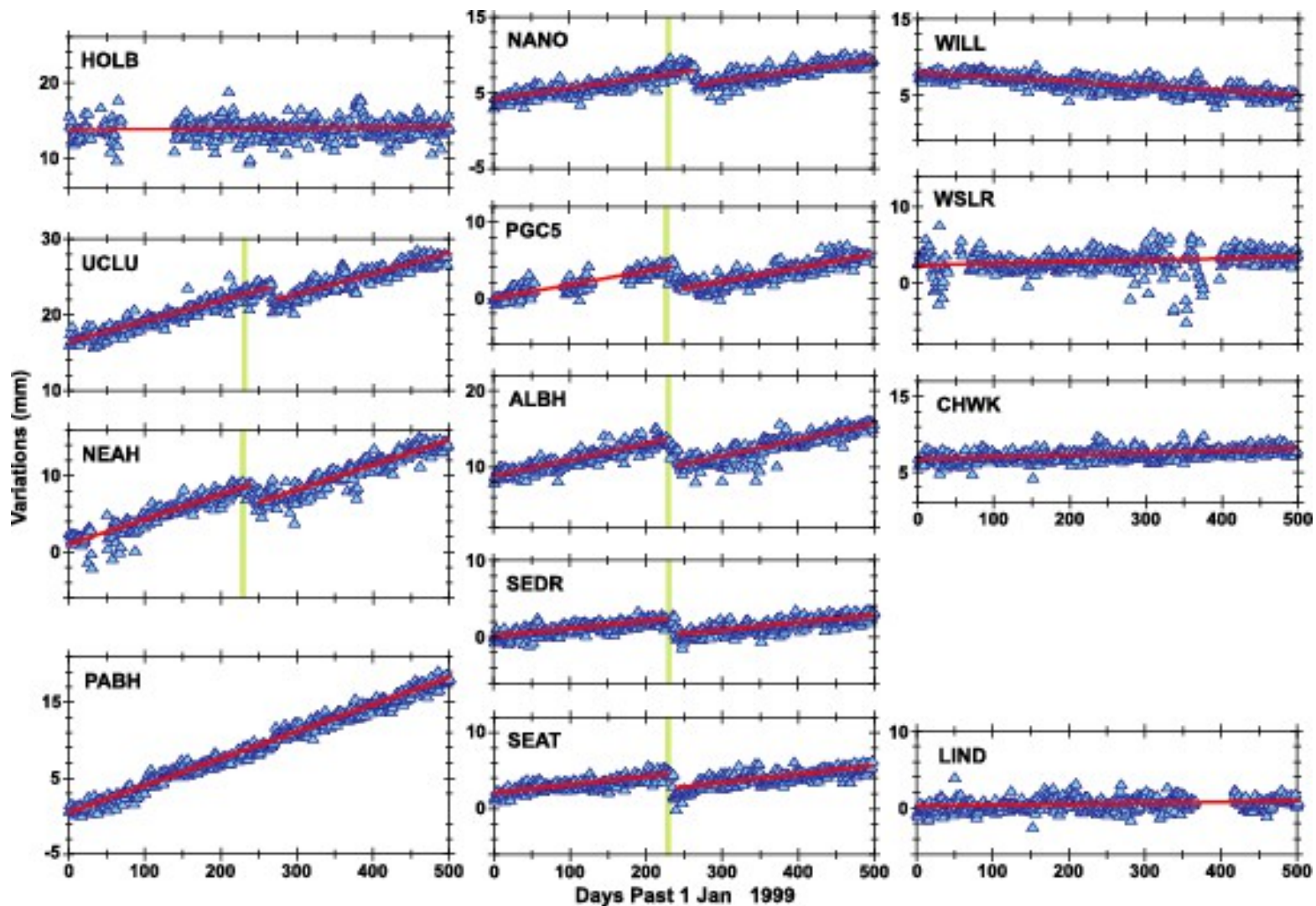
Silent slip on Cascadian subduction zone

Dragert et al., Science, 292, May 2001



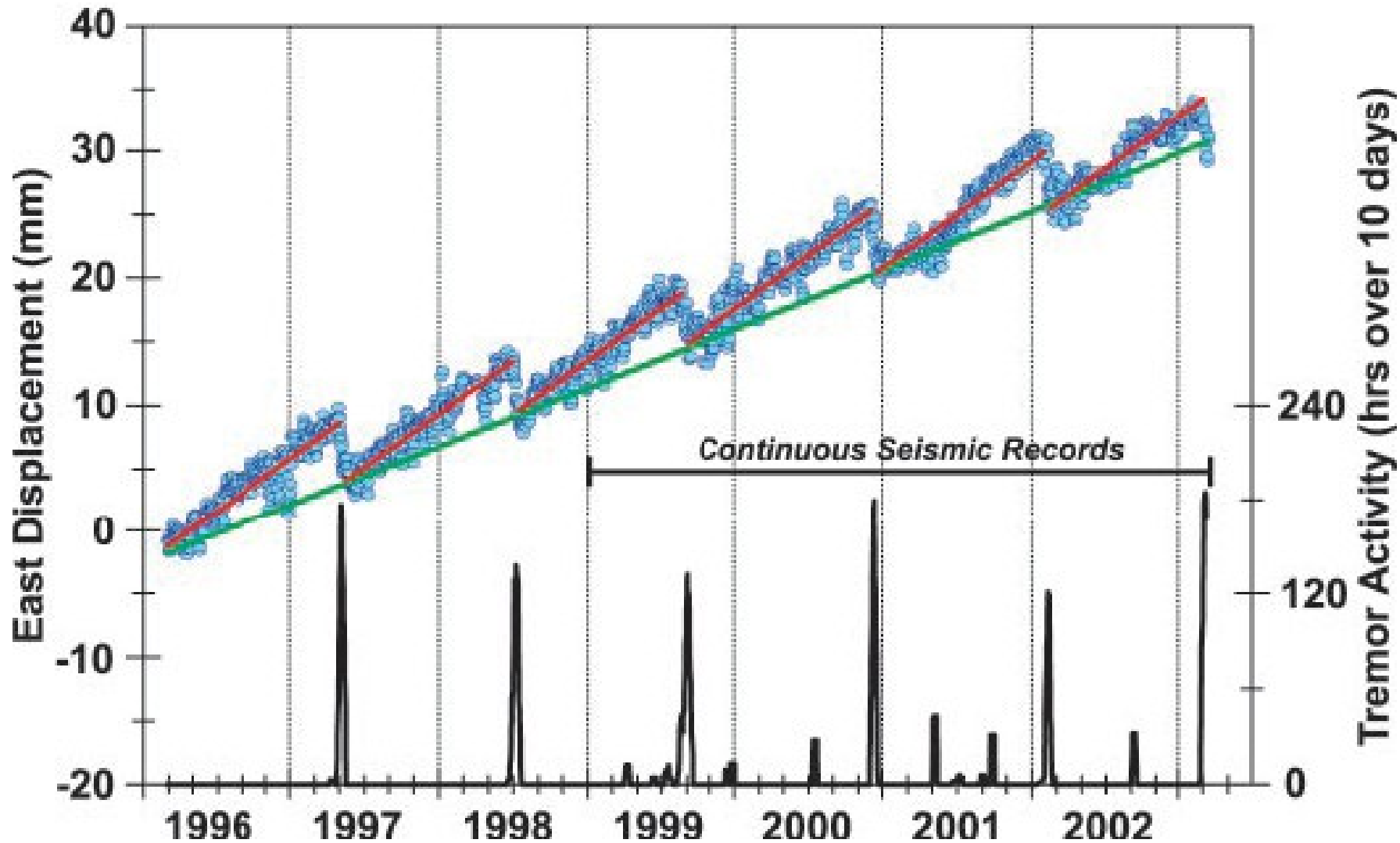
Jump in GPS stations time series

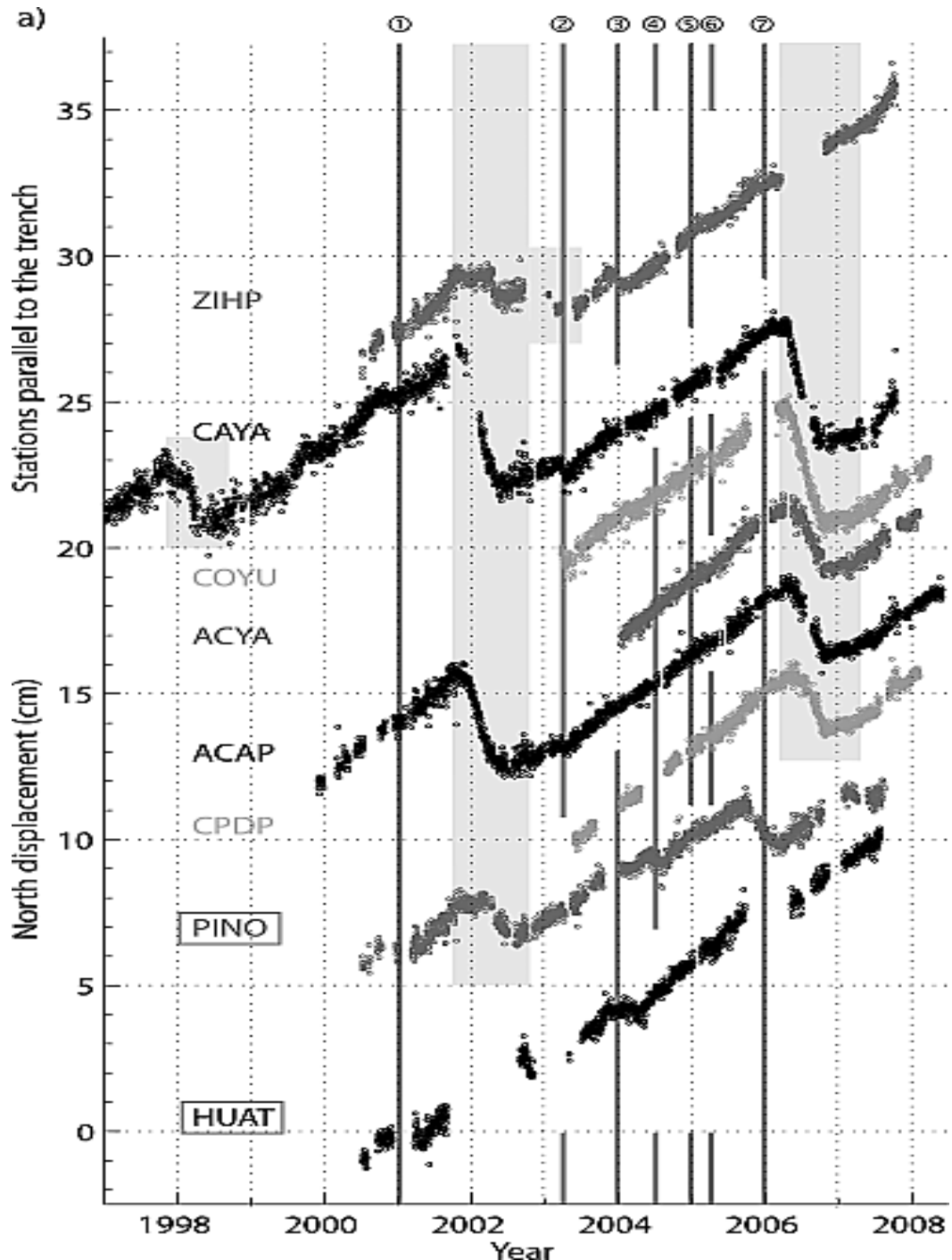
Dragert et al., Science, 292, May 2001



Repeated features, related to tremor

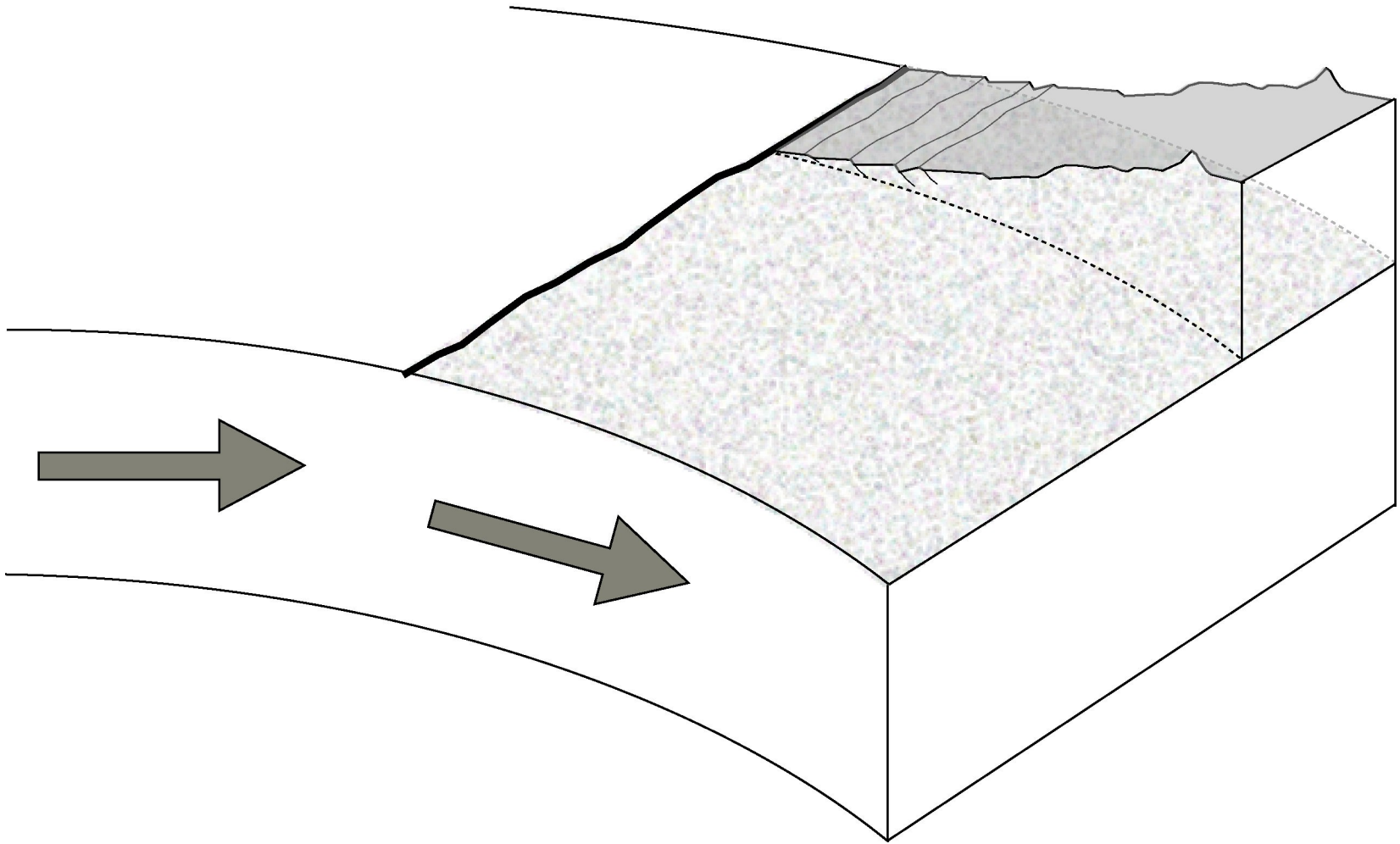
Rogers and Draggert, Science, 300, June 2003



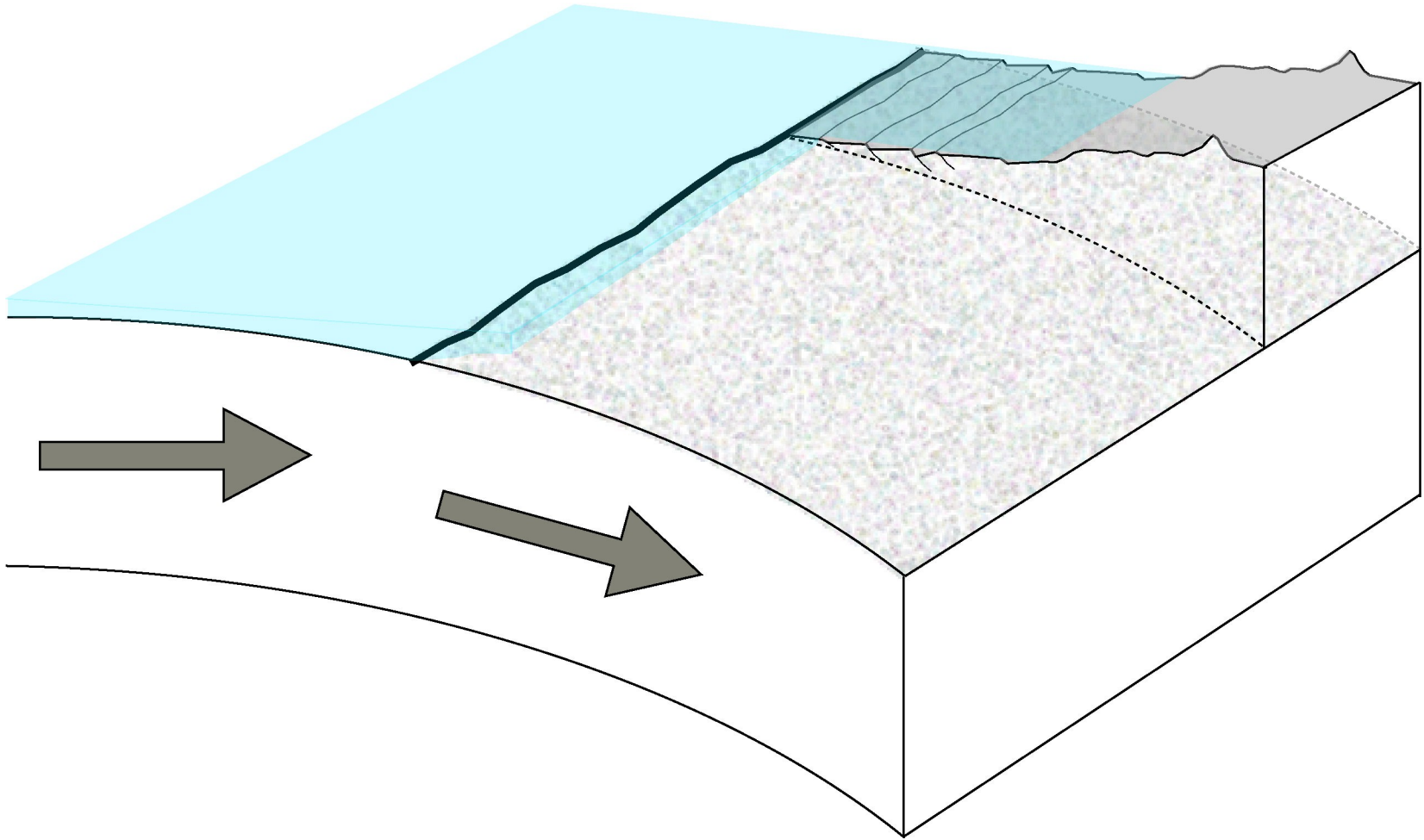


Au Mexique
aussi

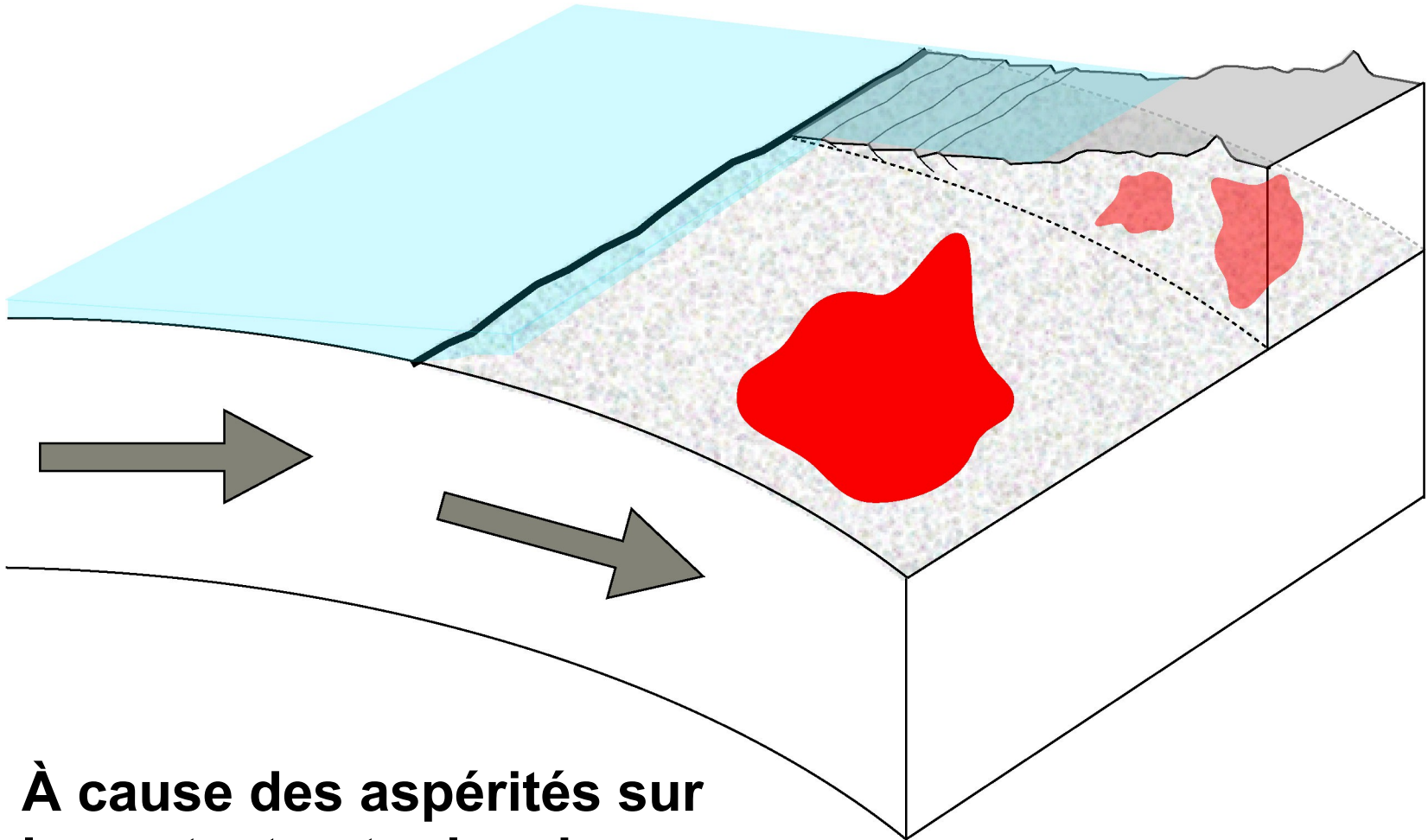
Pourquoi y a-t-il des séismes ?



Pourquoi y a-t-il des séismes ?

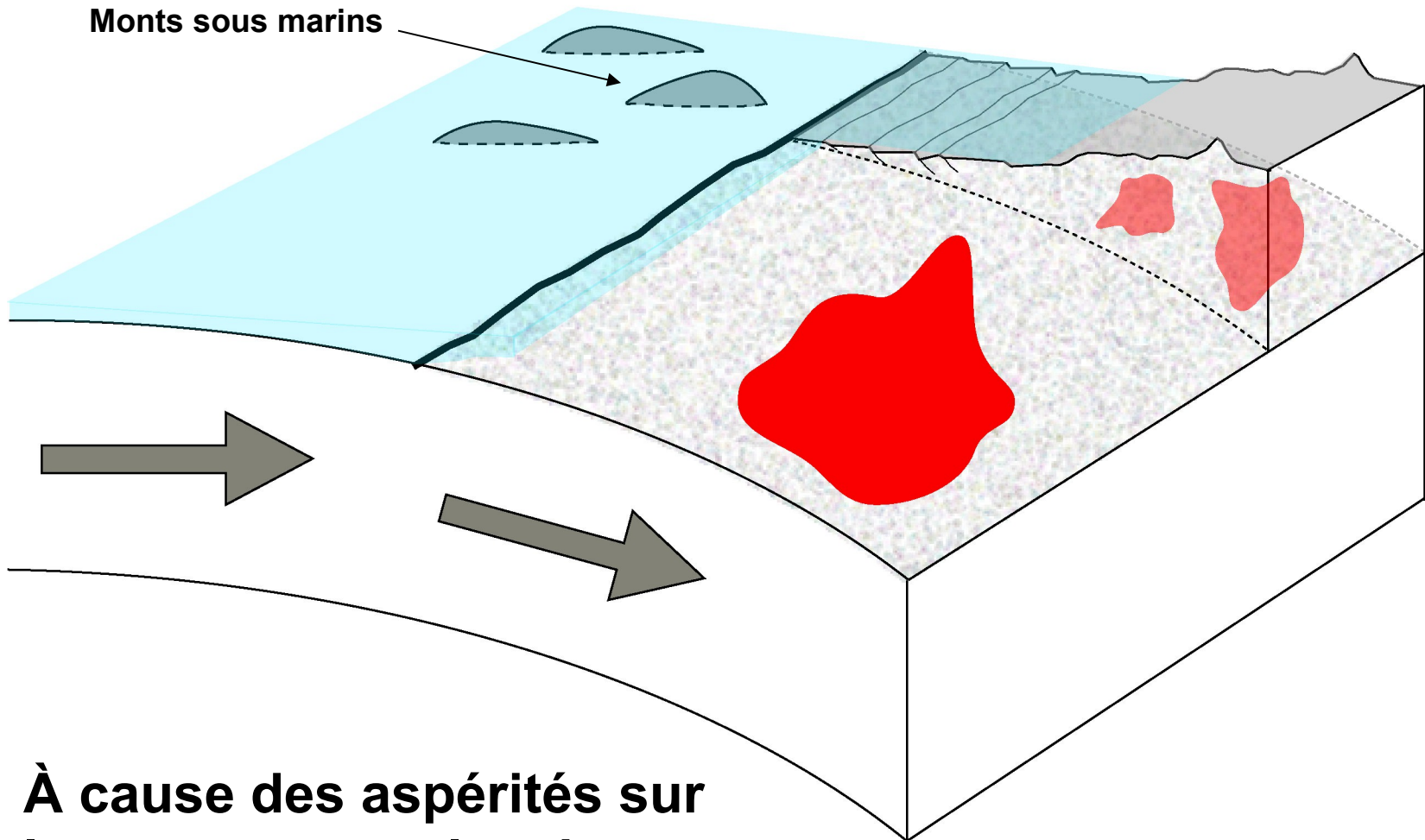


Pourquoi y a-t-il des séismes ?



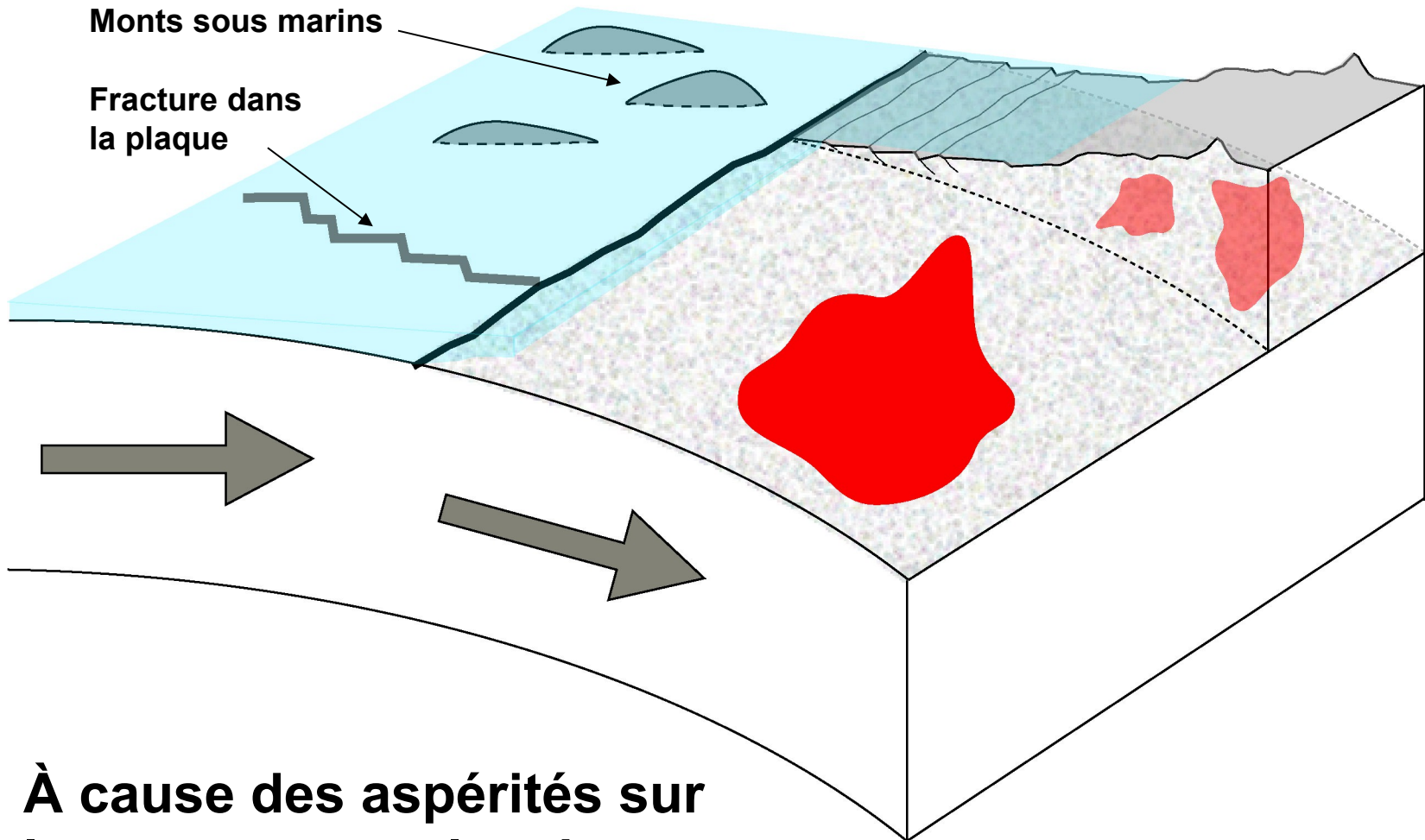
**À cause des aspérités sur
le contact entre les deux
plaques**

Pourquoi y a-t-il des séismes ?



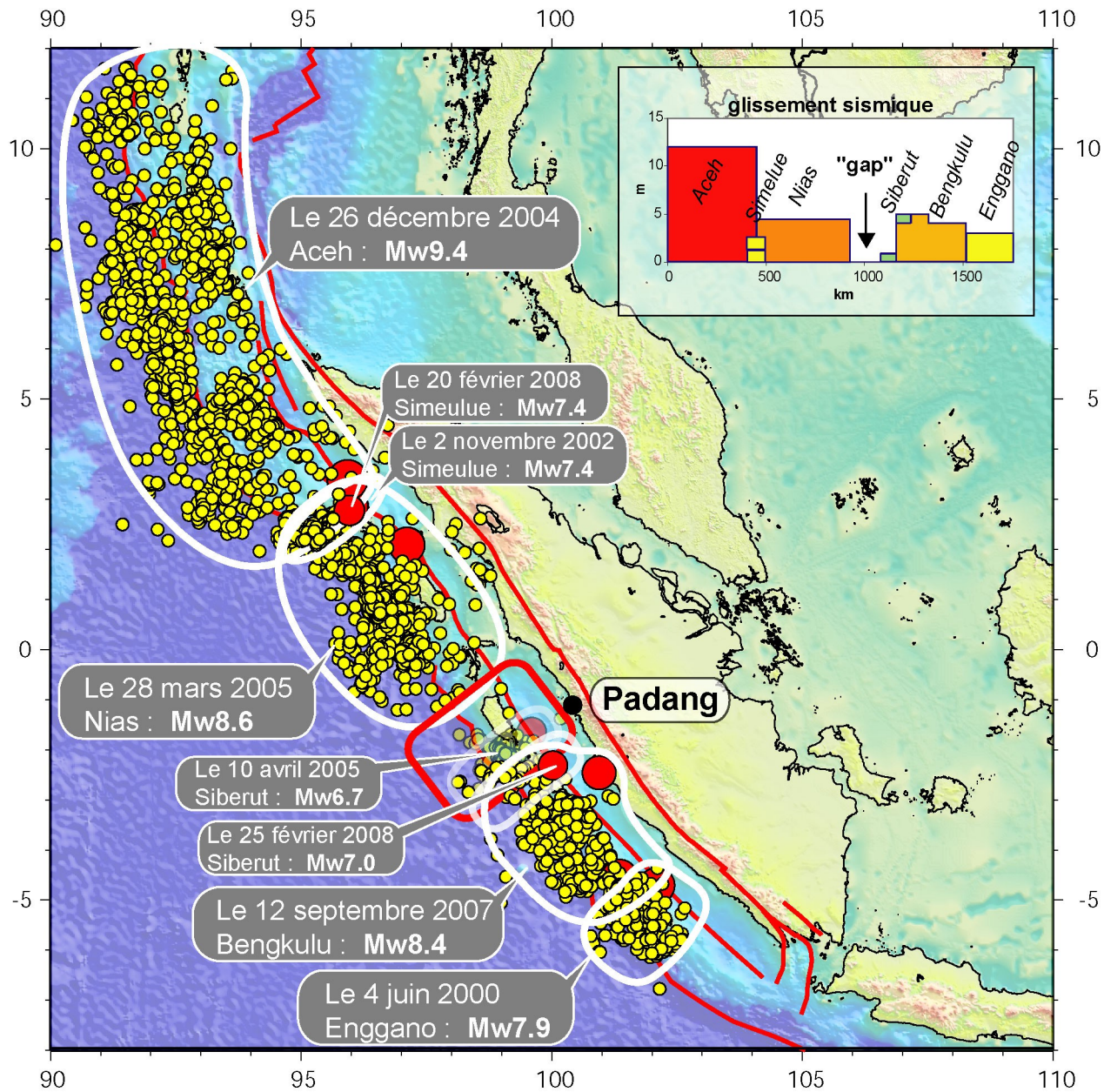
**À cause des aspérités sur
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Pourquoi y a-t-il des séismes ?

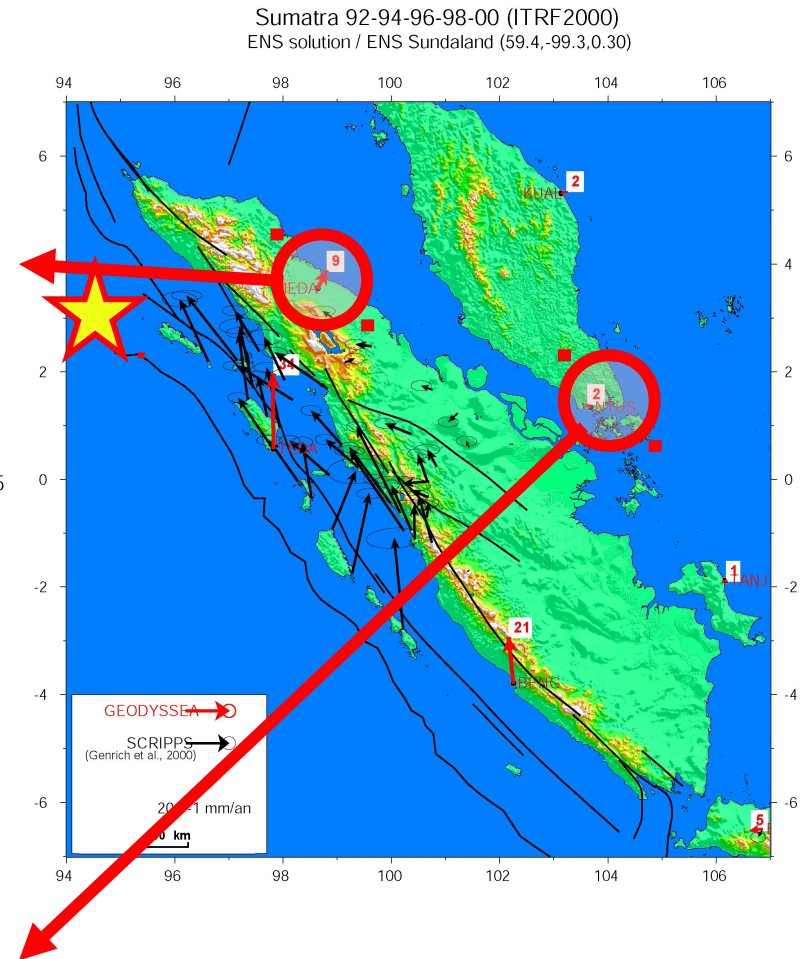
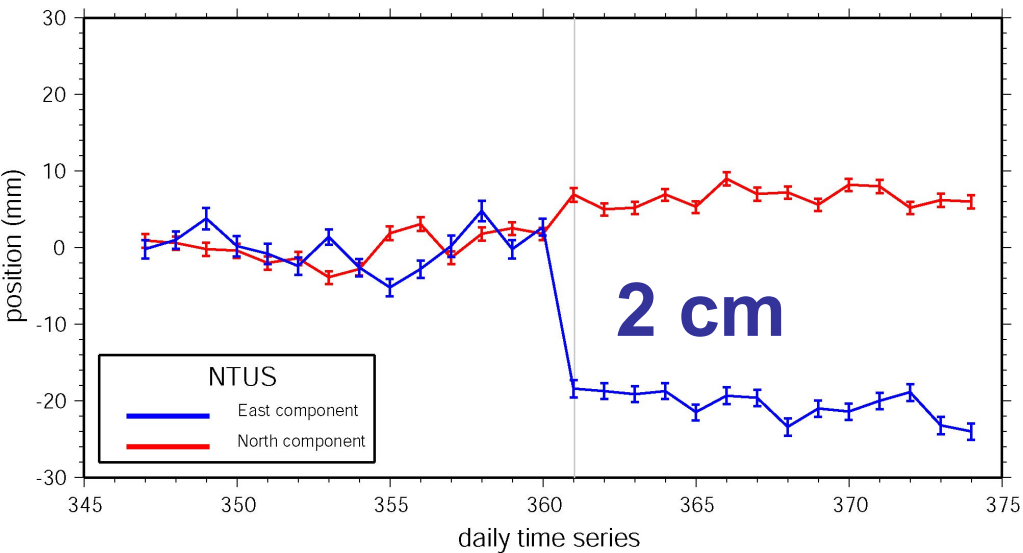
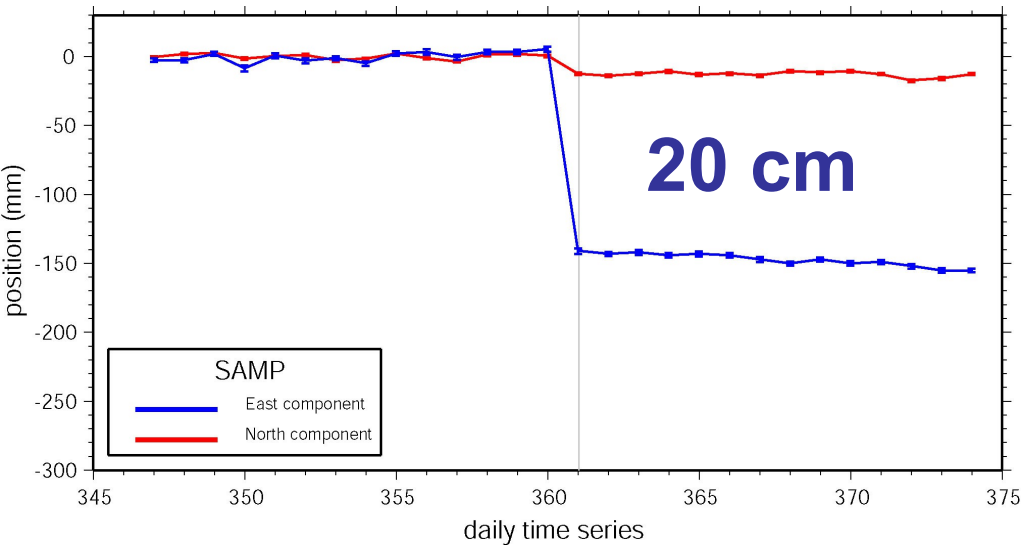


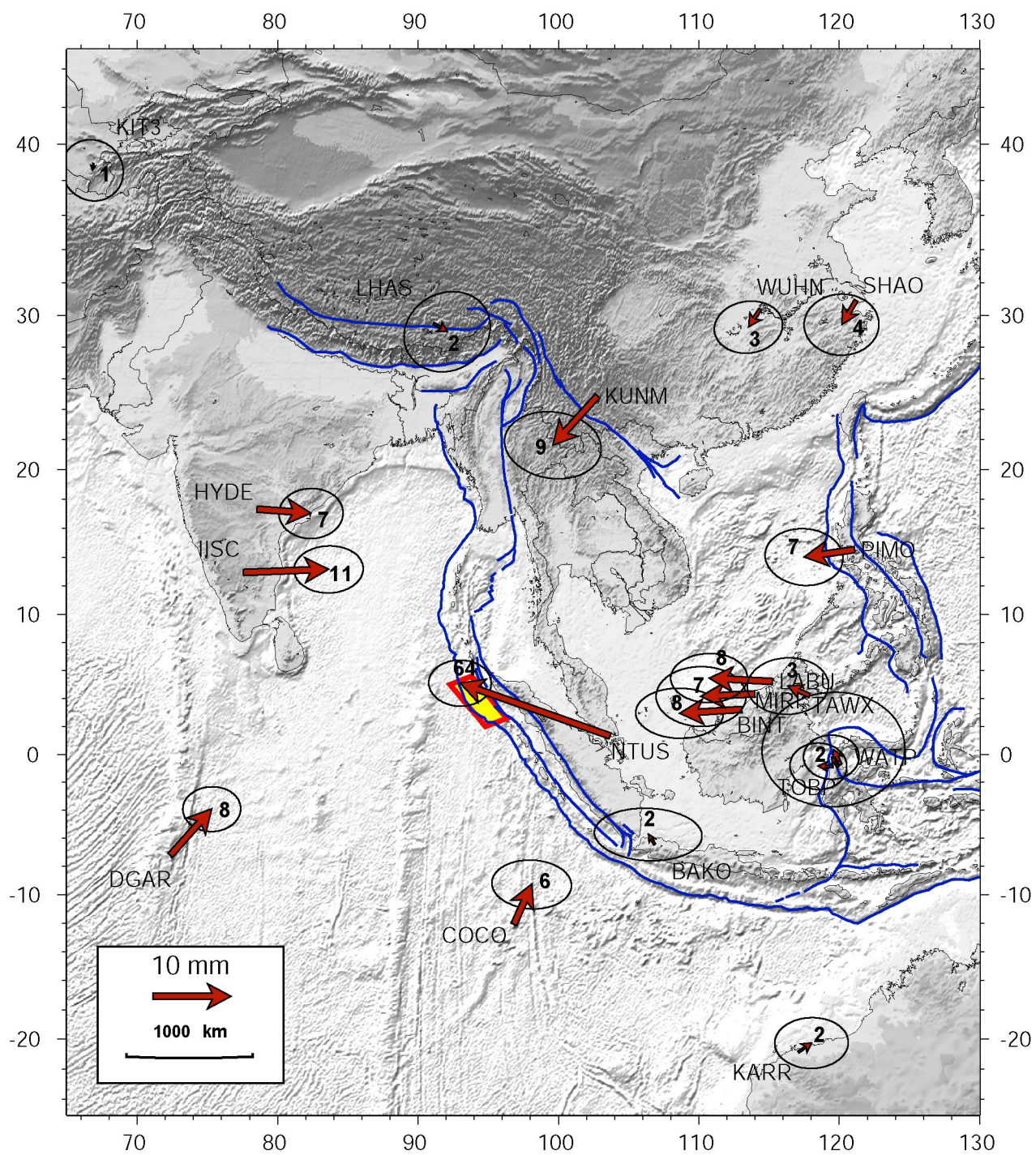
**À cause des aspérités sur
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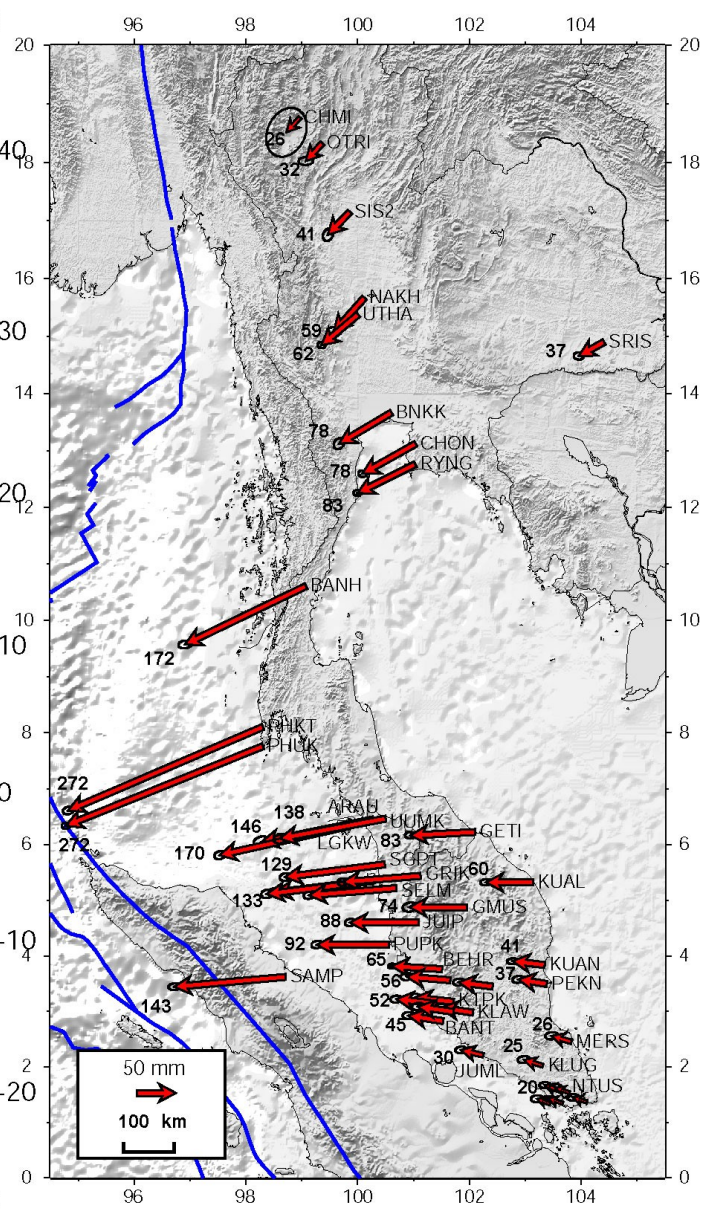
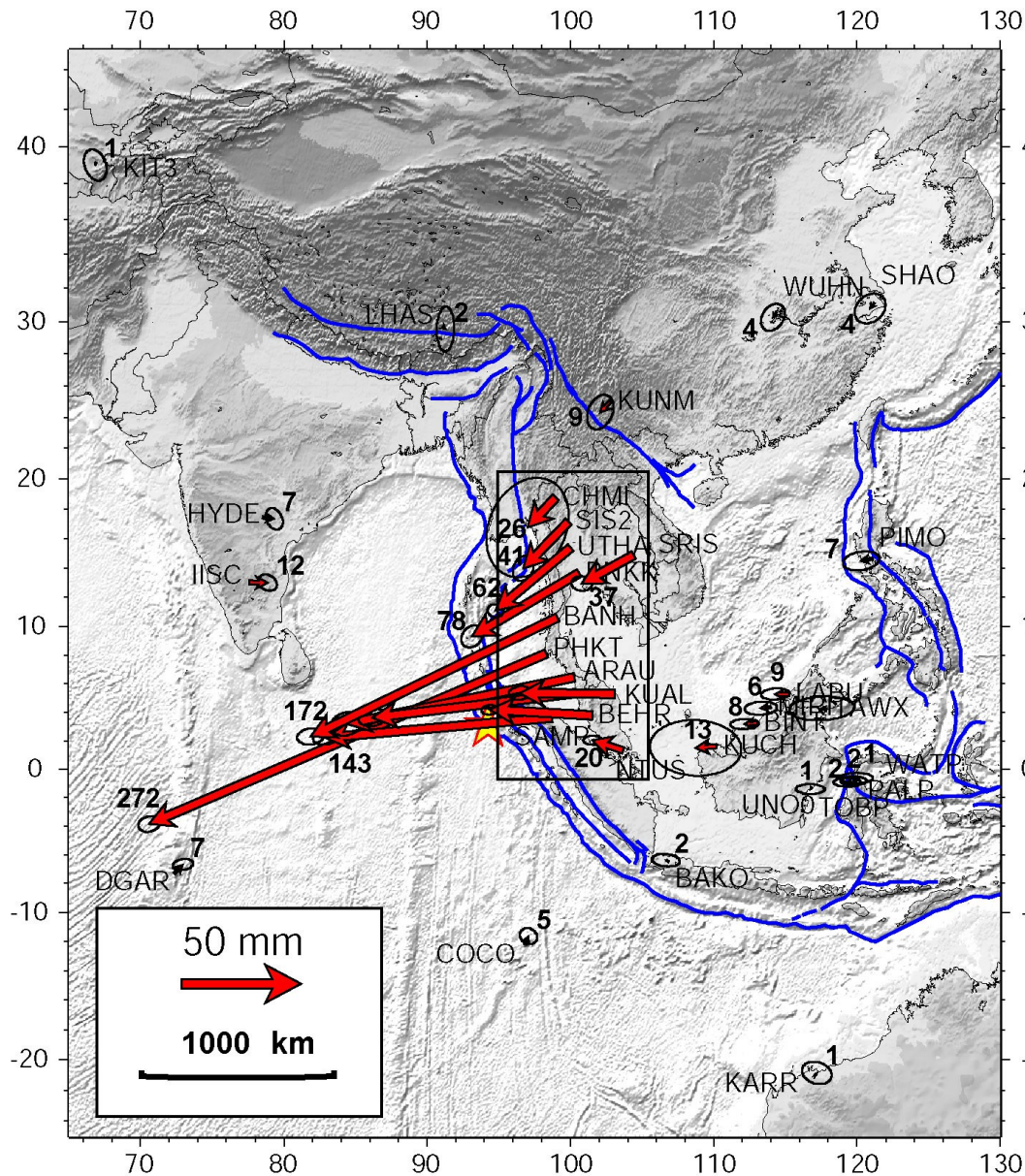
Le(s) séisme(s) de Sumatra

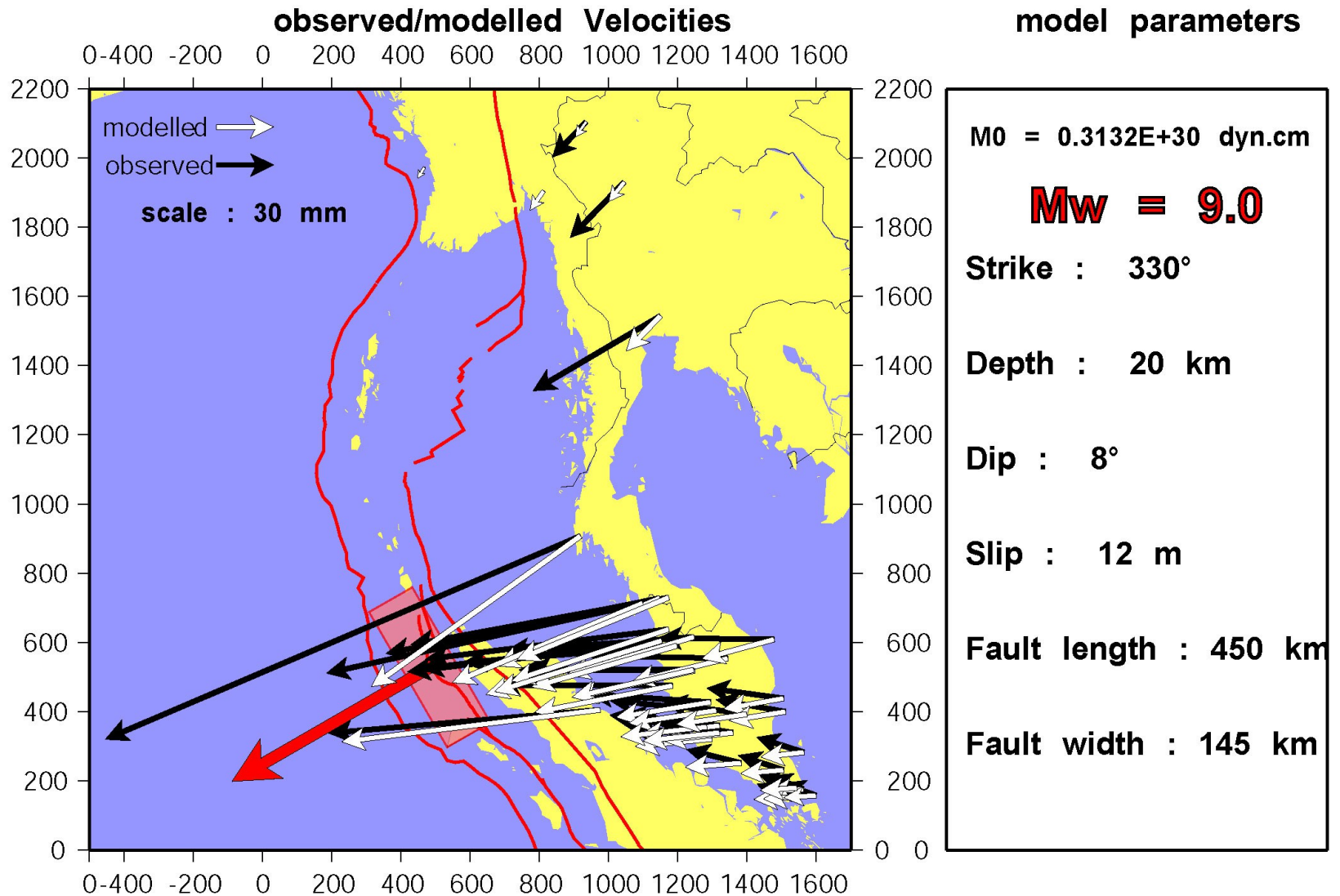


Le séisme de Sumatra du 25 décembre 2004

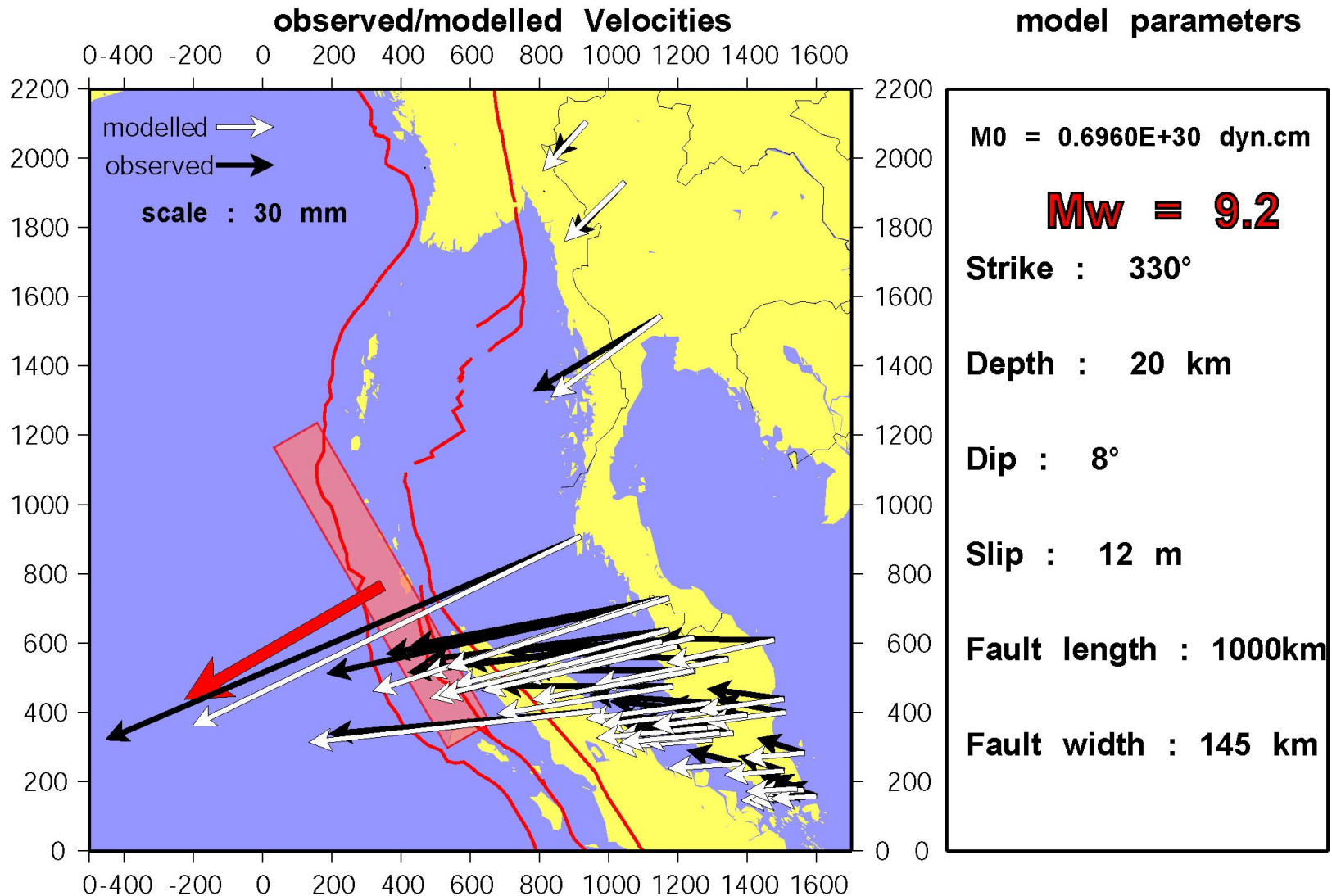




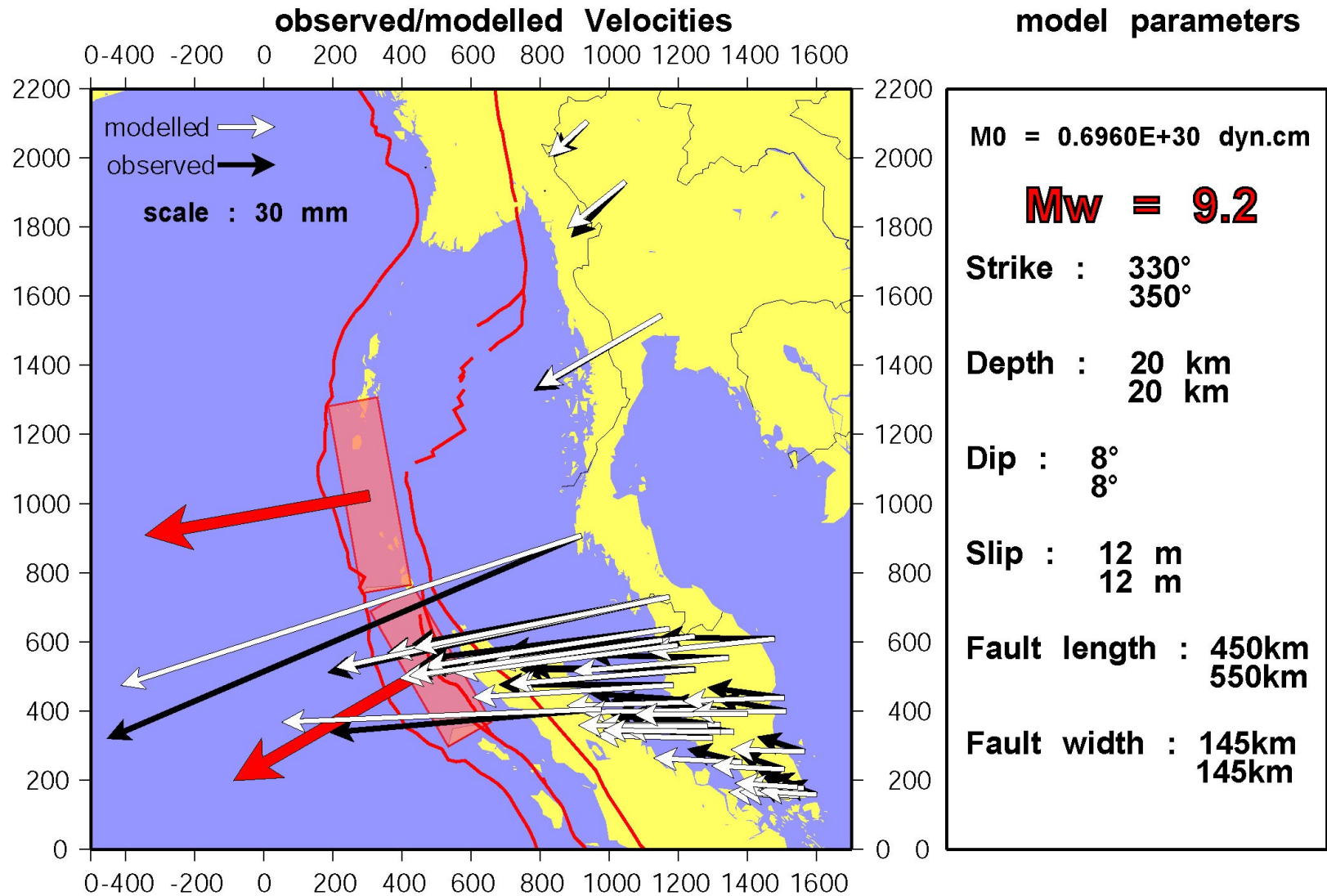




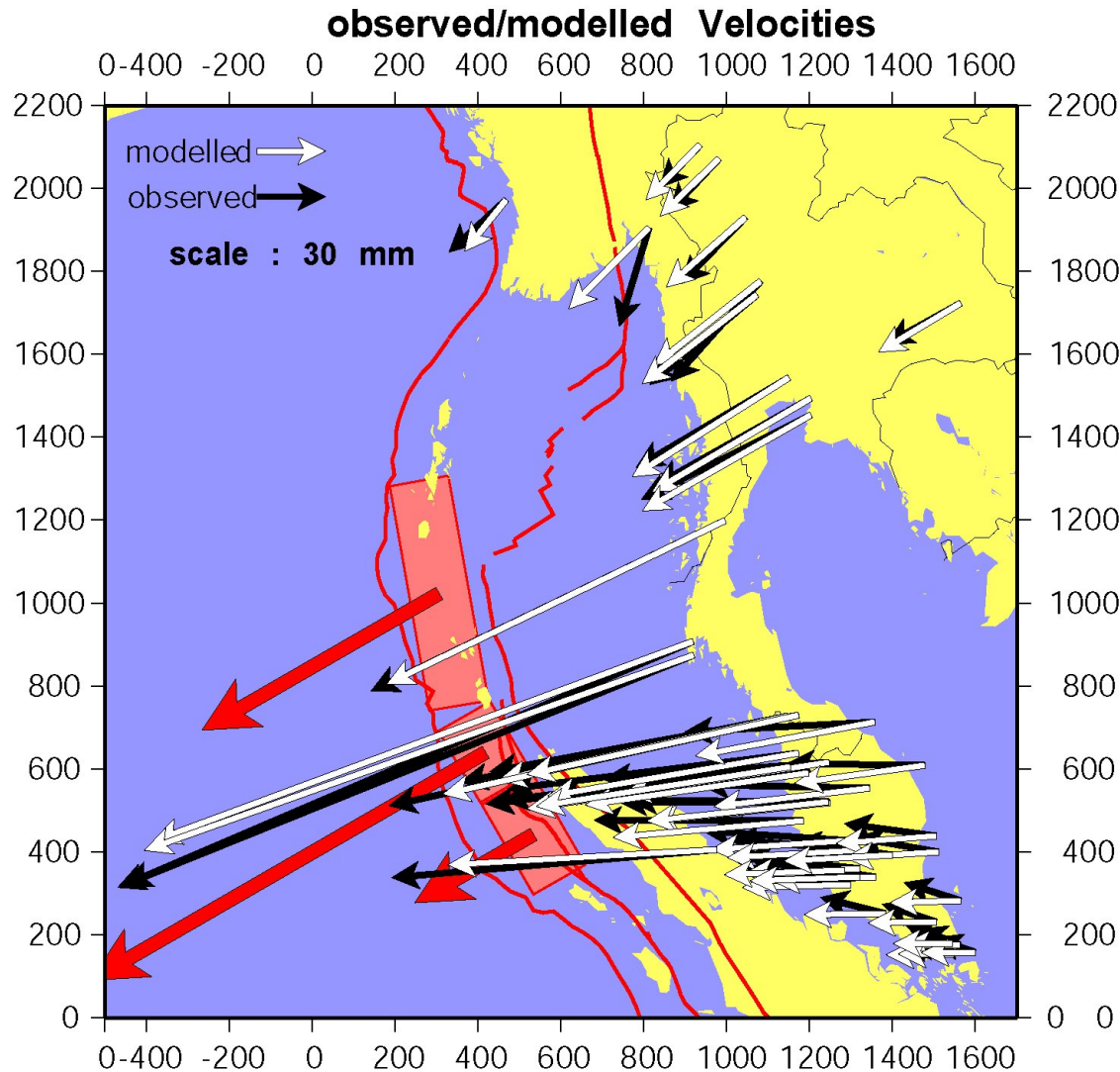
**A rupture of 450 km length gives the reported magnitude ($M_w=9.0$)
but it does not fit the observed deformation**



**A rupture of 1000 km length is required to fit far field deformation
it corresponds to a larger magnitude $M_w=9.2$**



Curvature of the trench must be taken into account to fit observed directions in Northern Malaysia



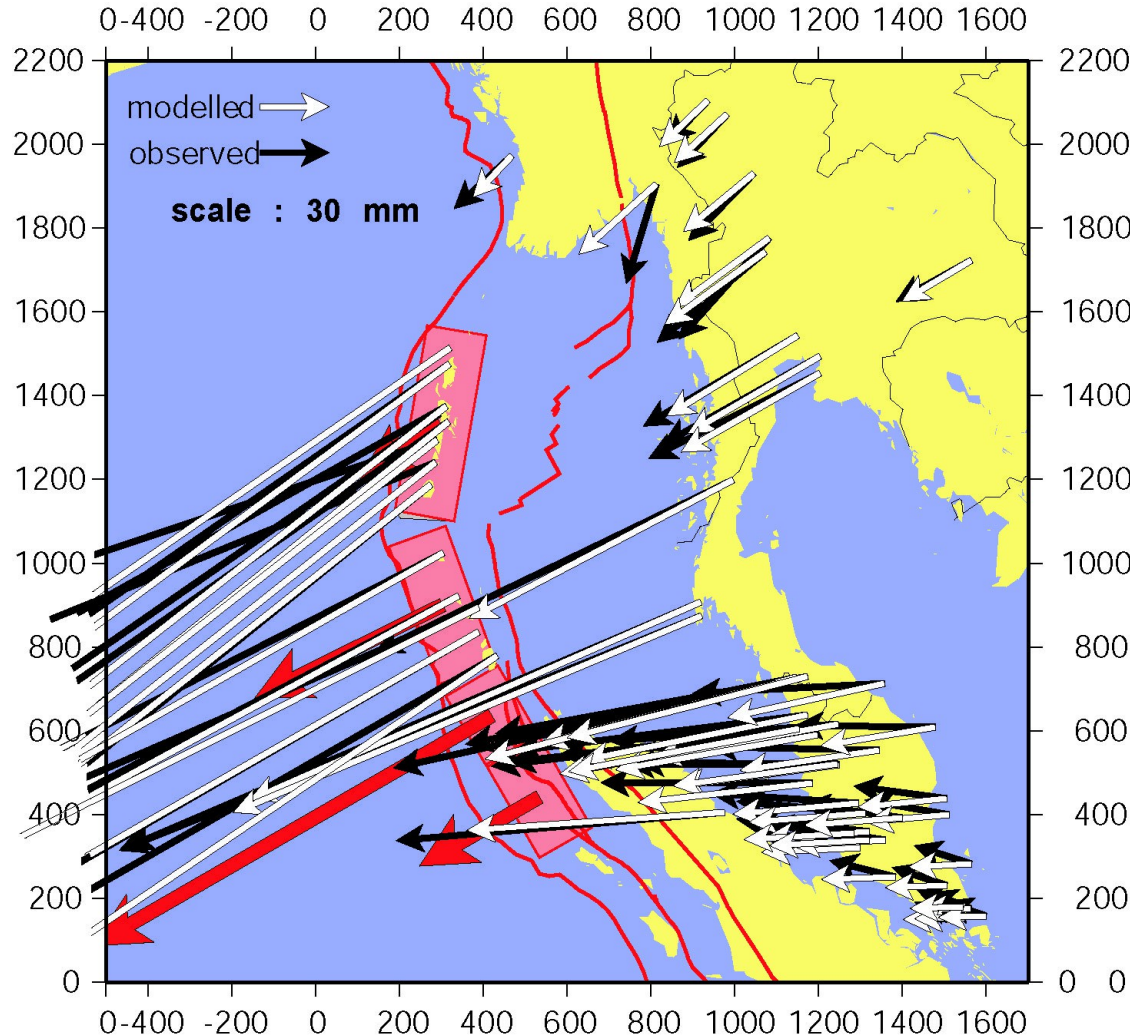
model parameters

M0	7.018E+29 dyn.cm
Mw	9.2
Strike	330° 330° 350°
Depth	20 km 20 km 20 km
Dip	8° 8° 8°
Slip	6 m 20 m 12 m
Fault length	250km 200km 550km
Fault width	145km 145km 145km

New Myanmar data can be fit with previous models, but...

mod

observed/modelled Velocities



$M_0 = 6.020E+29$ dyn.cm

Mw = 9.2

Strike : 330°
330°
340°
10°

Depth : 20 km

Dip : 8°

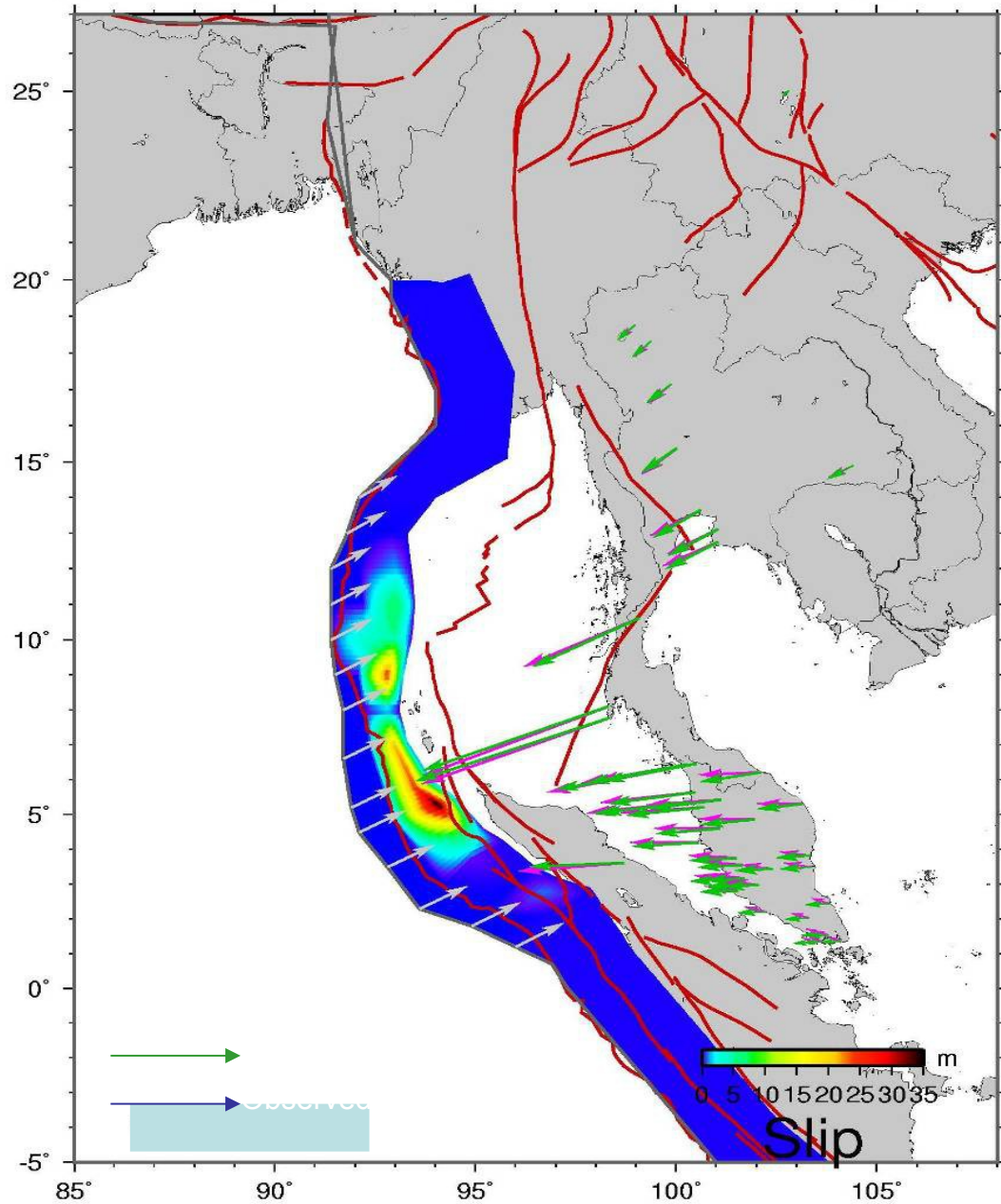
Slip : 6 m
20 m
9 m
4 m

Fault length : 250km
200km
350km
450km

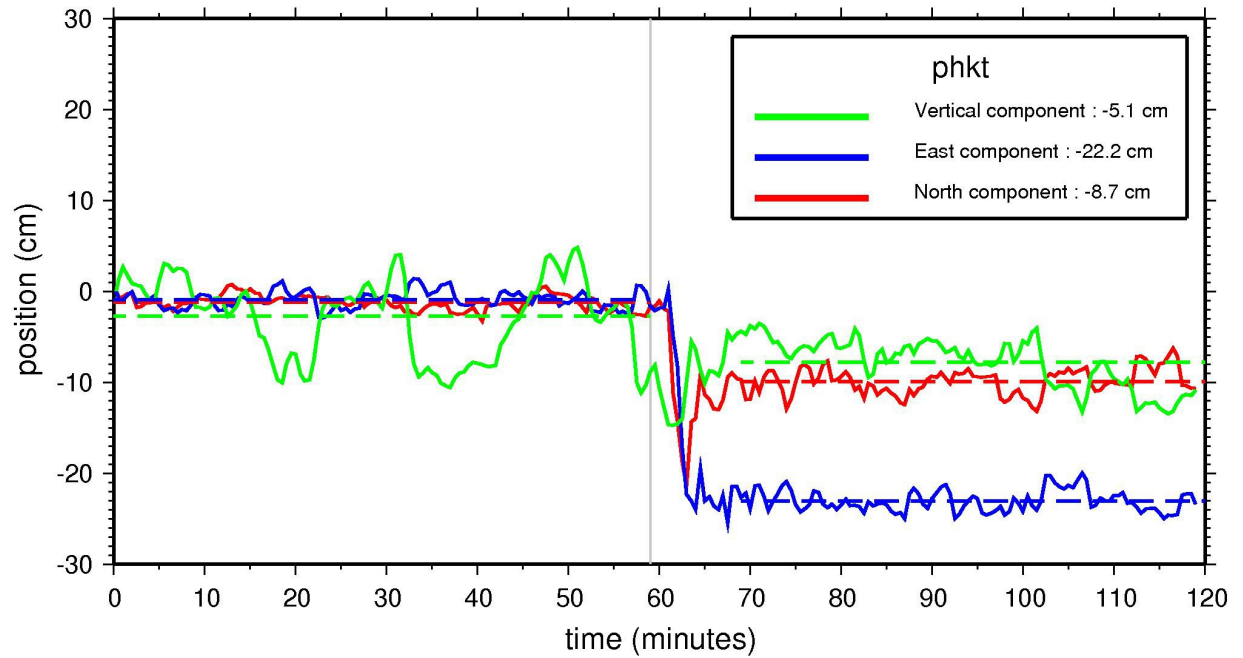
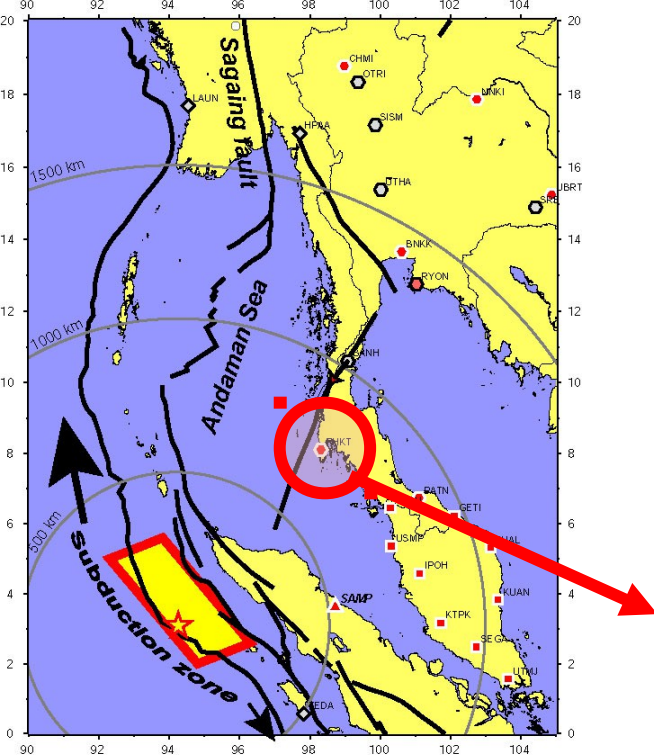
Fault width : 145km

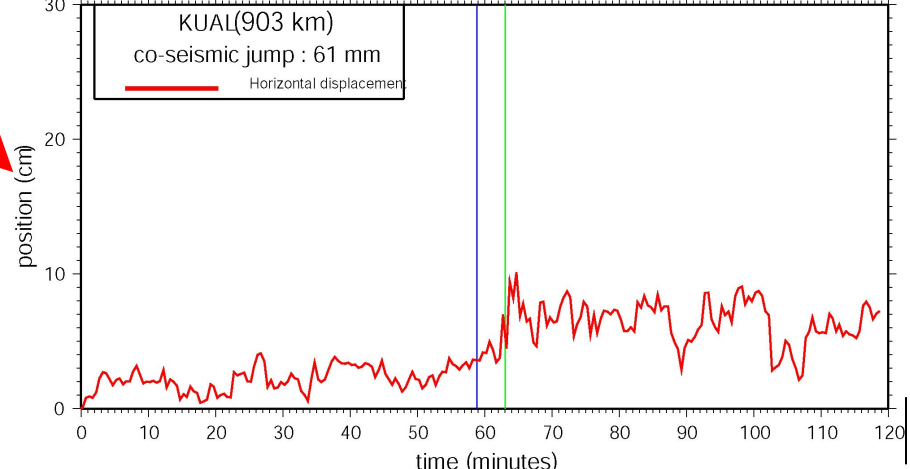
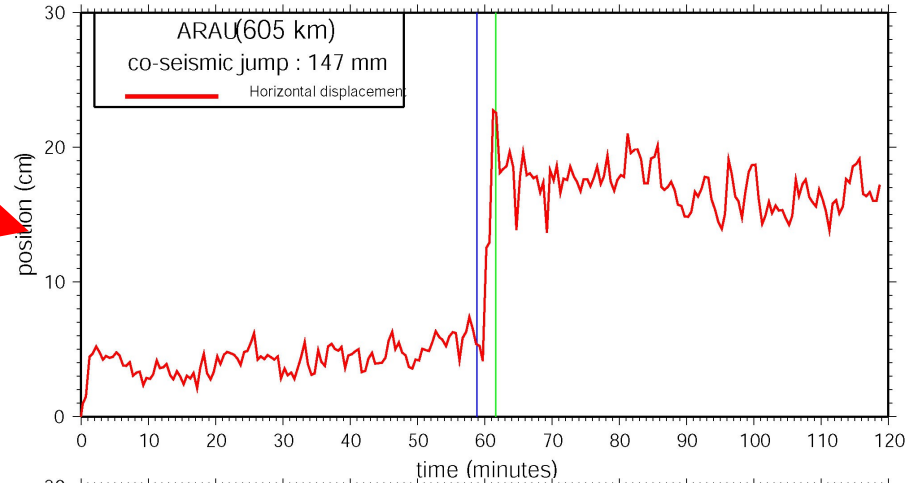
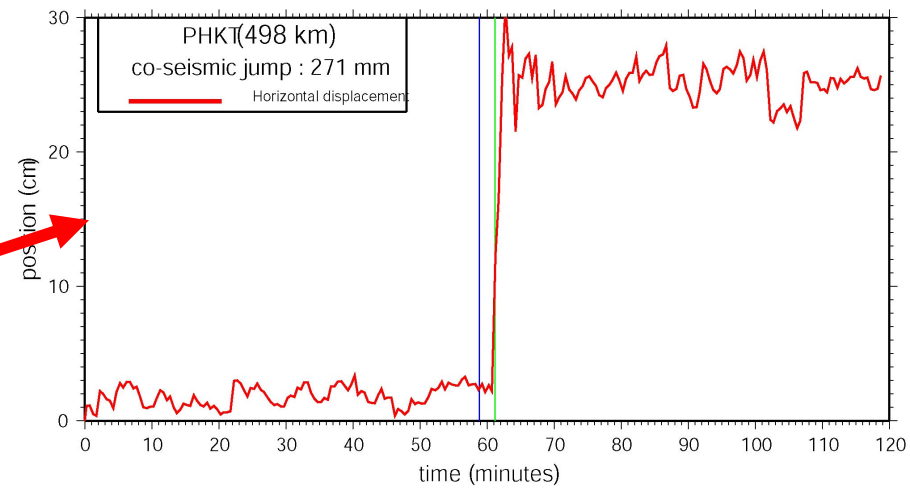
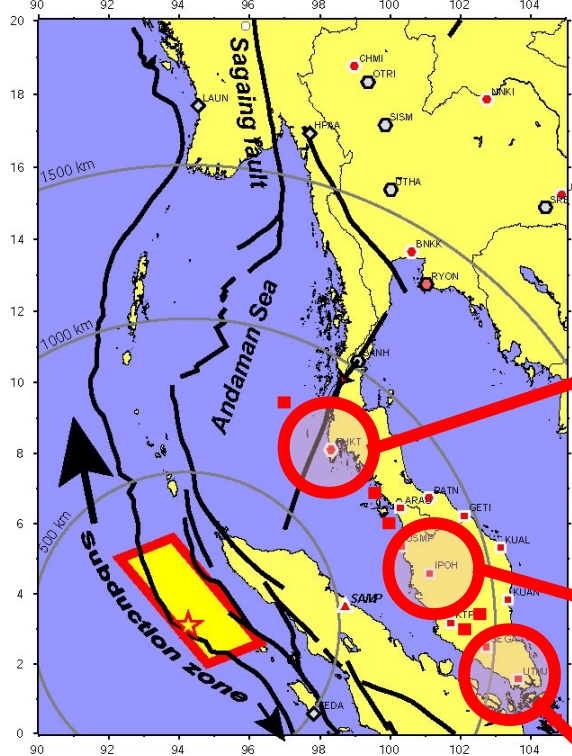
... Andaman data requires longer plane AND oblique slip

Full inversion of slip on fault



Kinematic solution at Phuket



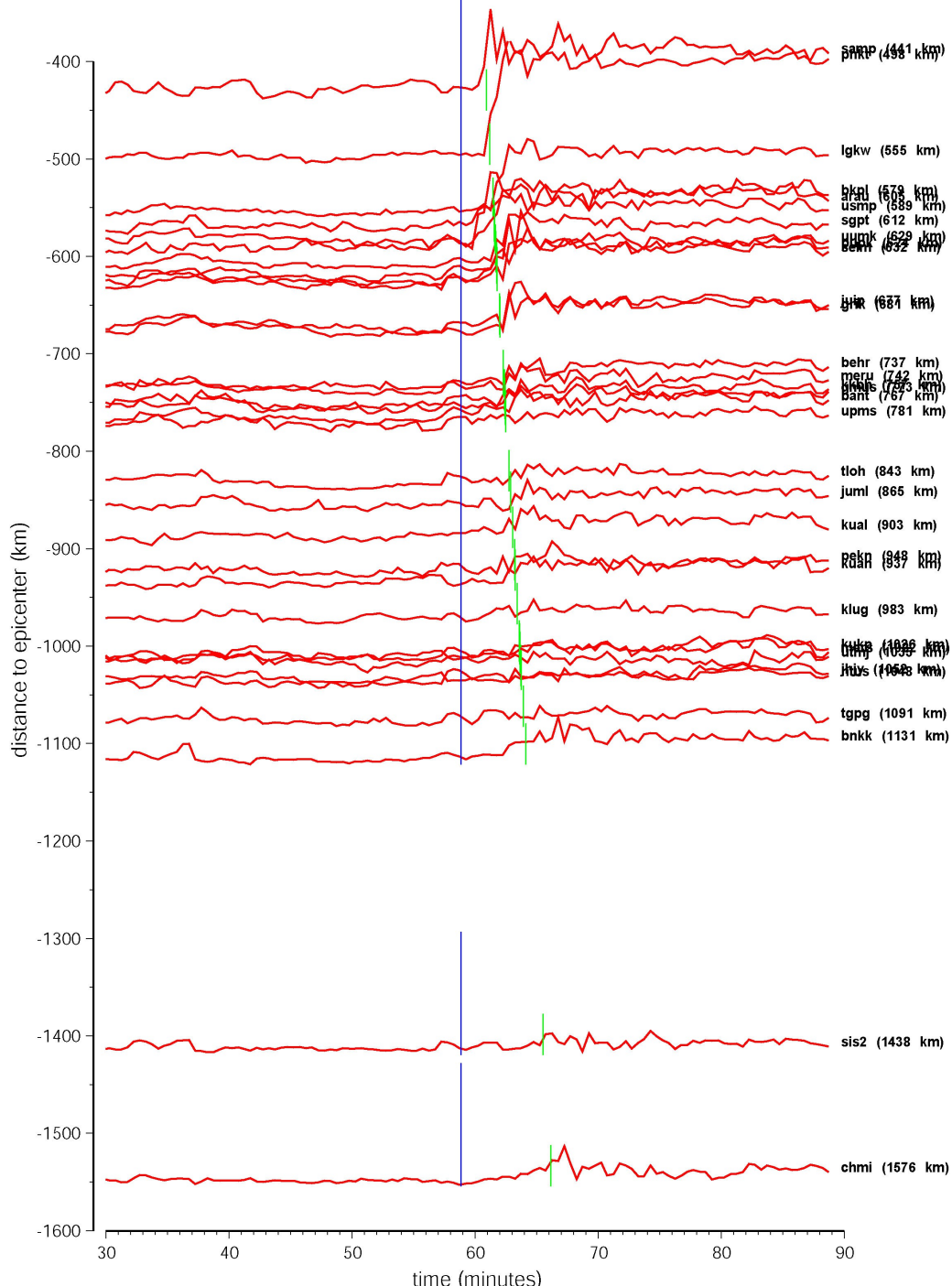


“Kinematic” (epoch-by-epoch) positioning of the GPS station show the co-seismic step...

...and allow to determine the displacement arrival time

It seems related to surface waves rather than P or S wave.....

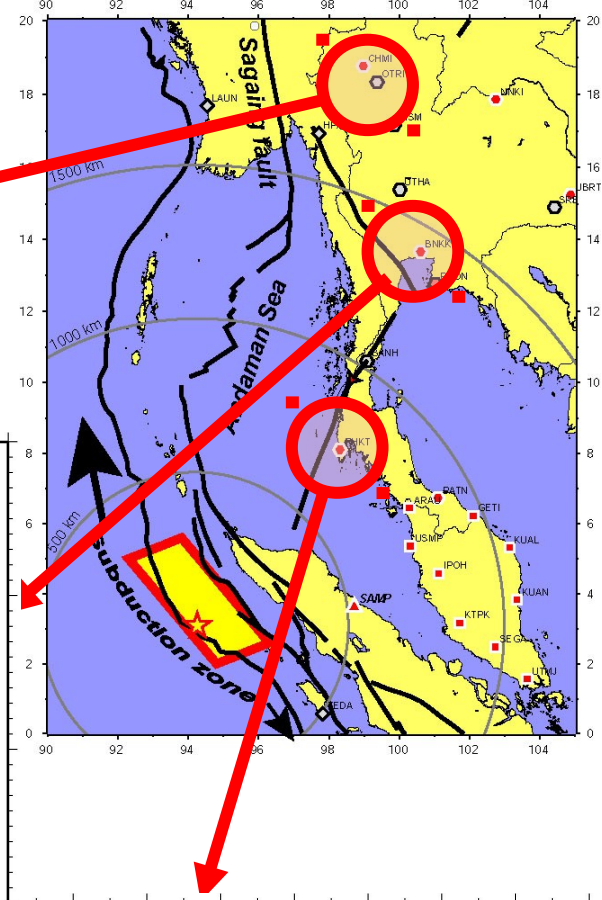
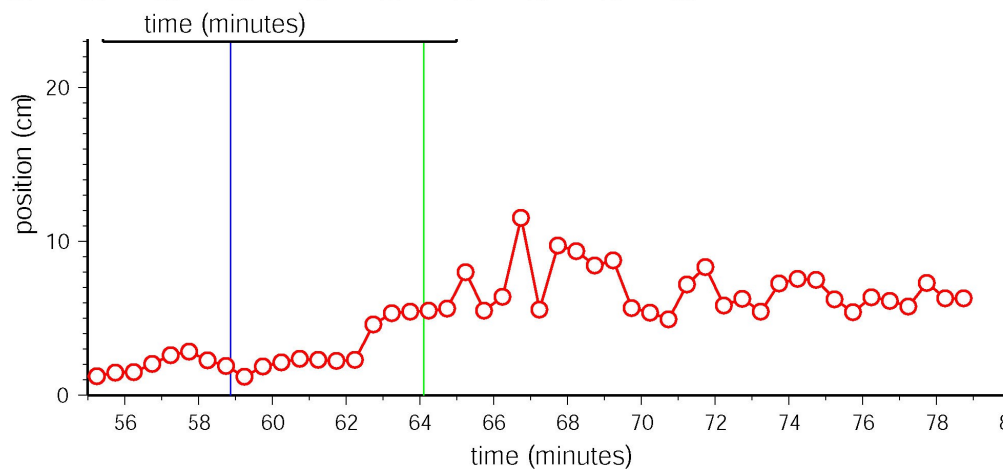
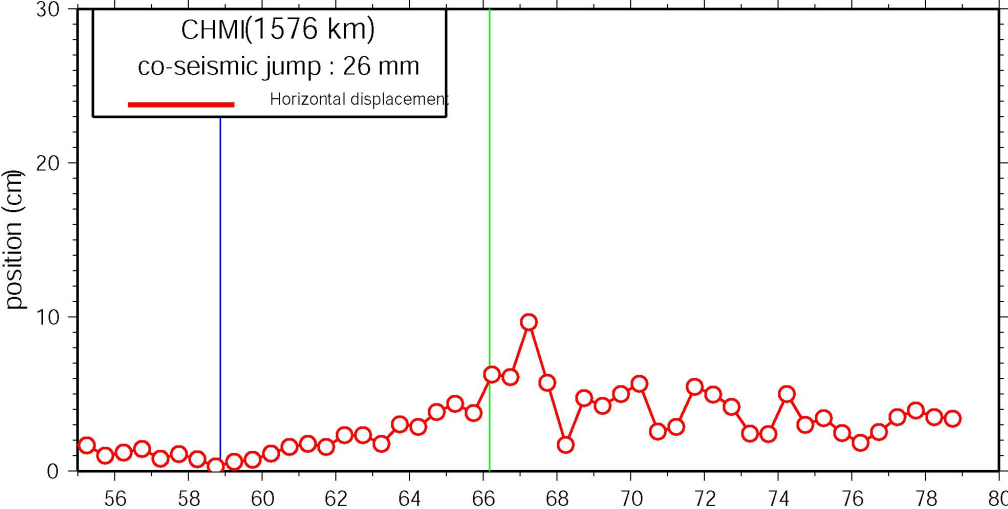
GPS constrained Epicenter : 95.0E 5.0N



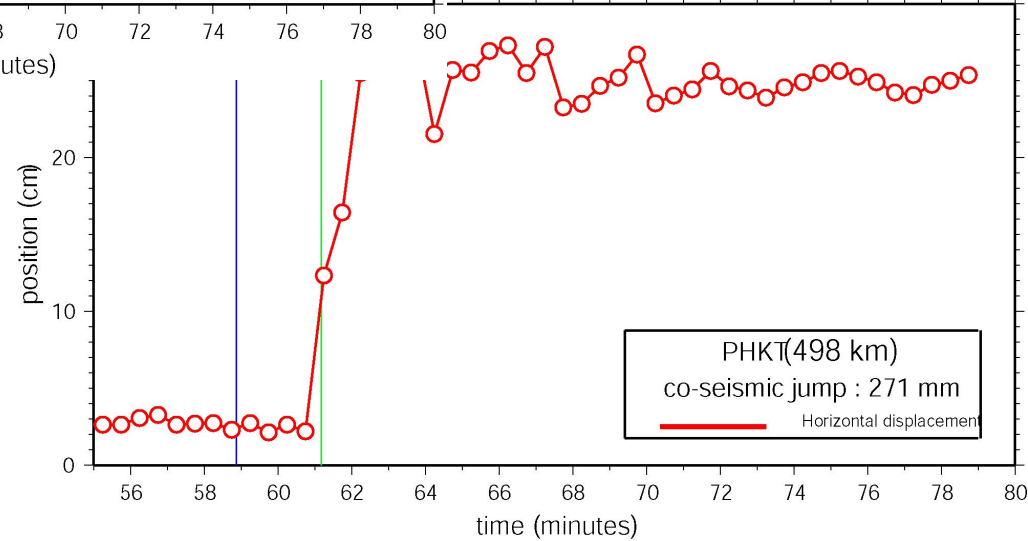
**Assuming a
velocity of 3.6
km/s for seismic
waves**

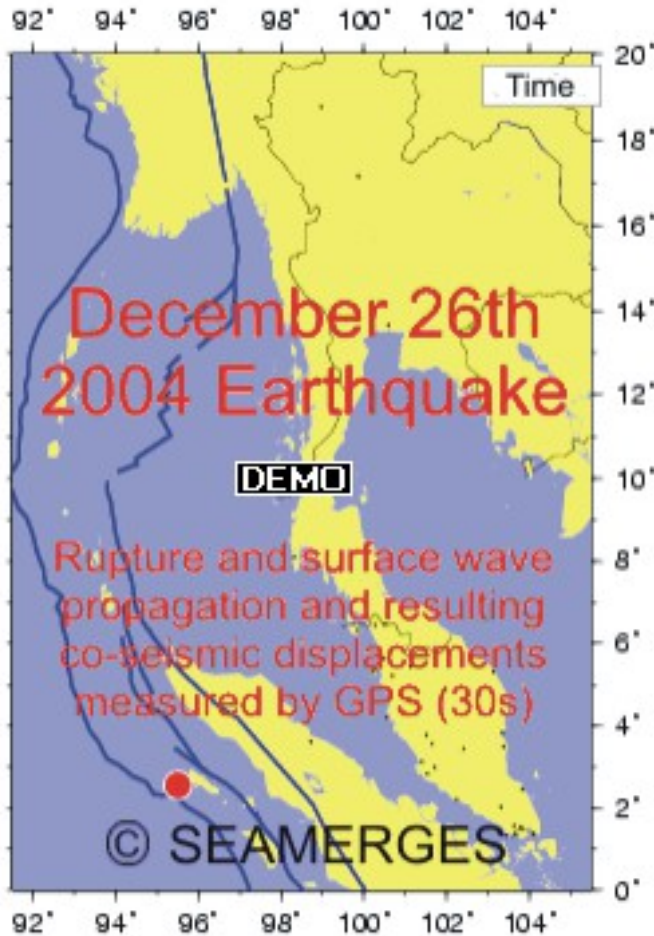
**relocation of the
source of the
seismic energy is
needed to match
and sort arrival
times at stations**

**Again, a
relocation of 200
km to the north is
requested**



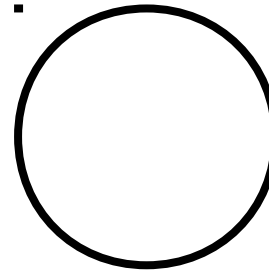
Indication of source directivity is pointed by larger "rise times" at northern stations





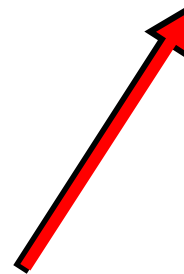
<http://www.deos.tudelft.nl/seamerges>

● rupture



Seismic **surface**
waves propagation

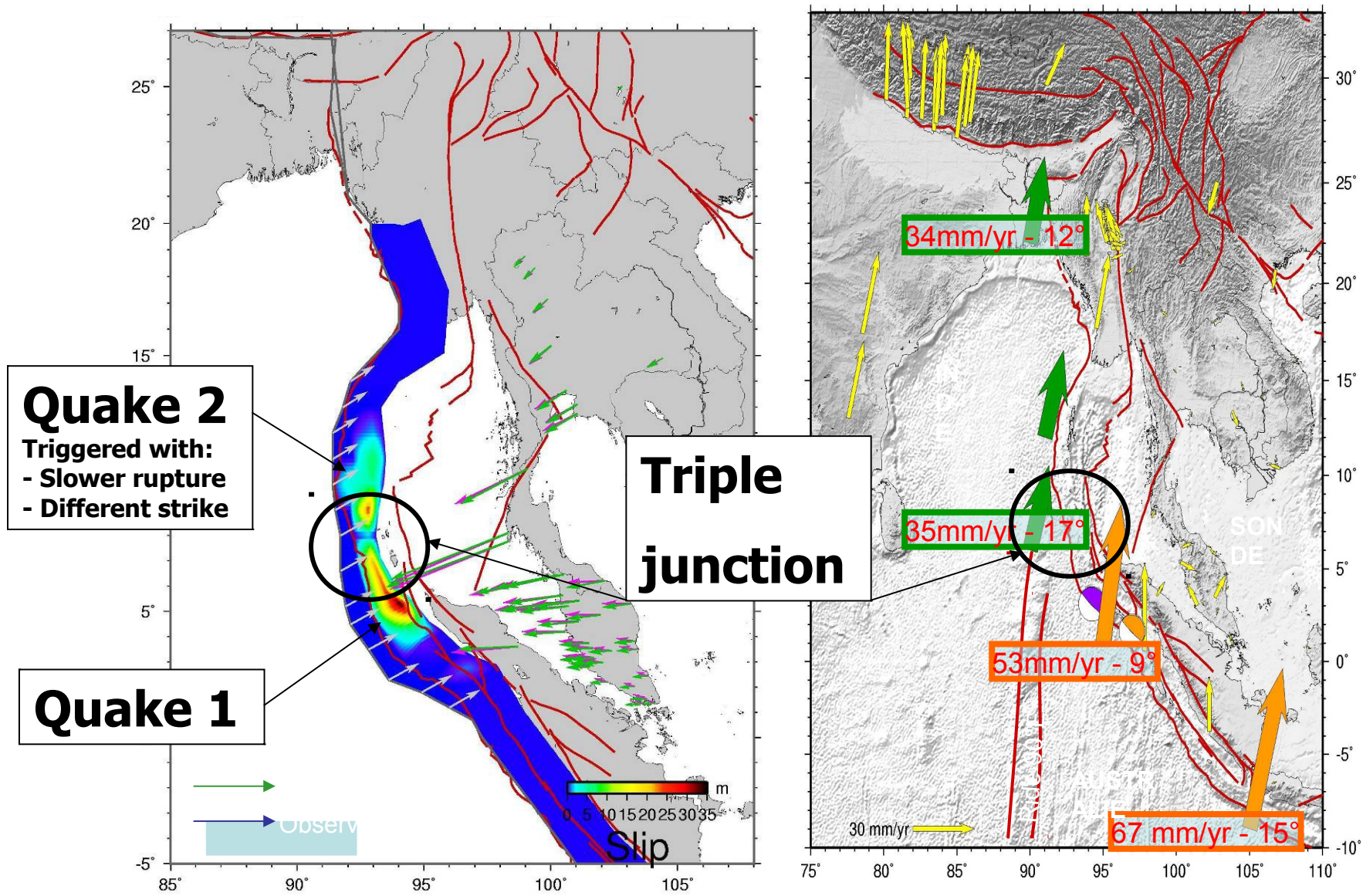
(3.7 km/s)



GPS stations
displacements

Rupture Propagation:
3.7 km/s initially (South)
30s stop $\sim 8^\circ$ lat
1.8 km/s onward (North)

GPS cinématique => vitesse de rupture



GPS cinématique+statique => 2 ruptures

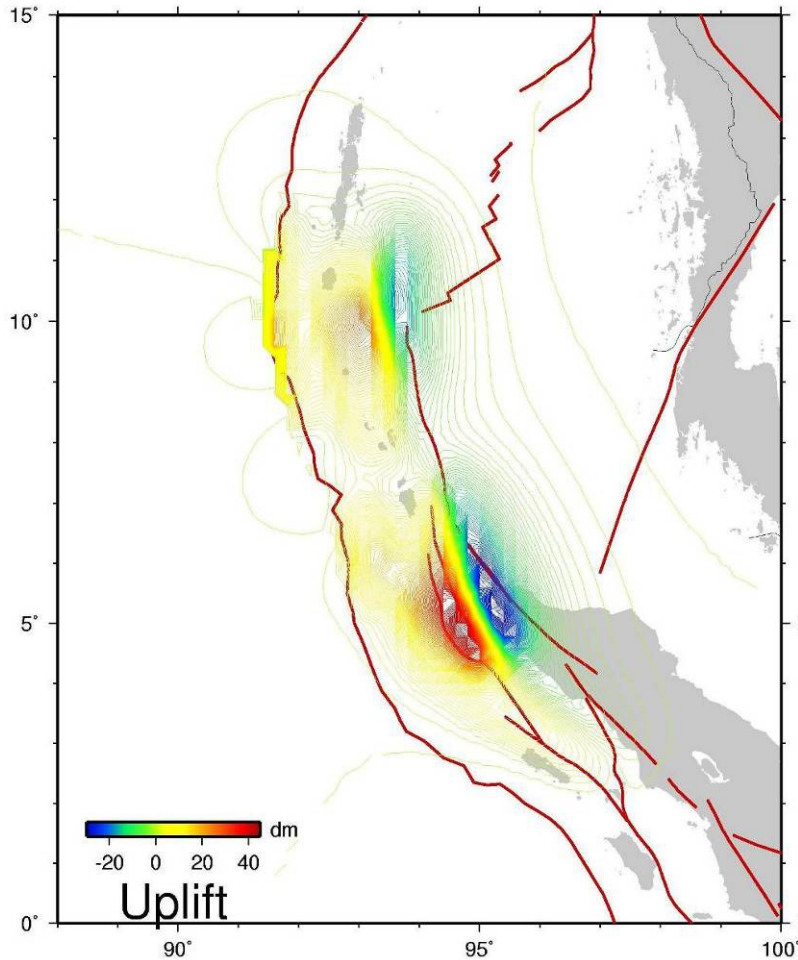
Vertical motions predicted by the models

- 4 m of uplift
- 2 m of subsidence



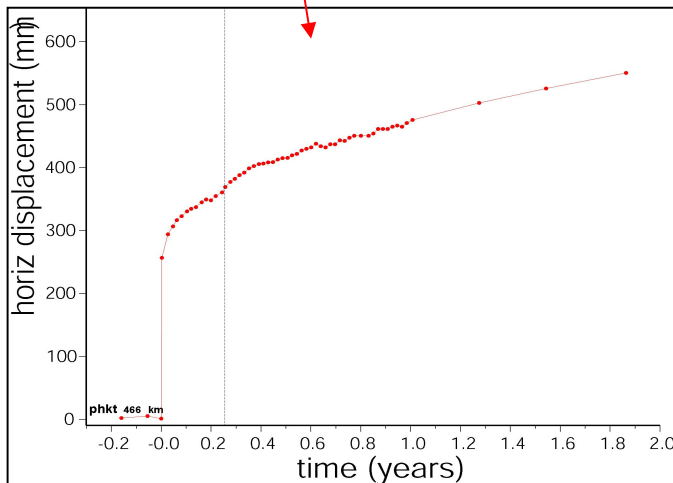
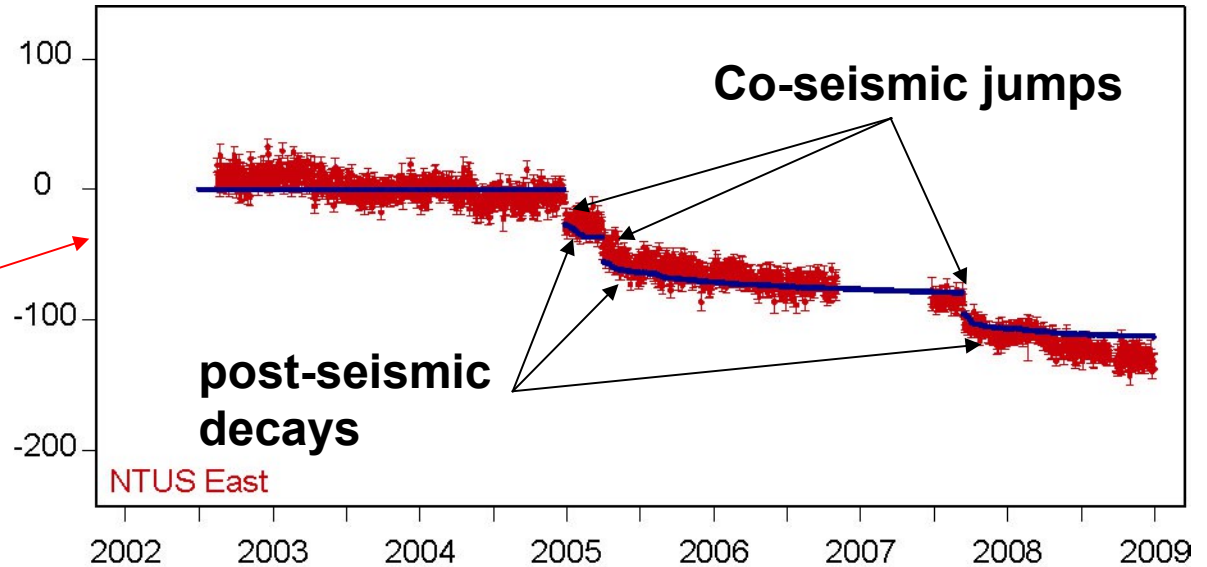
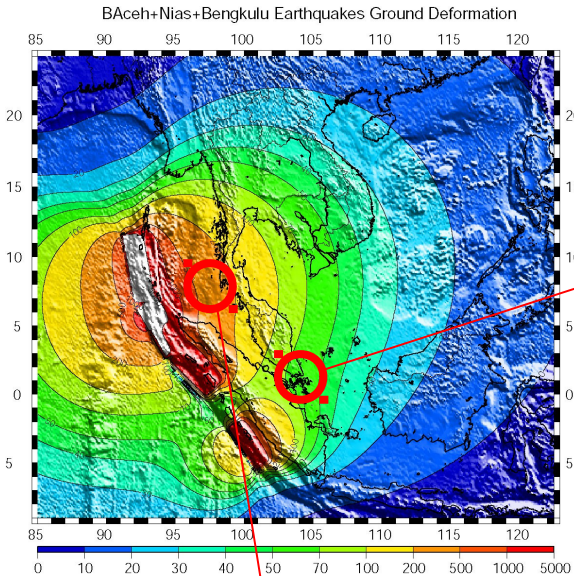
Tsunami modélisation

Pietrzack et al., 2007



GPS cinématique+statique => modèle de Tsunami

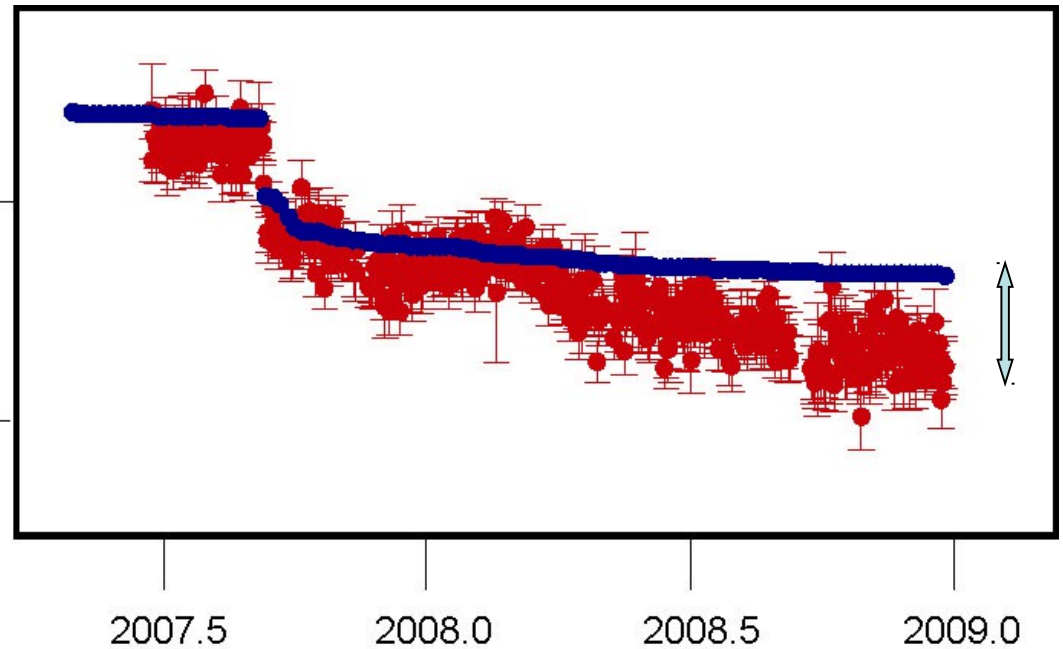
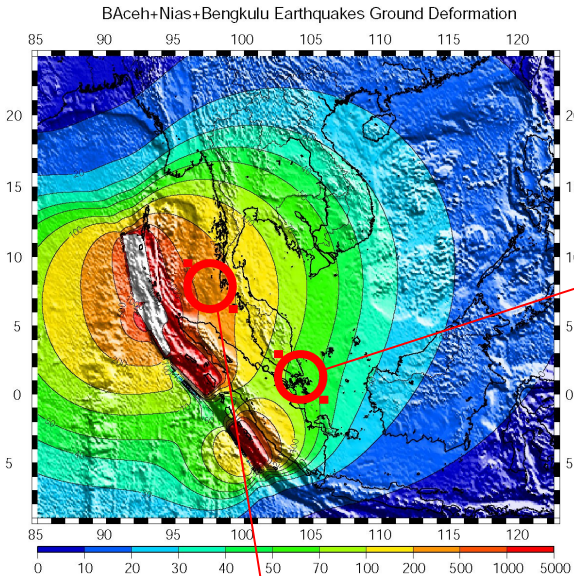
post-seismic deformations



rates are different from what they were before... and they keep changing with time

Reference networks in these countries have to be redone... continuously !!

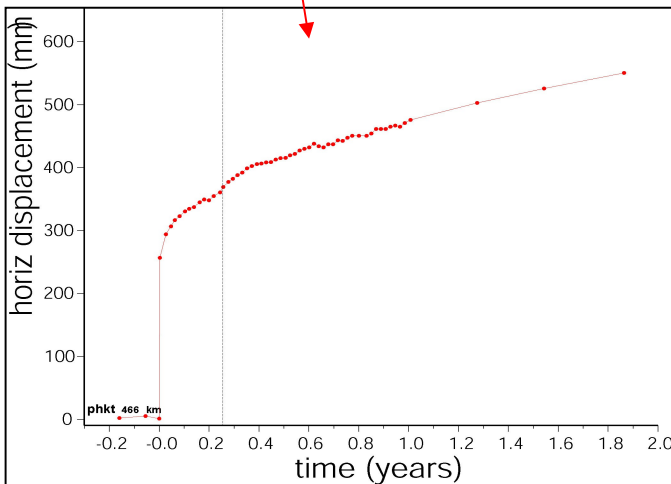
post-seismic deformations



Models don't really match the data

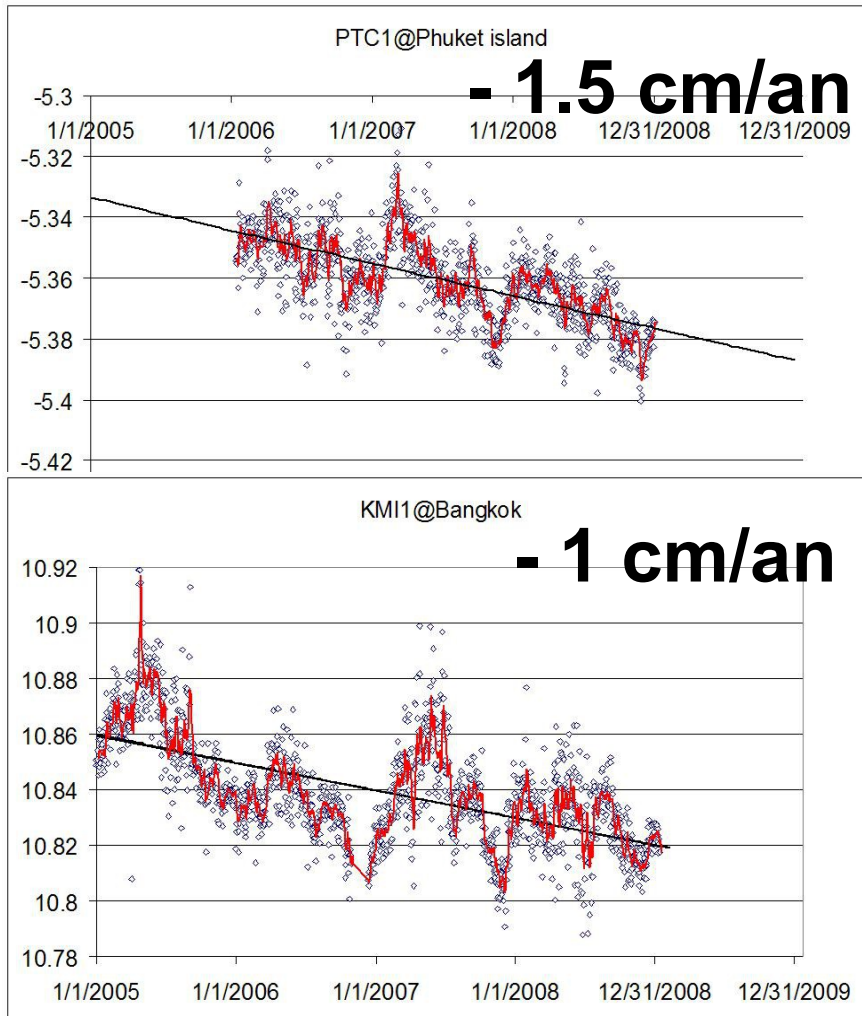
- Unknown type of slip
- Unknown long term effects

More continuous measurements are needed, most probably for much more than a decade

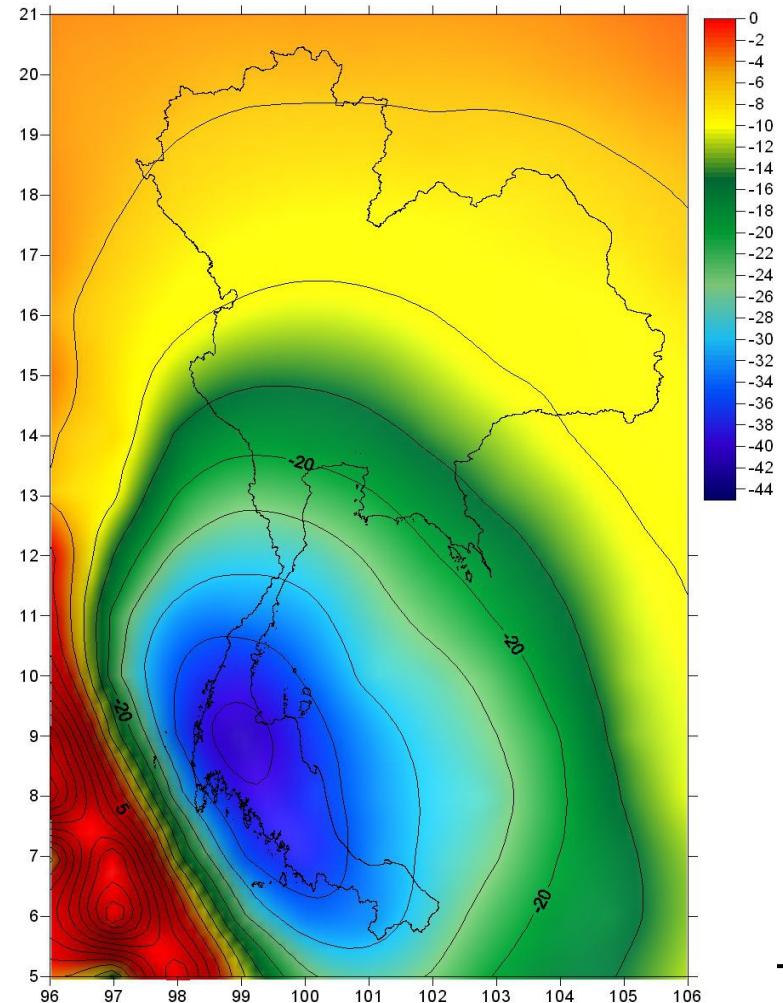


Subsidence en Thaïlande déclenchée par le séisme de Sumatra

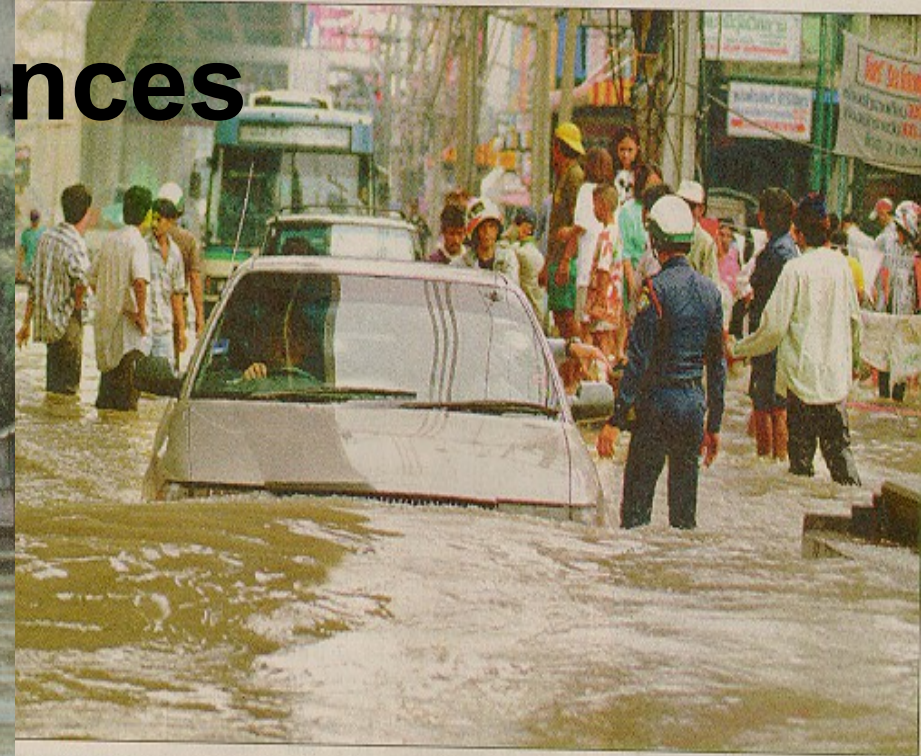
Les Mesures



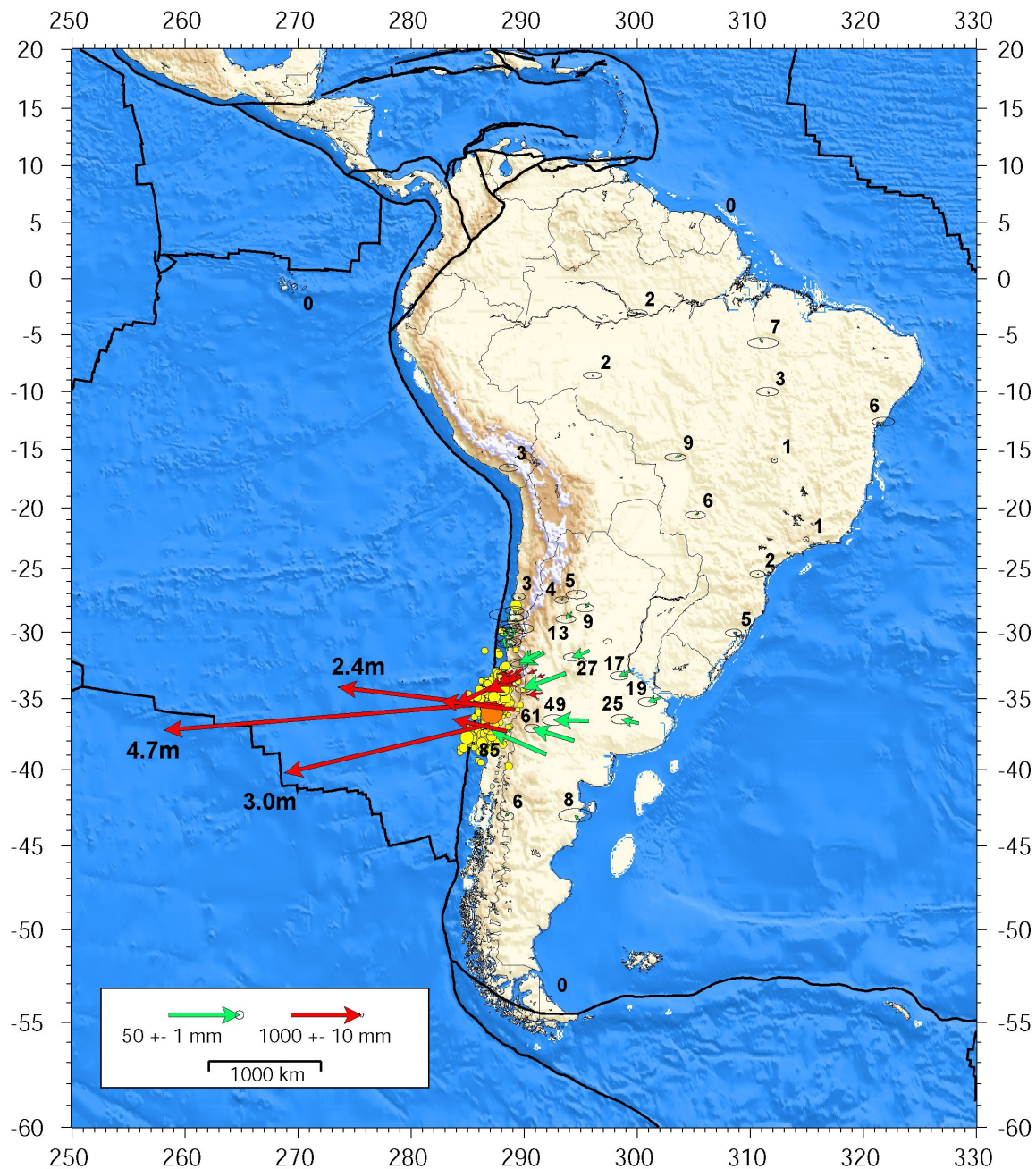
Le Modèle



Les conséquences



Le séisme du Chili – 27 fev. 2010



South America
cGPS solution:
59 stations

IGS : 5

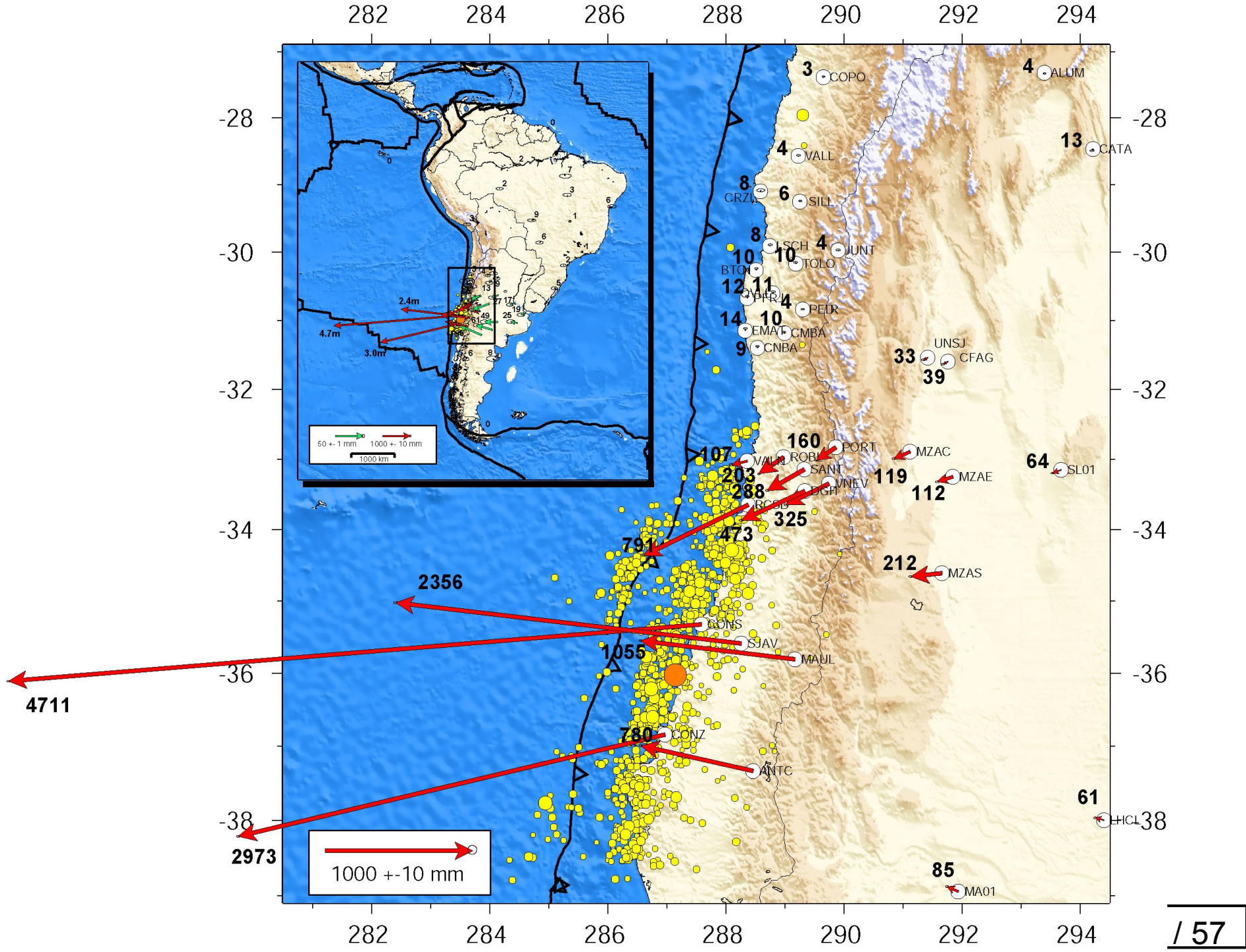
RAMSAC : 16

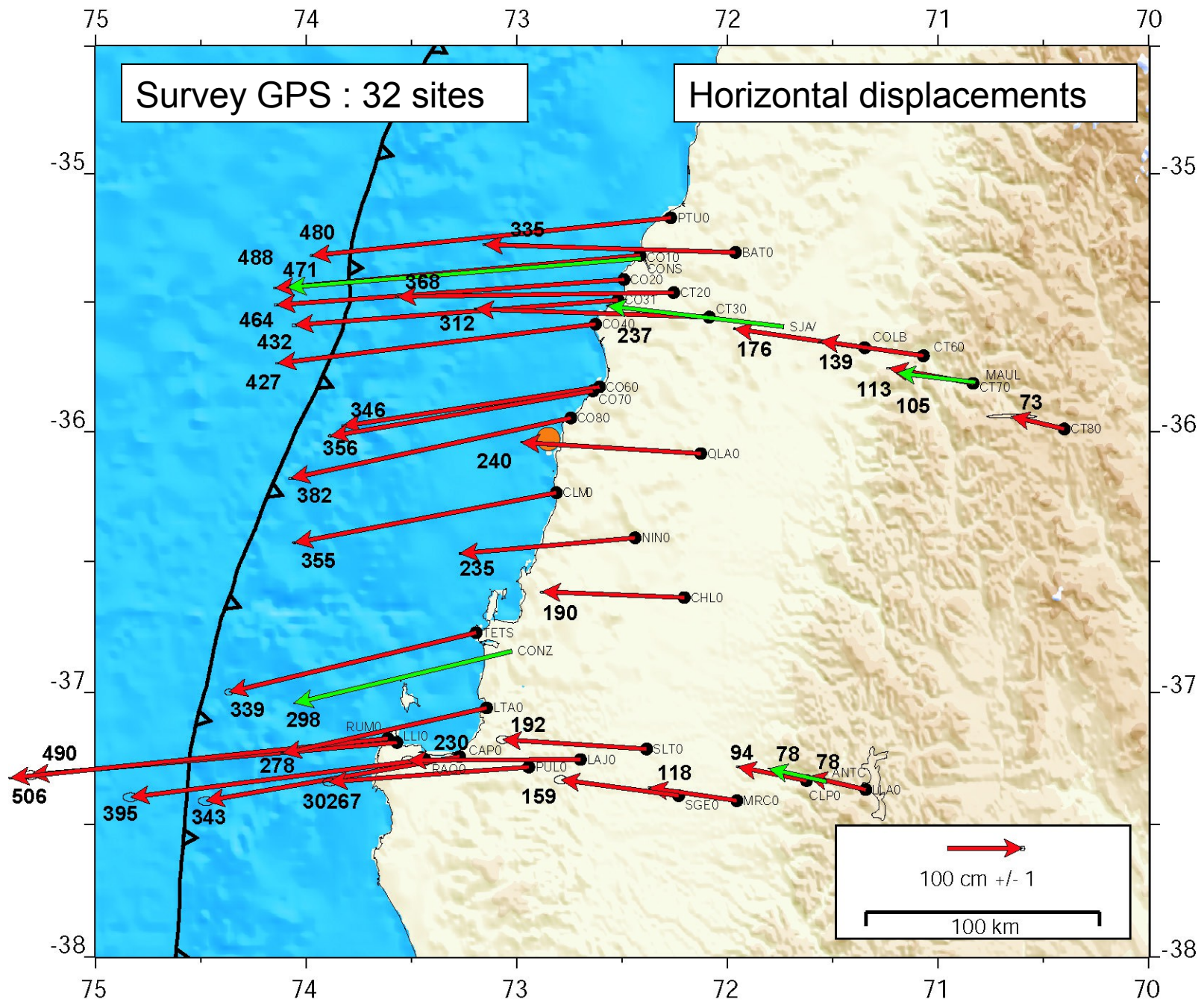
RBMC : 10

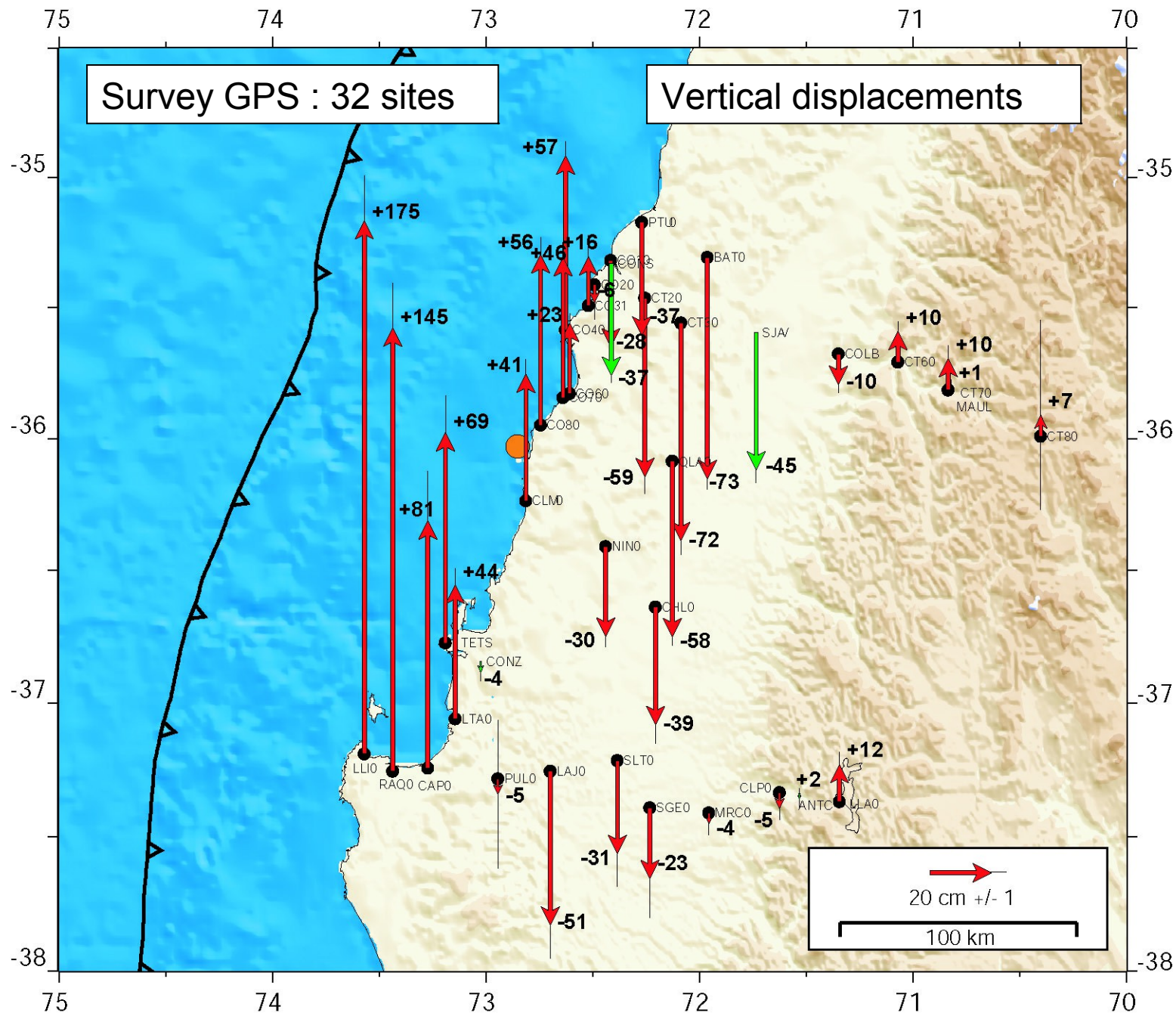
TIGO : 1

CAP : 5

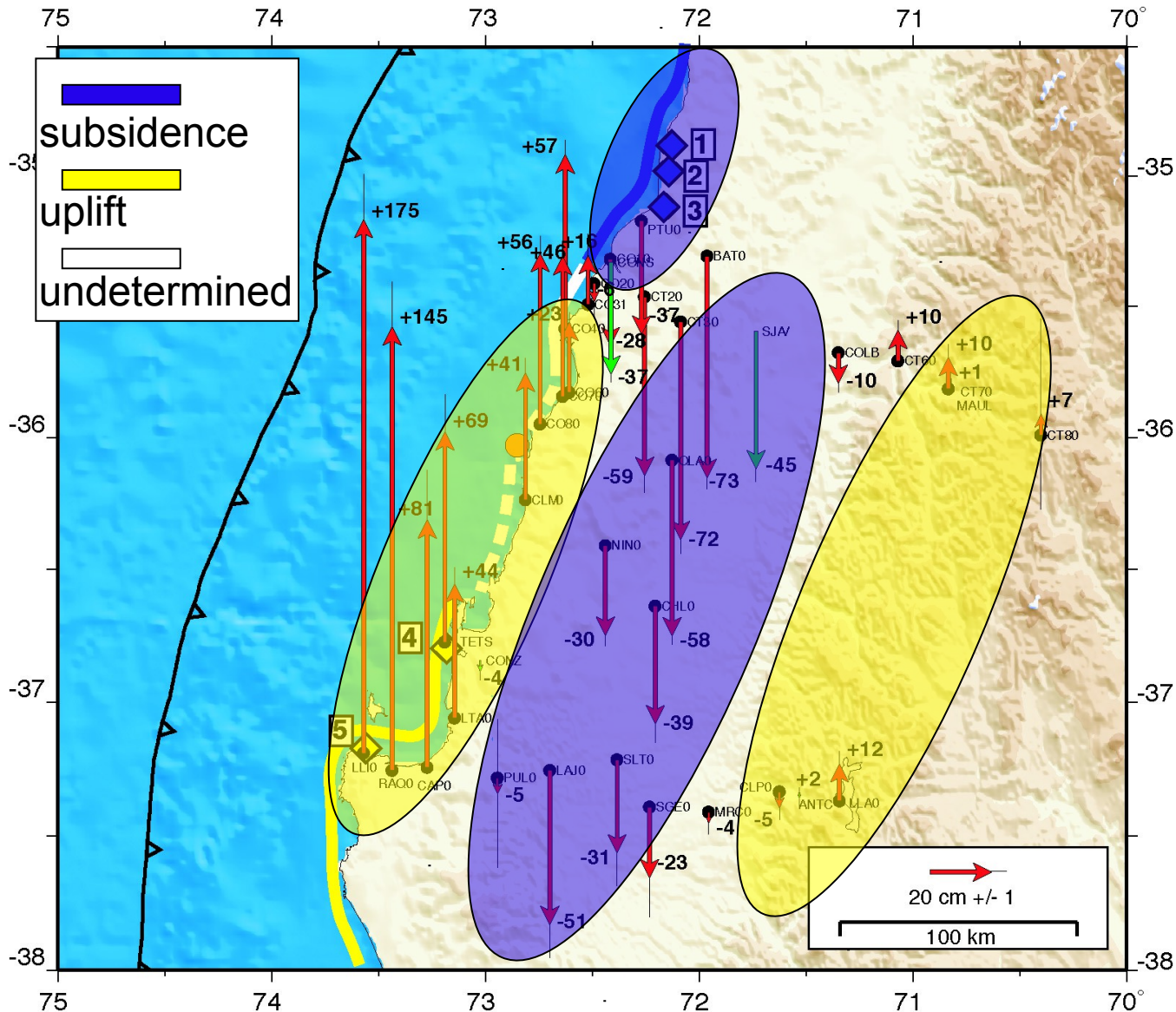
LIA-MdB : 22







Vertical displacements : comparison with field observations



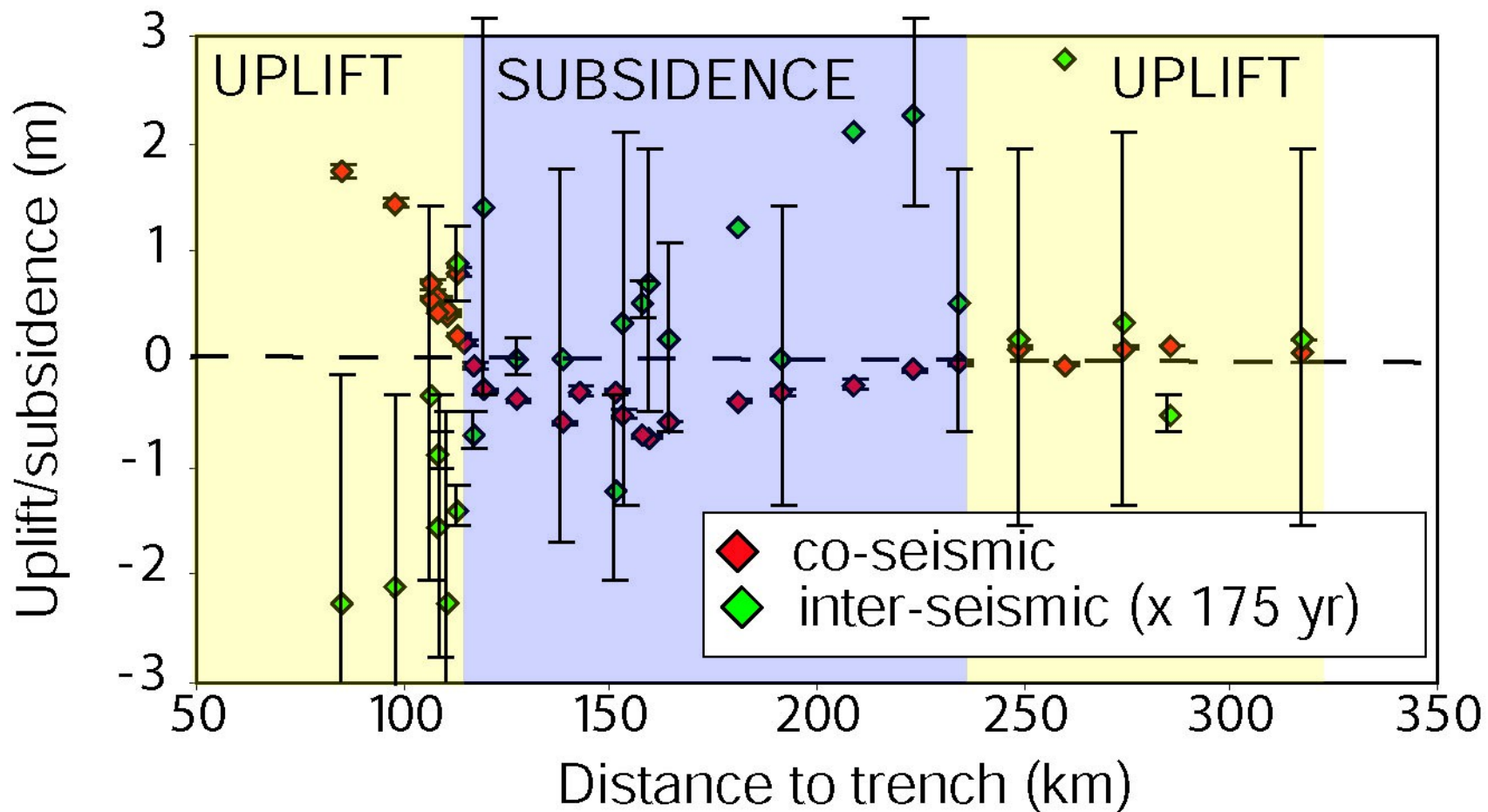
1- Caleta Duao
(high tide line)
-0.8m – -1.6m

2- Caleta La Pesca
(high tide lines)
-1m – -2m

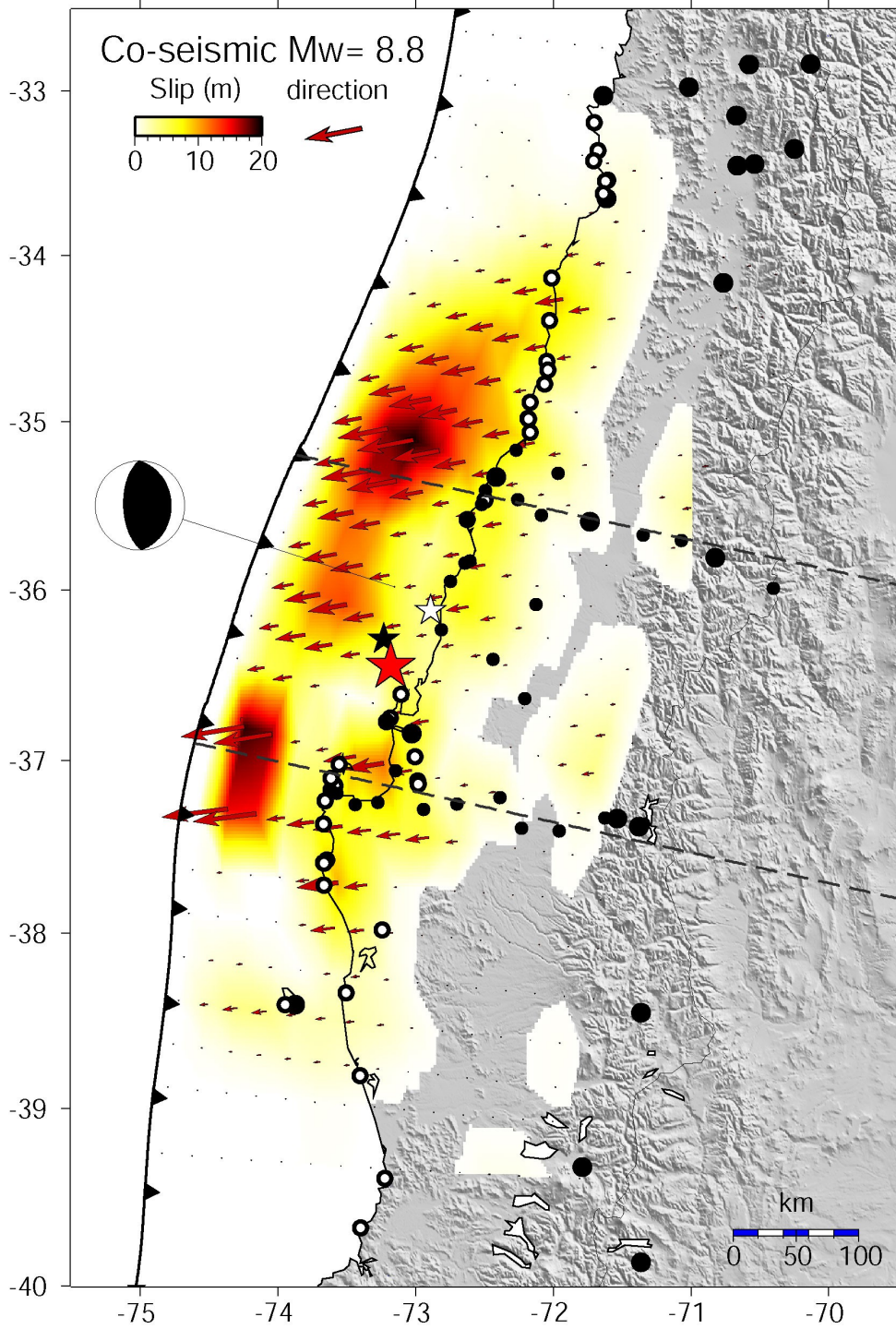
3- La Trinchera
(swimer testimony)
-1m – -1.5m

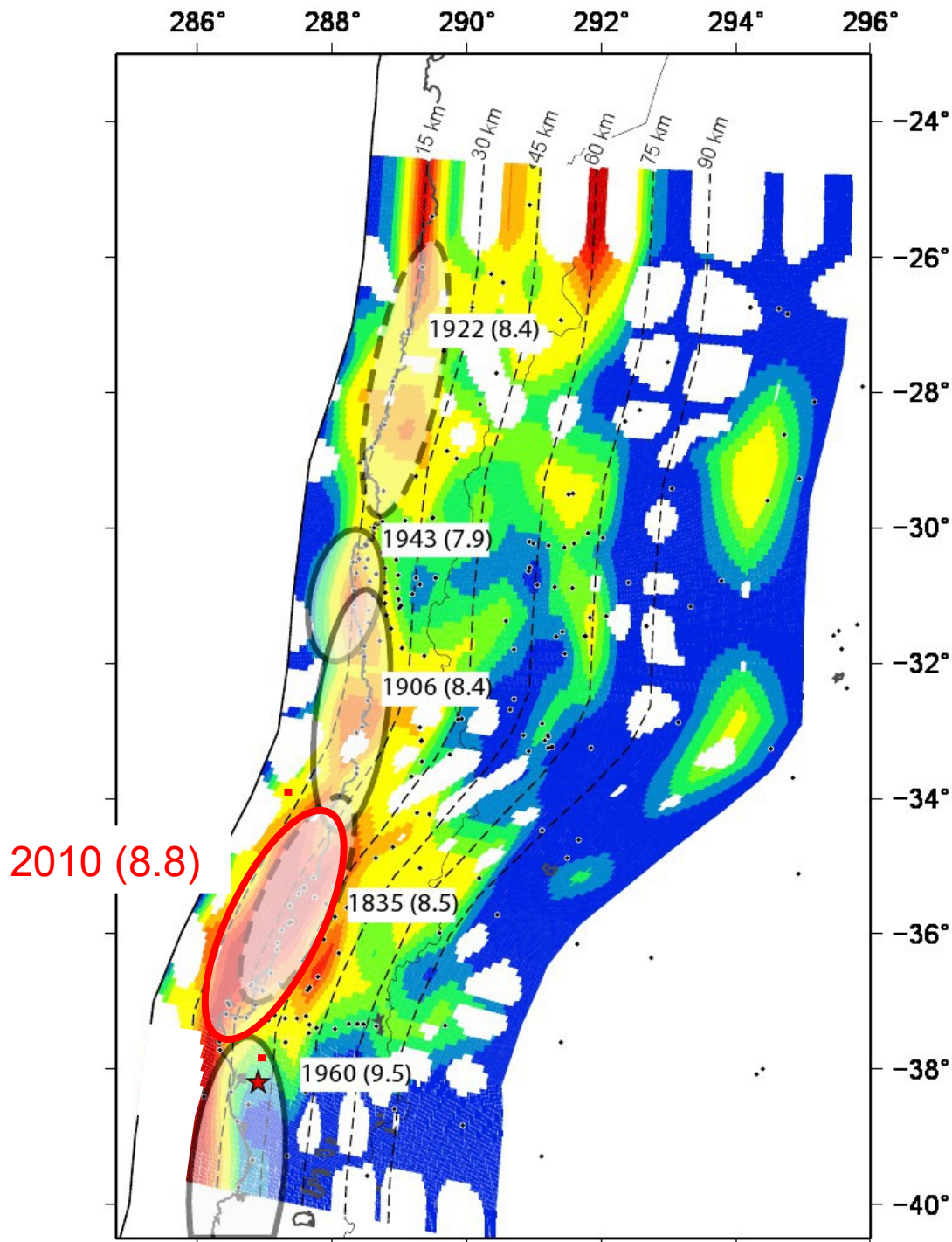
4- Caleta Chome
(high tide lines)
+0.6m – +0.8m

5- Caleta Lavapié
(high tide line)
~+2m



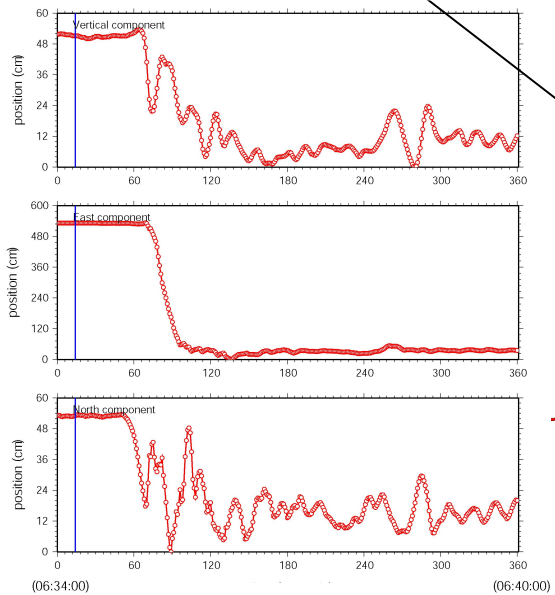
Le modèle de glissement sur la faille contraint par toutes les données: (GPS + INSAR + déplacement relatif du sol)



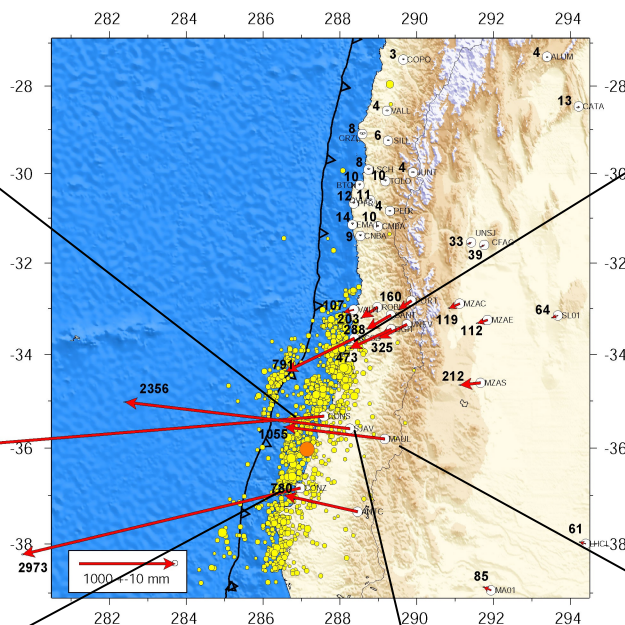
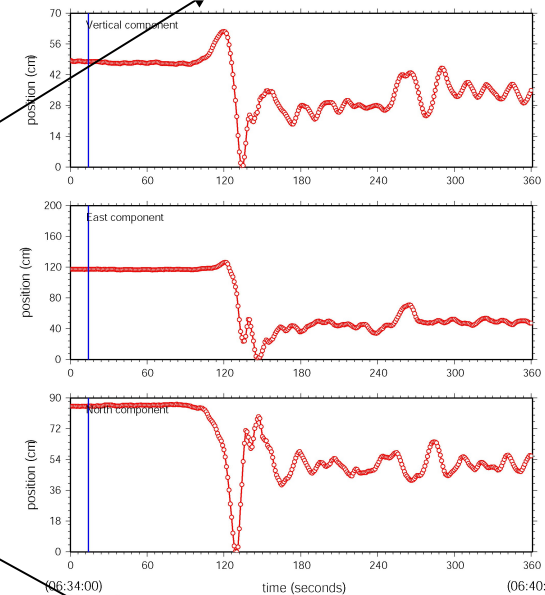


Très bonne
Corrélation
entre la
zone
couplée et
les ruptures
sismiques

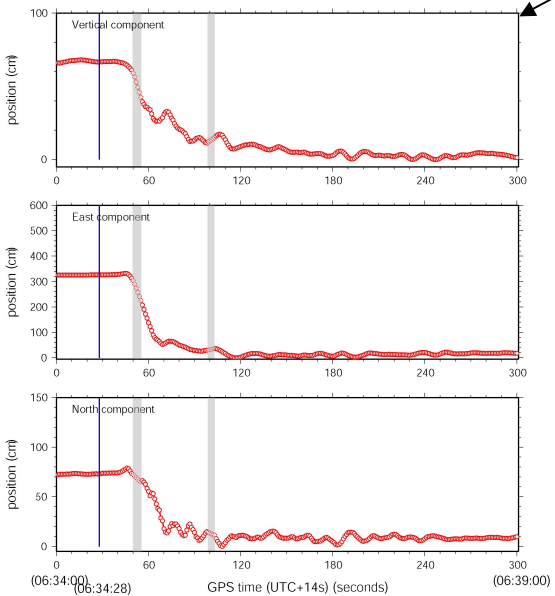
Constitucion



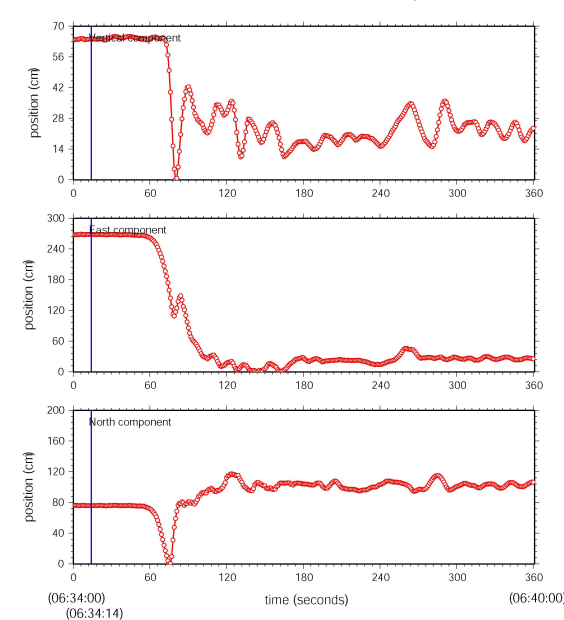
Santo Domingo



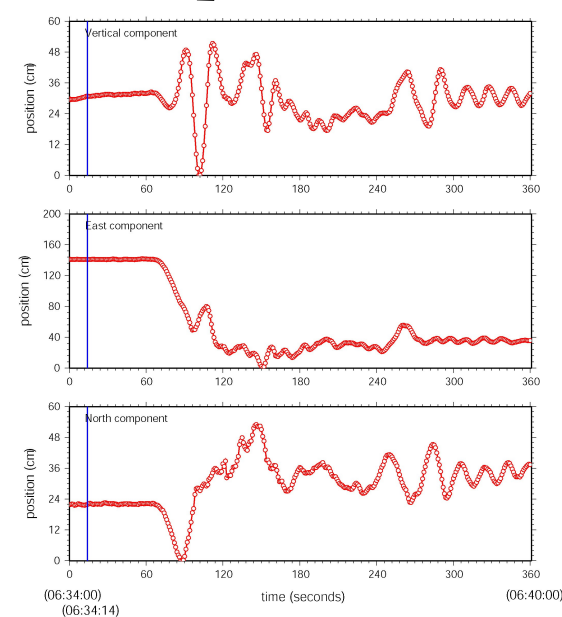
Concepcion



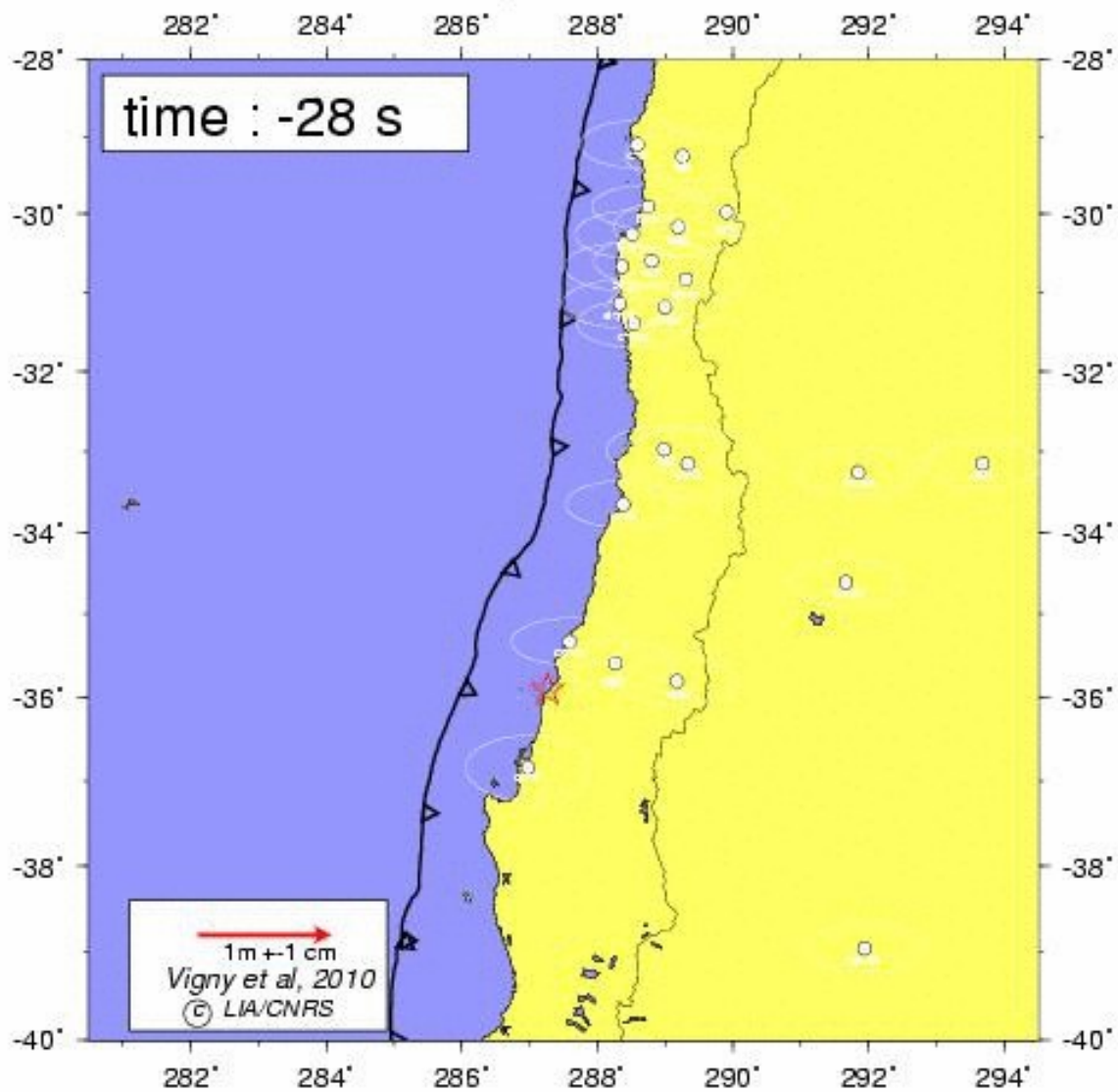
San Javier

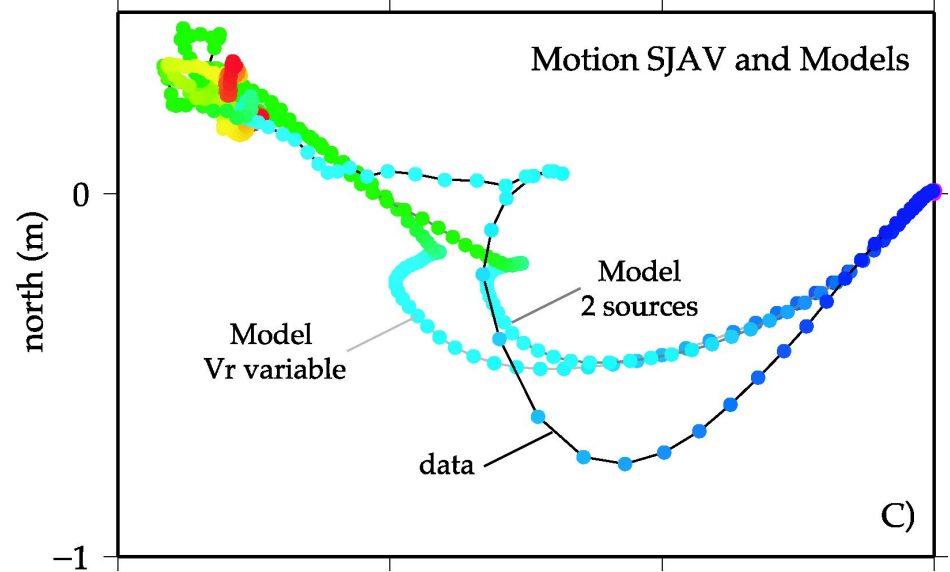
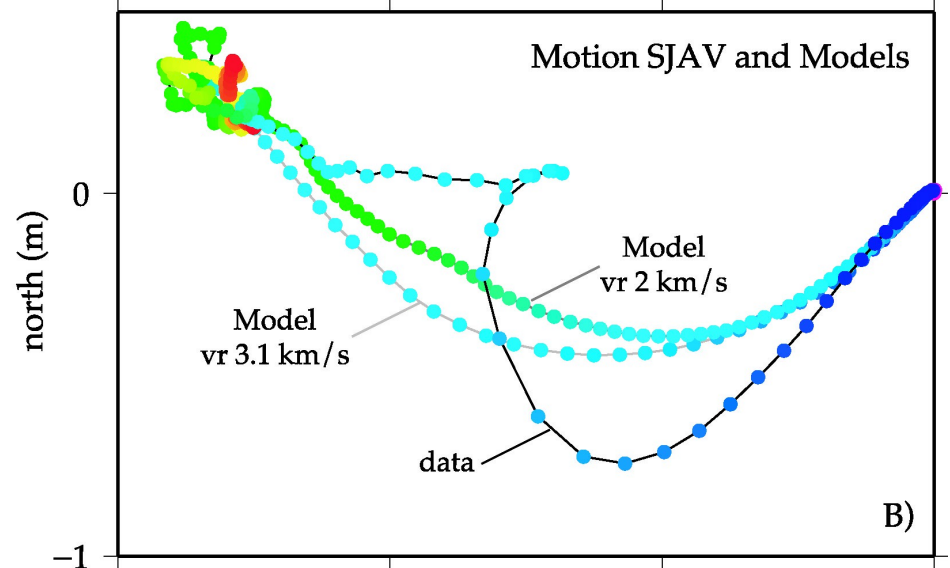
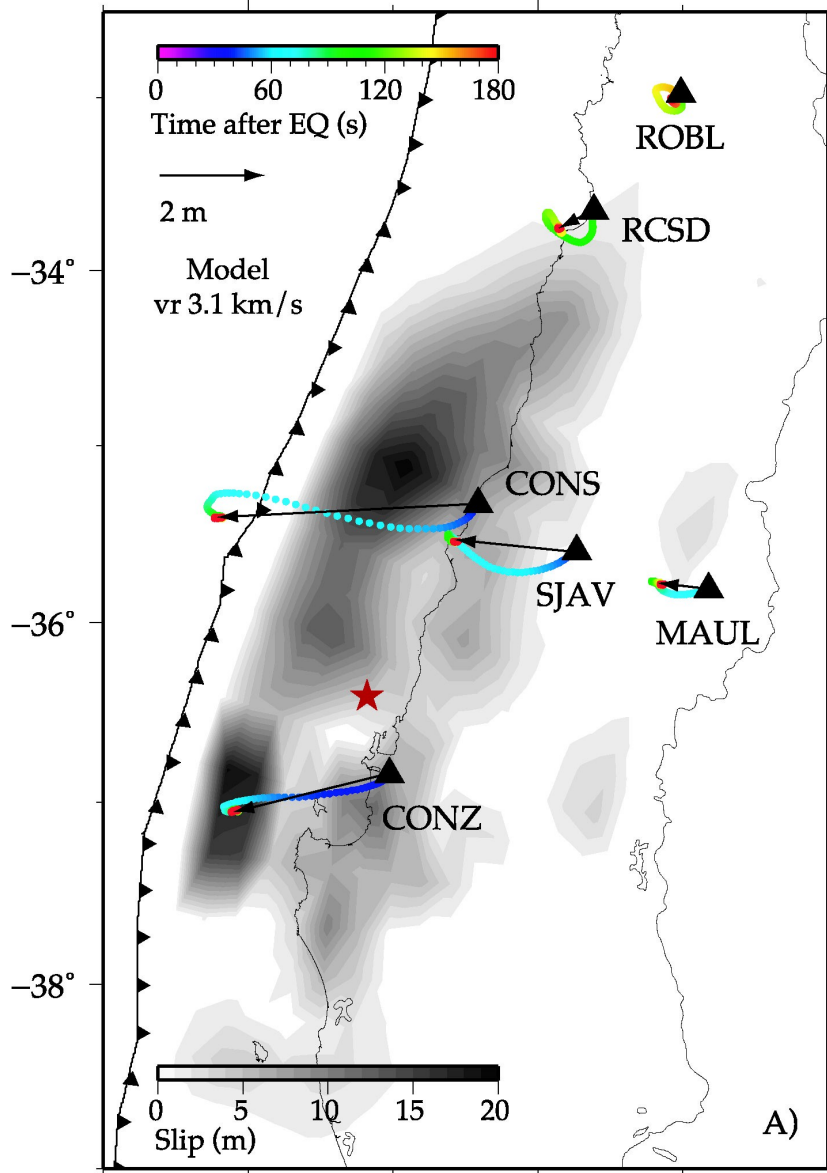


Colbun



Maule Eq 27-Feb-2010

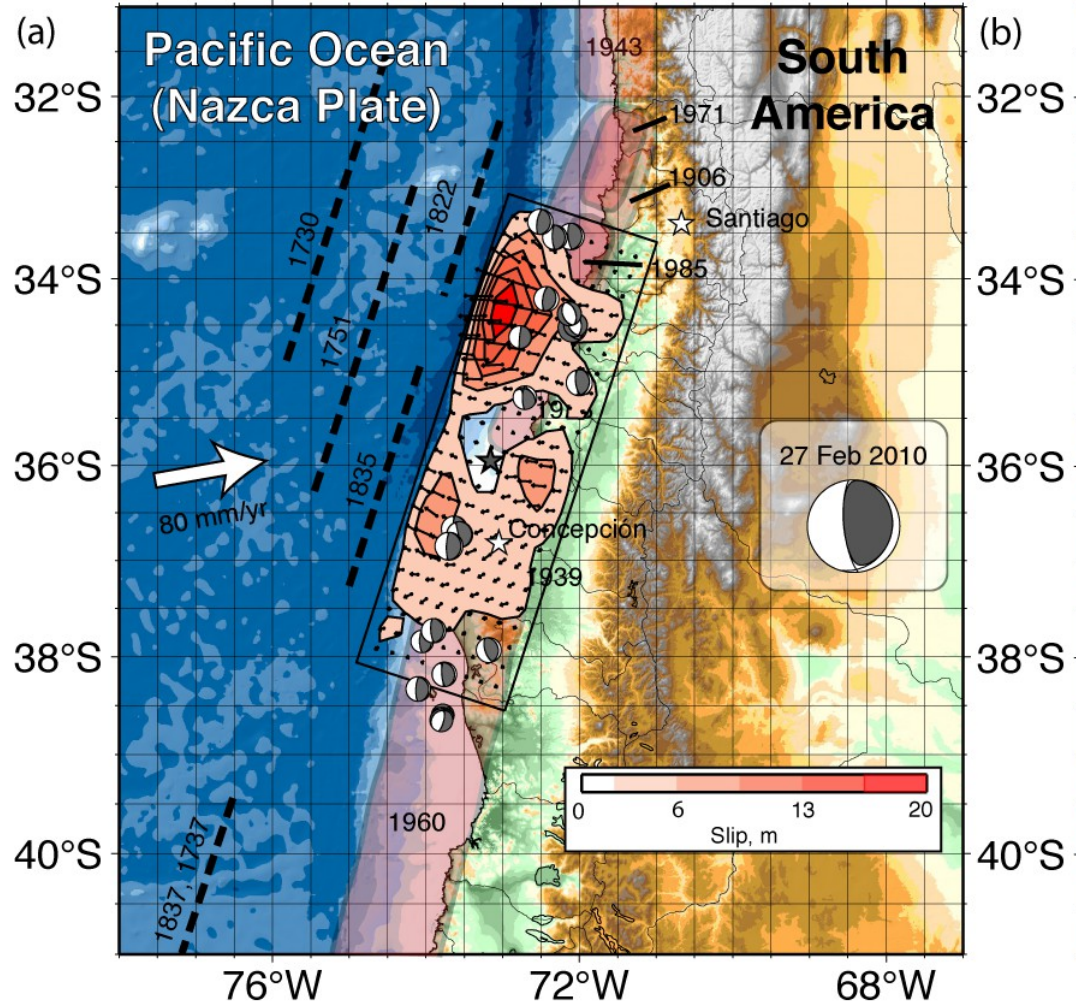




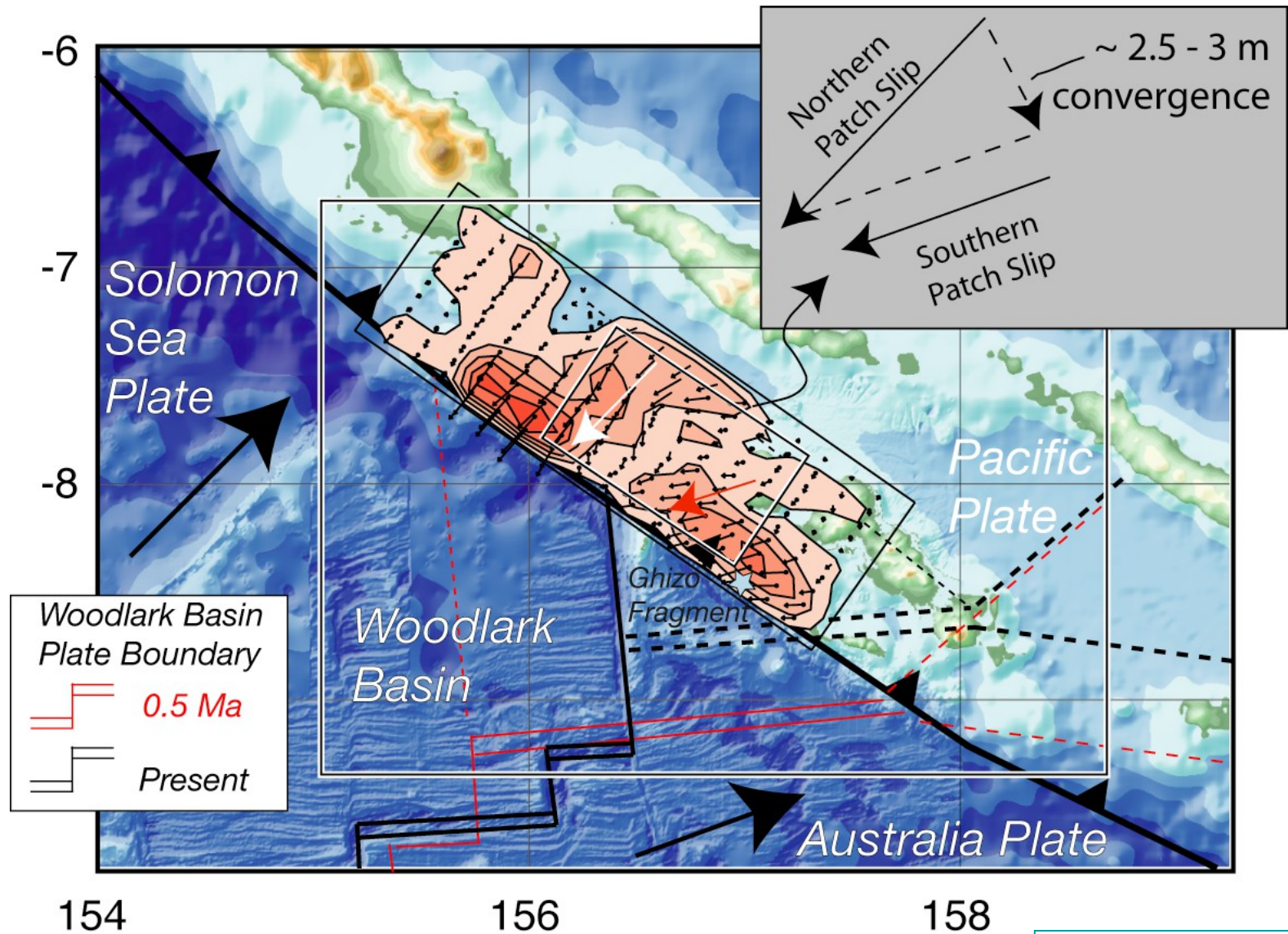
GPS haute fréquence: 1 point par seconde => trajectoire des stations pendant le séisme. Comparaisons avec modèles théoriques

Ruptures superficielles ?

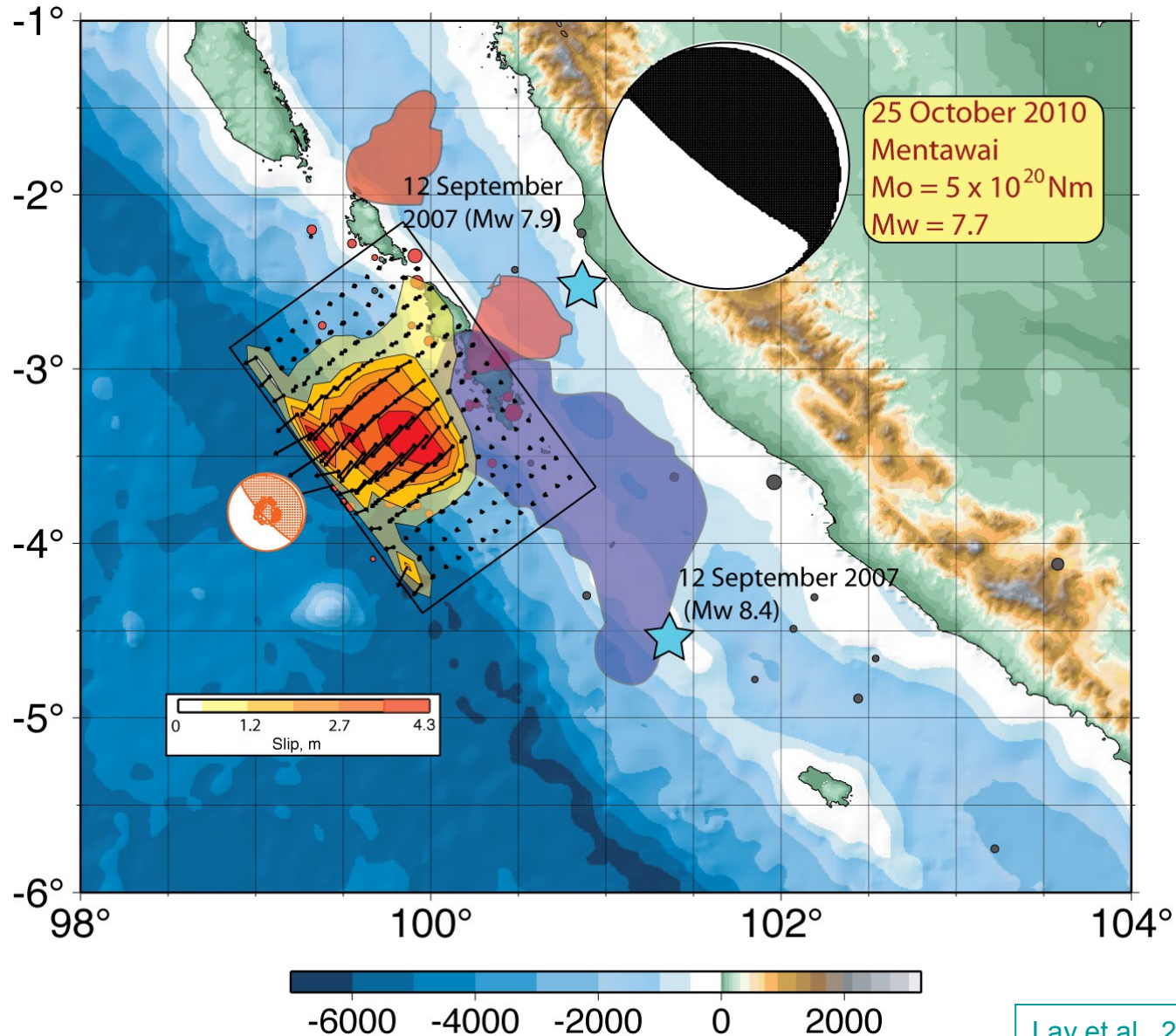
Chili (Maule) $M_w=8.8$, 27 Février 2010



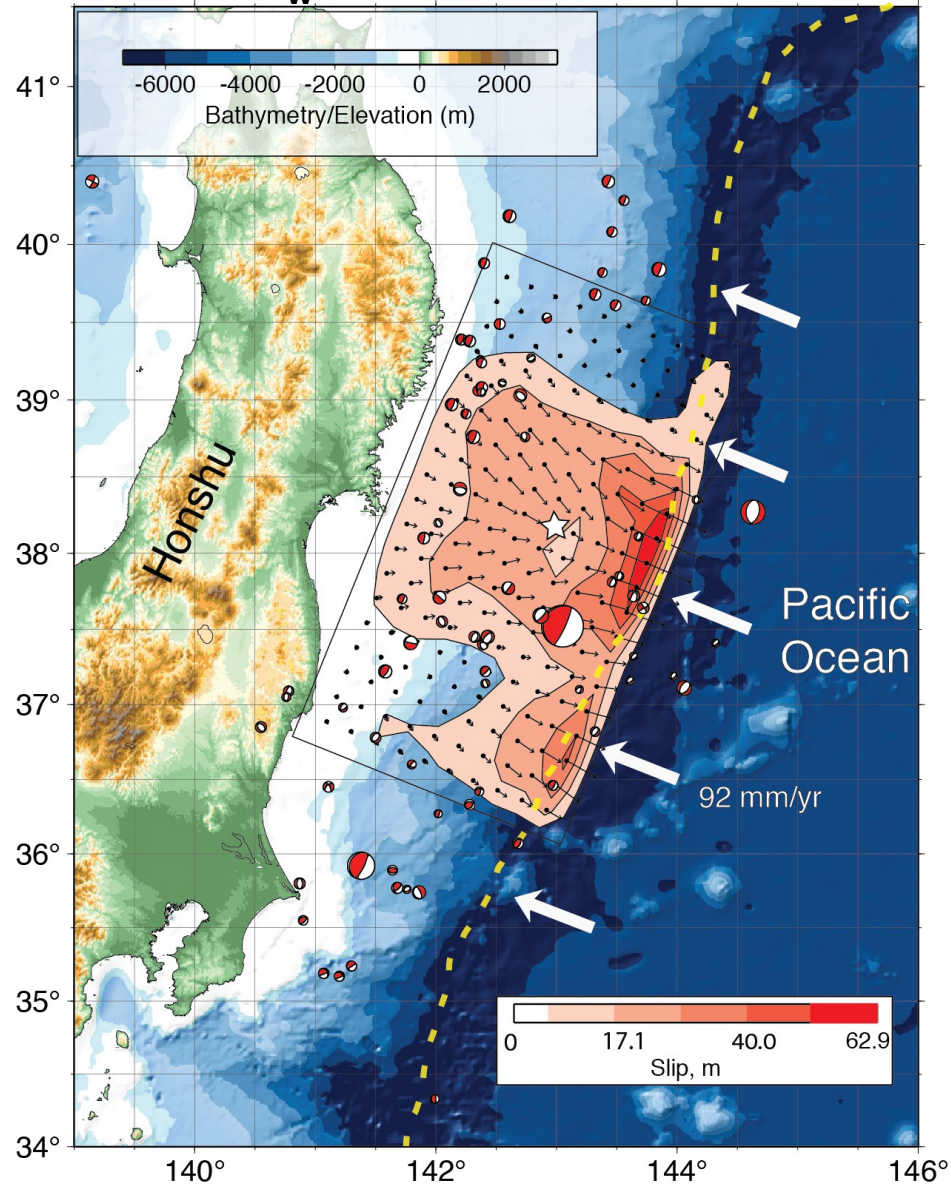
Iles Solomon $M_w=8.1$, 1 Avril 2007



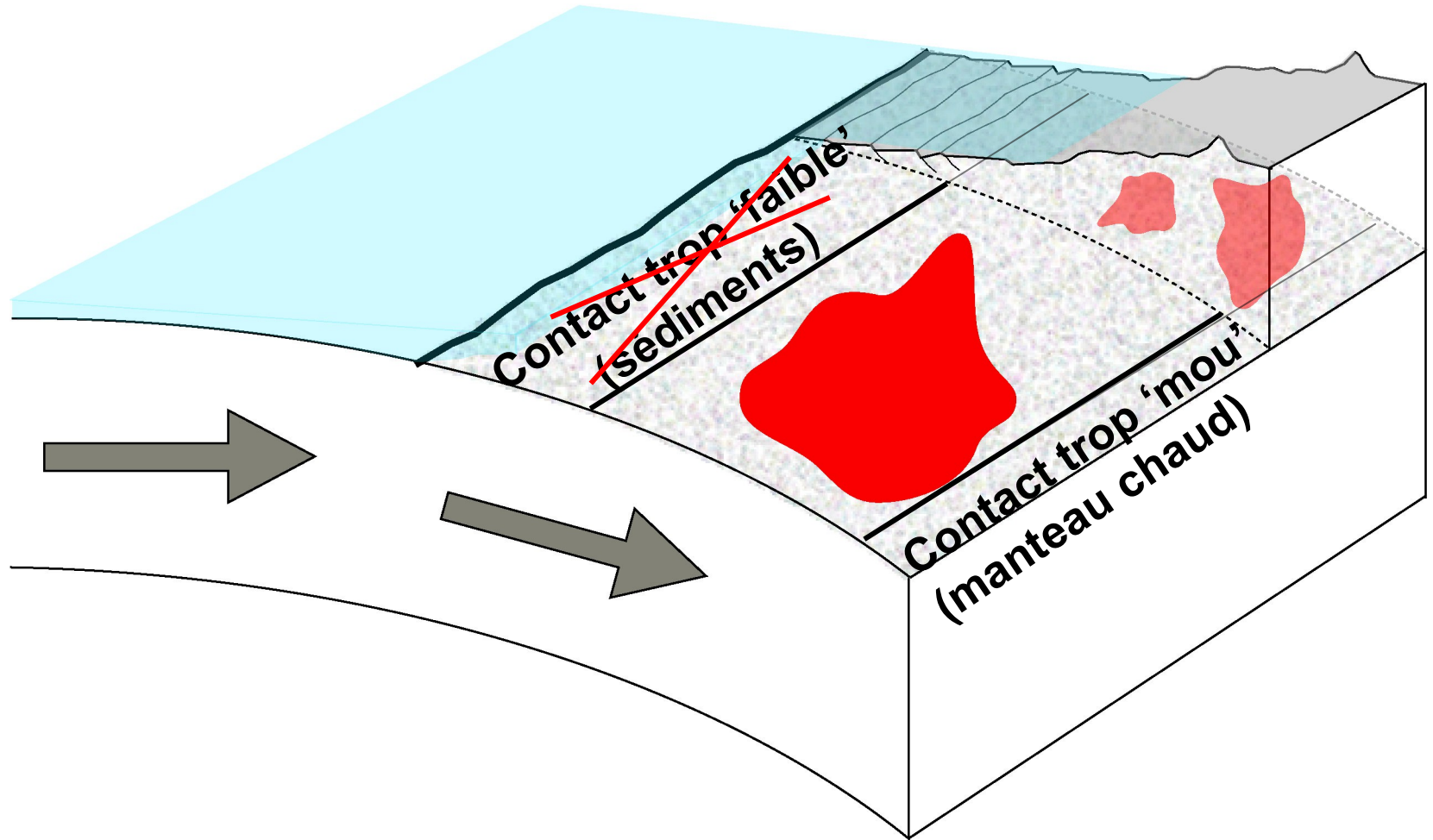
Iles Mentawai $M_w=7.7$, 25 Octobre 2010



Tohoku $M_w=9.0$, 11 mars 2011

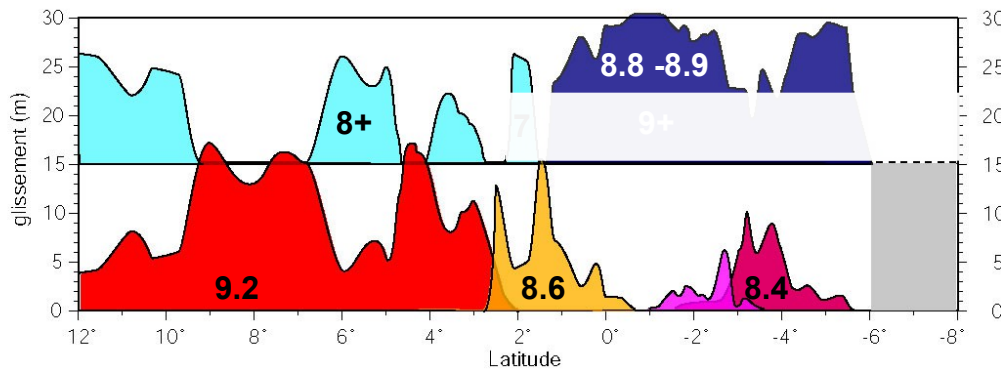


Donc : la zone superficielle de contact casse aussi

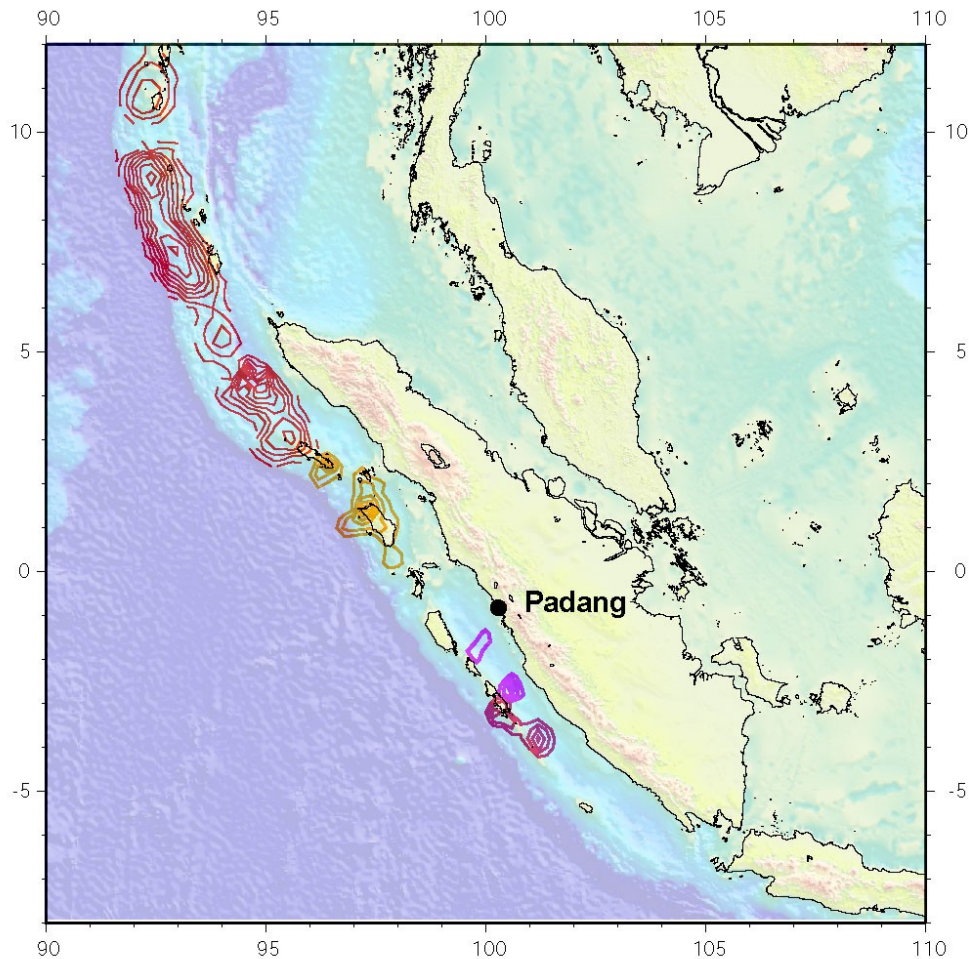


Quantification de l'aléa sismique

Où ?

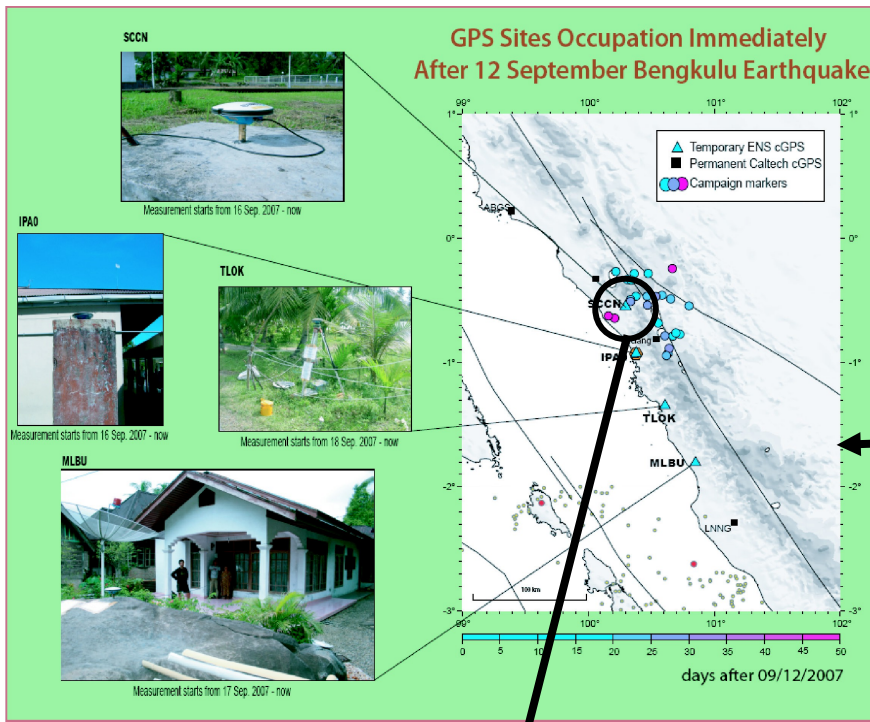


7.5 m uniform slip
 =
 250 years at 3 cm/y

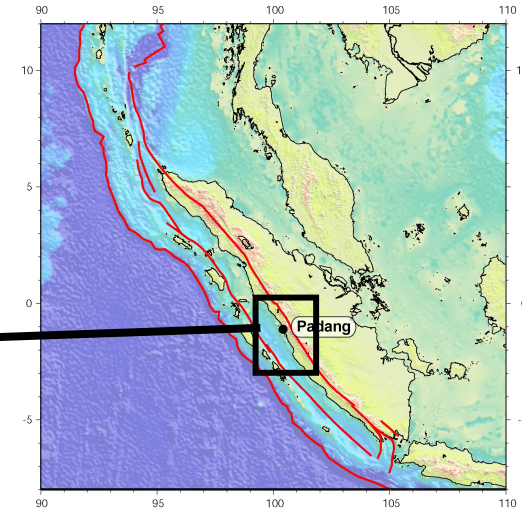


Chen Ji (Caltech)
slip distribution
Inverted from
seismic & GPS data

GPS Sites Occupation Immediately After 12 September Bengkulu Earthquake



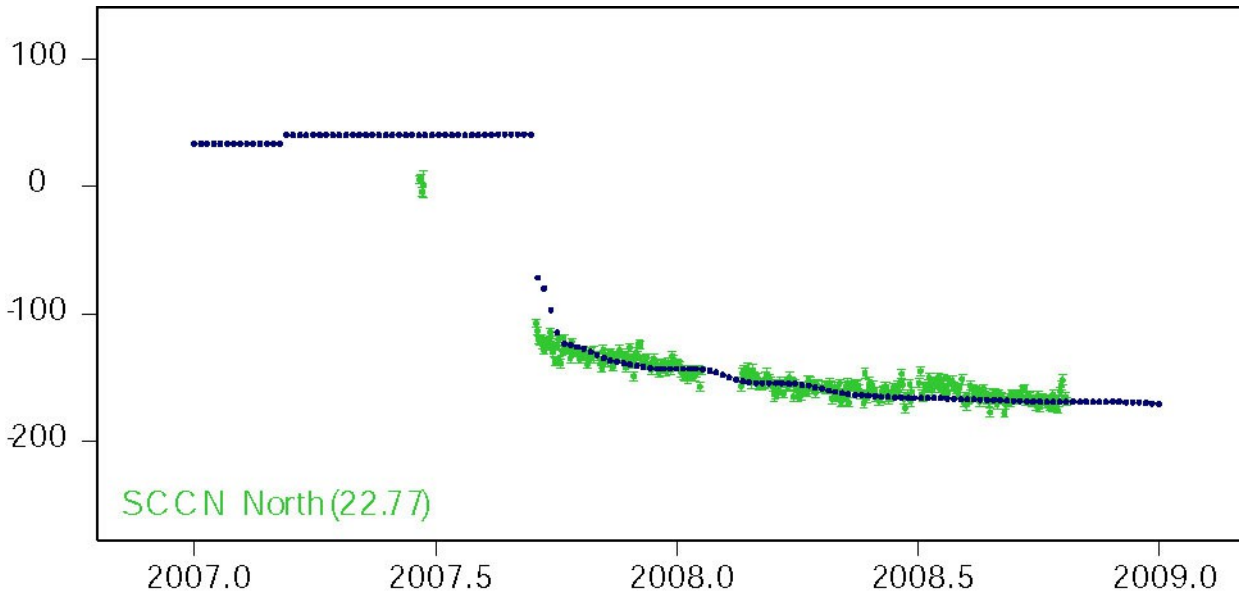
Padang still at risk



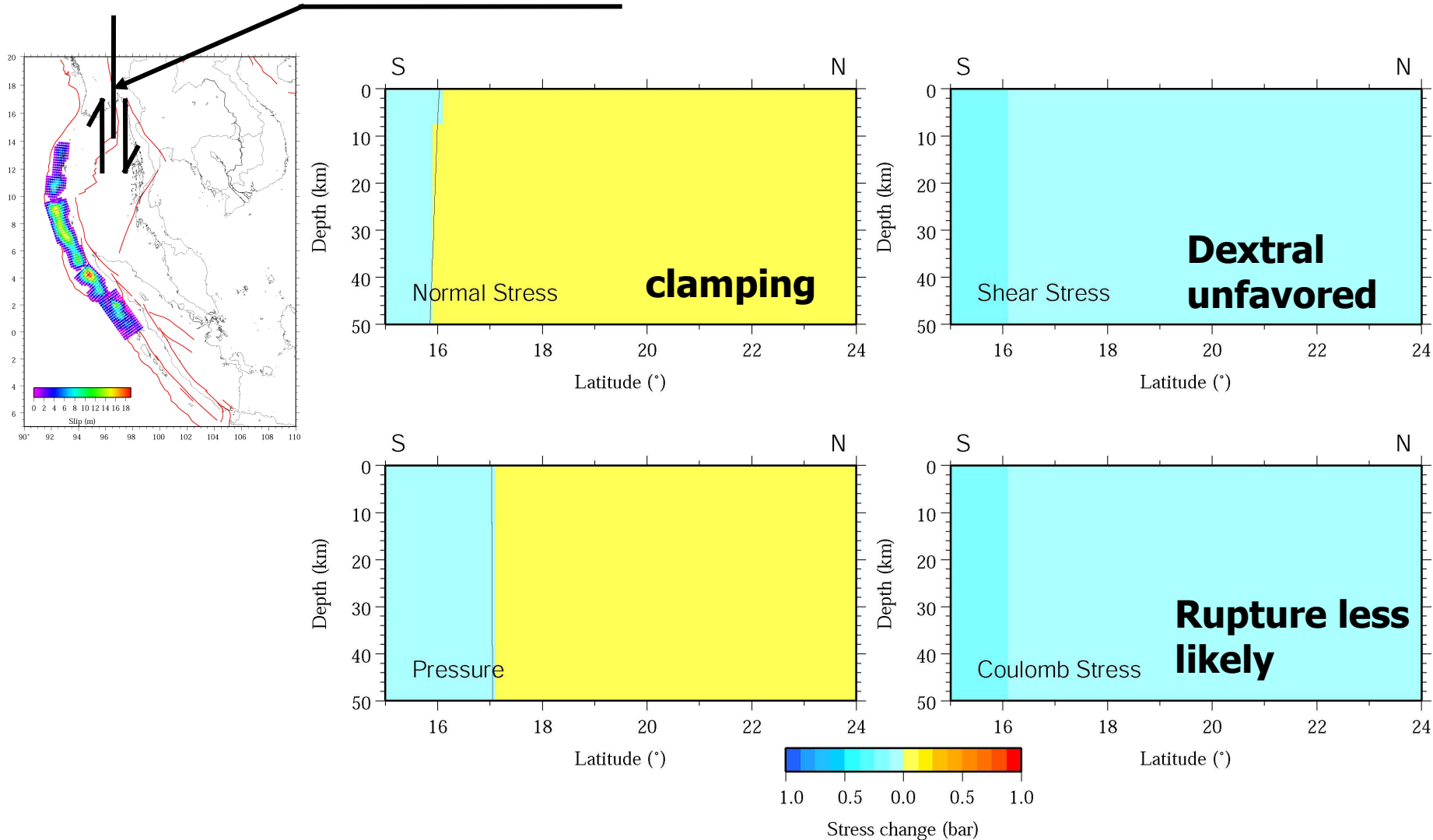
Two scenarios:

1. post-seismic « normal »
 => still locked
 => rupture later (the latest the biggest)

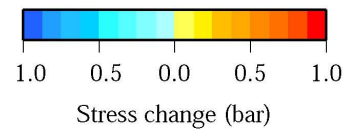
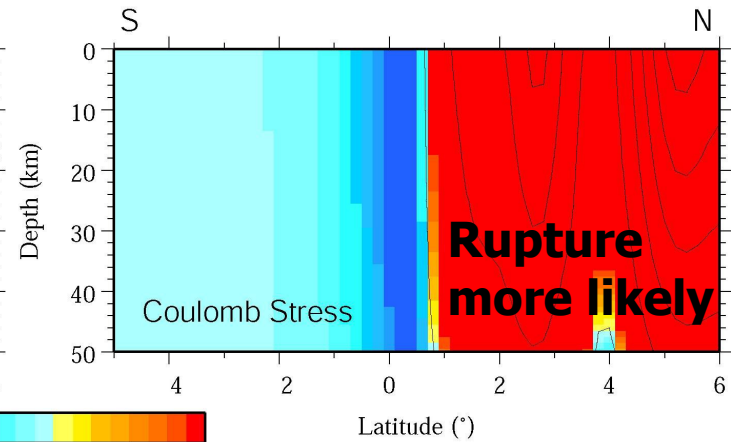
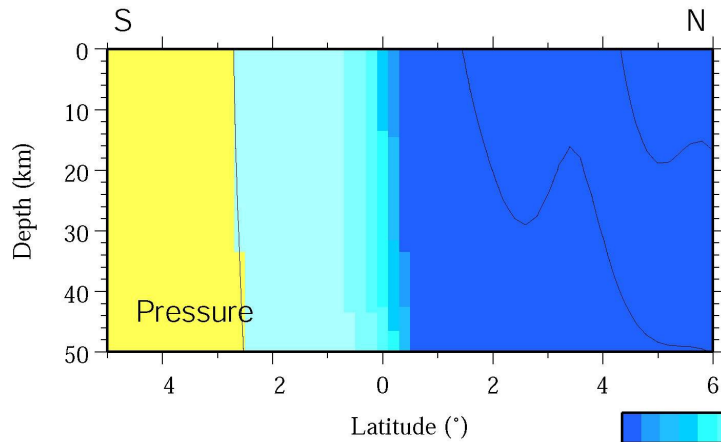
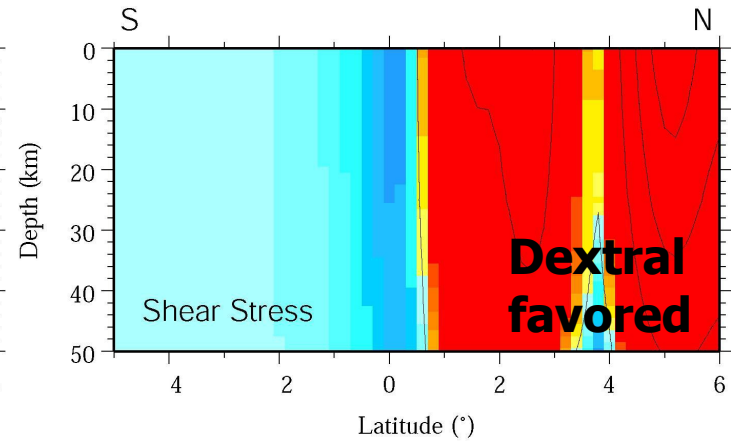
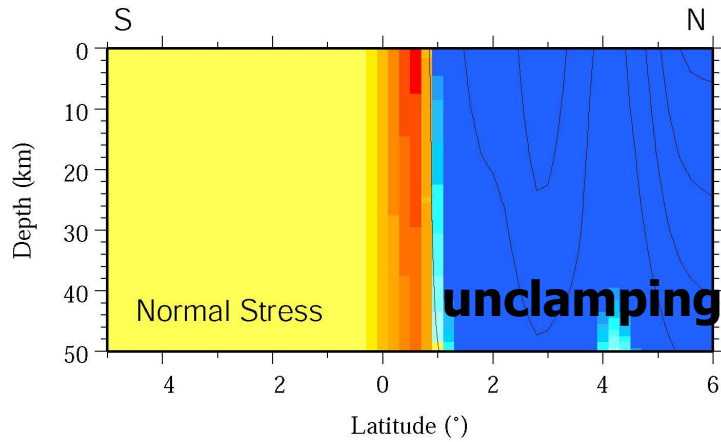
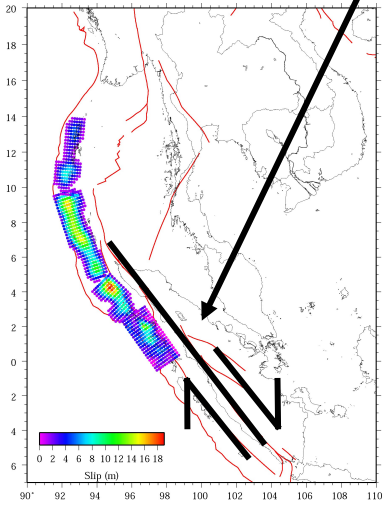
2. Post-seismic
 « more than normal »
 => silent dissipation of elastic deformation
 => smaller future rupture



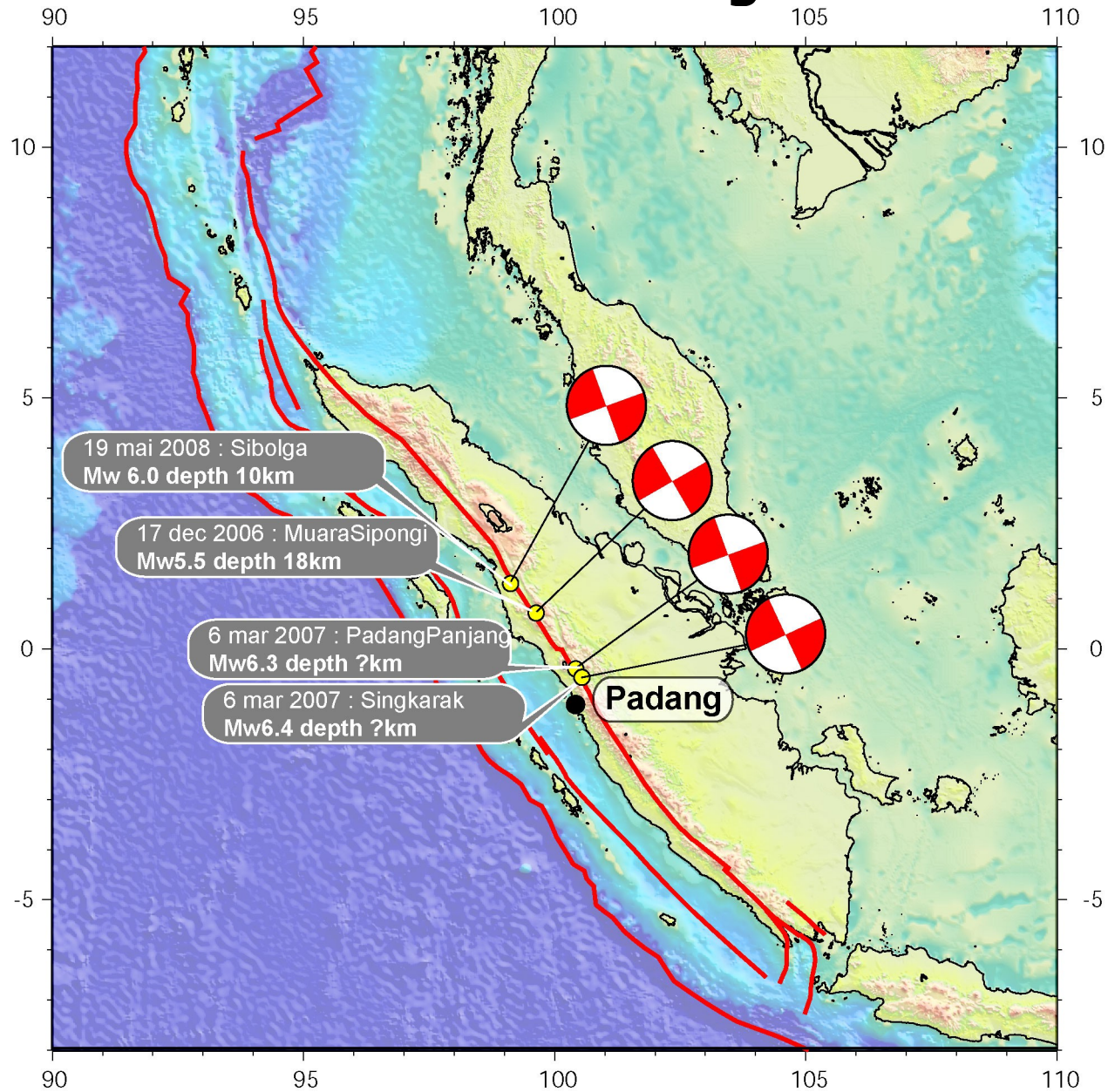
Loading of Sagaing fault: low everywhere and reverse!



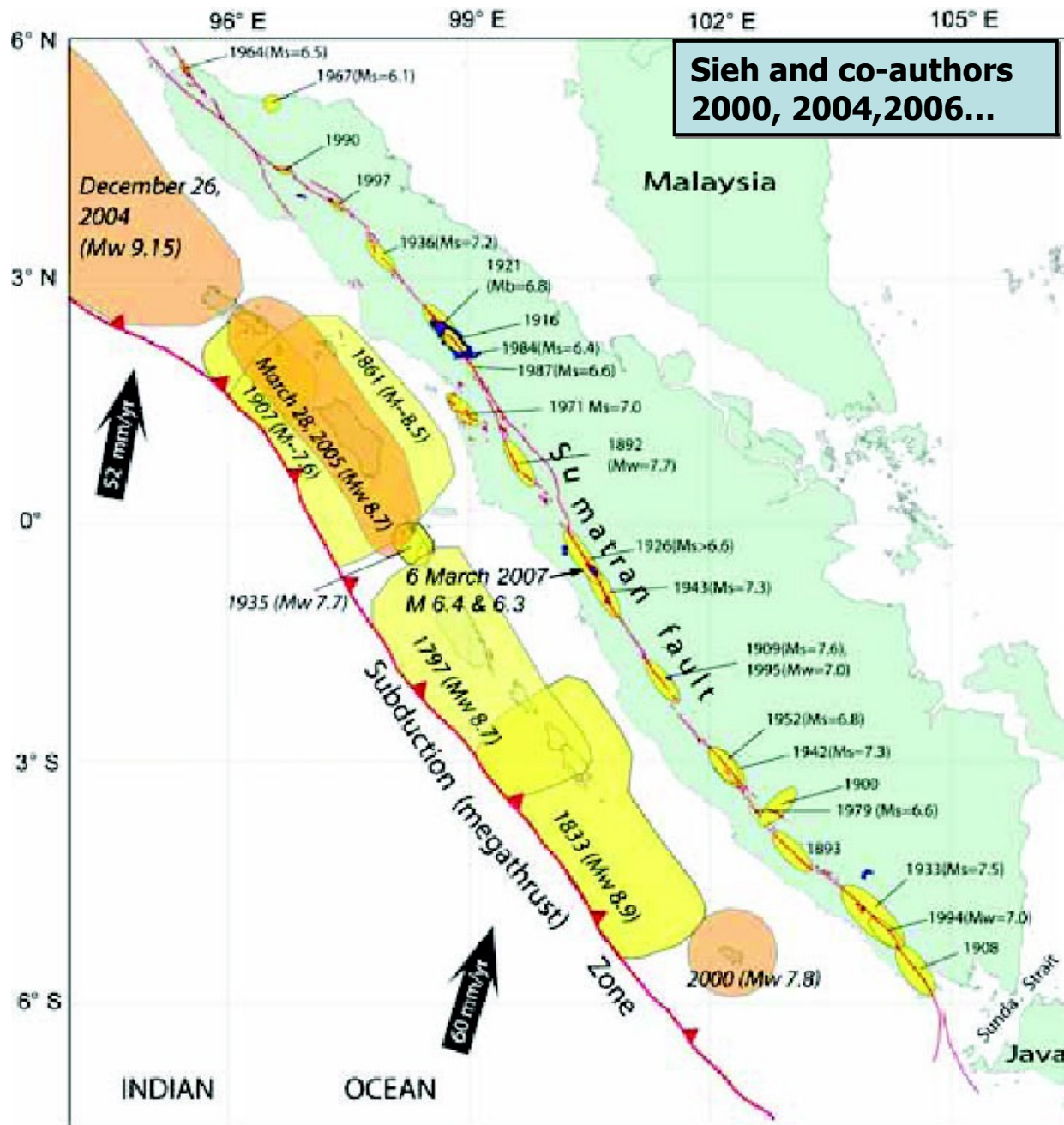
Loading of Great Sumatra fault: high above 0°N



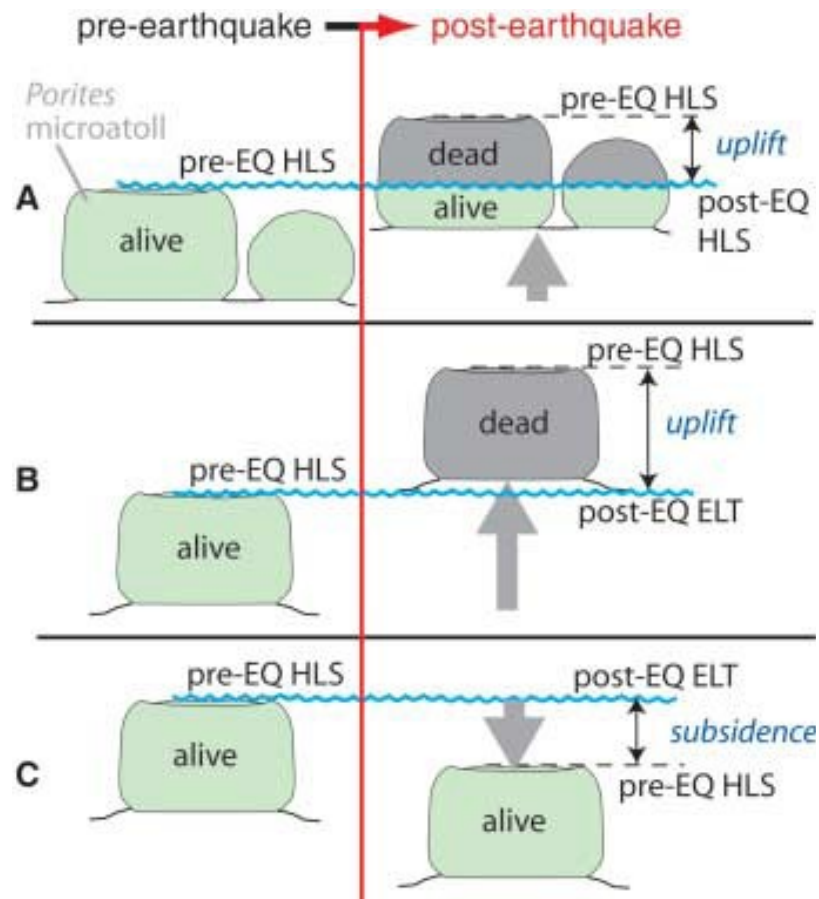
Loading of Great Sumatra fault: high above 0°N



QUAND ?



Coral reefs records



- Three scenarios for measuring vertical deformation using *Porites* coral microatolls.
- (A) Uplift recorded as the difference between pre- and post-earthquake highest level of survival (HLS).
 - (B) Uplift as separation between pre-earthquake HLS (pre-EQ HLS) and the model elevation of postearthquake extreme low tide (post-EQ ELT).
 - (C) Subsidence measured upward from pre-earthquake HLS to post-earthquake ELT.

Coral reefs records



Cross sections in vertical slabs of *Goniastrea* coral reef

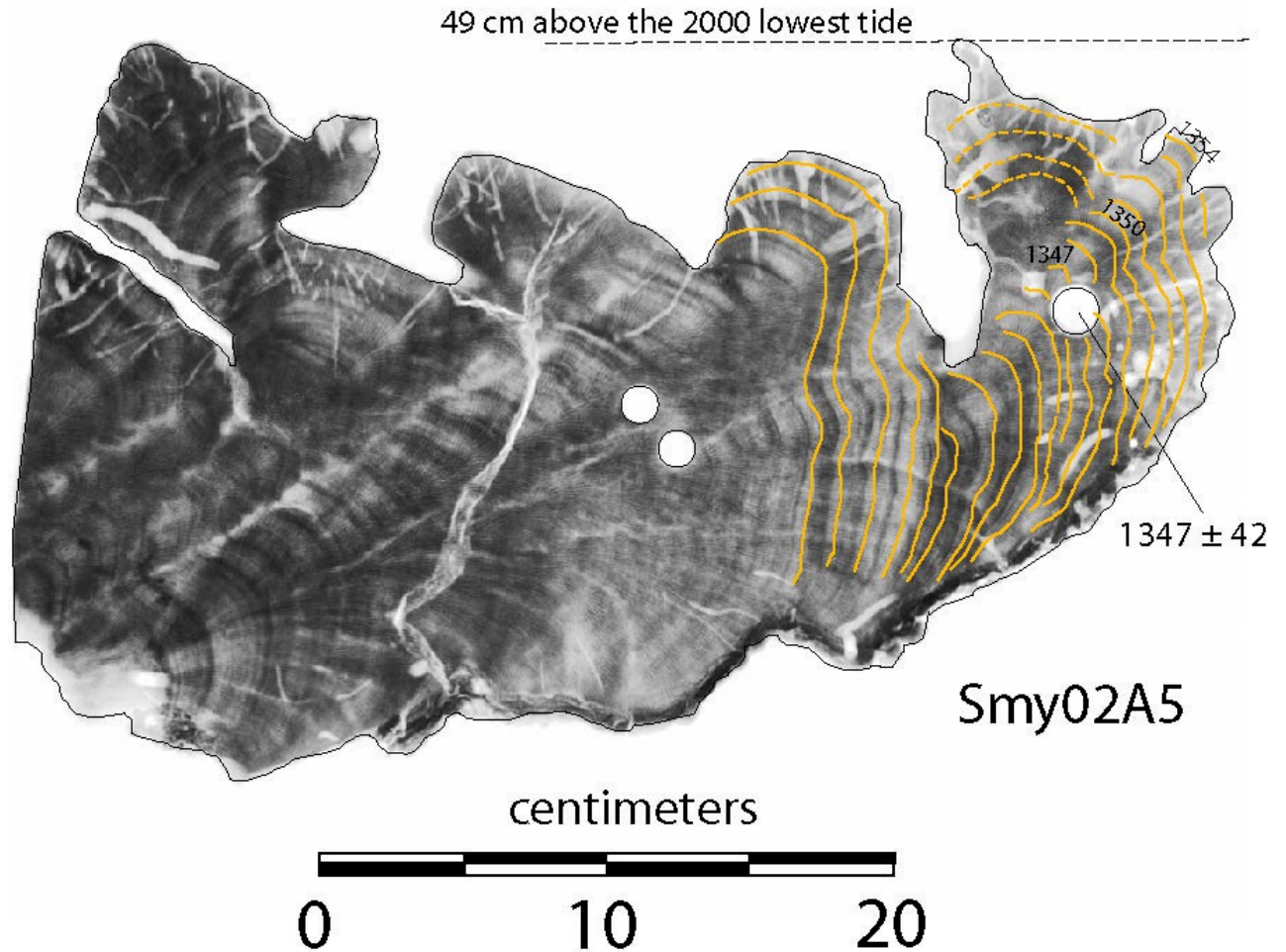


Figure S16. Cross-section of slab Smy02A5.

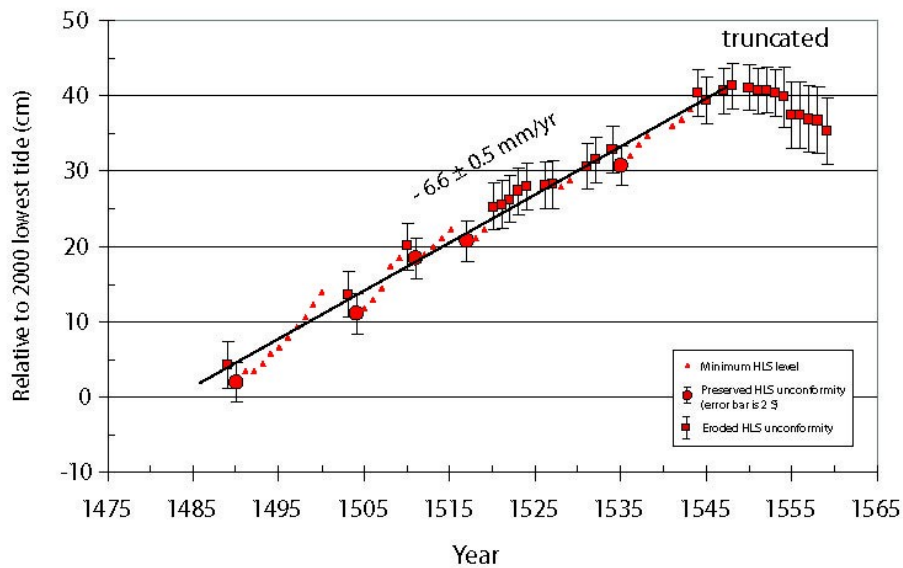
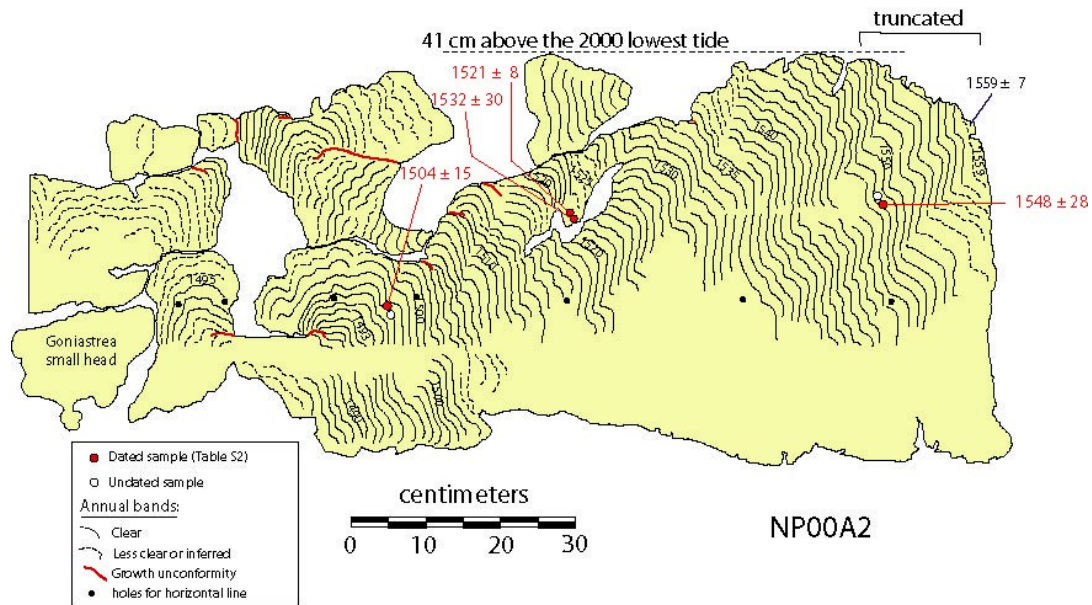
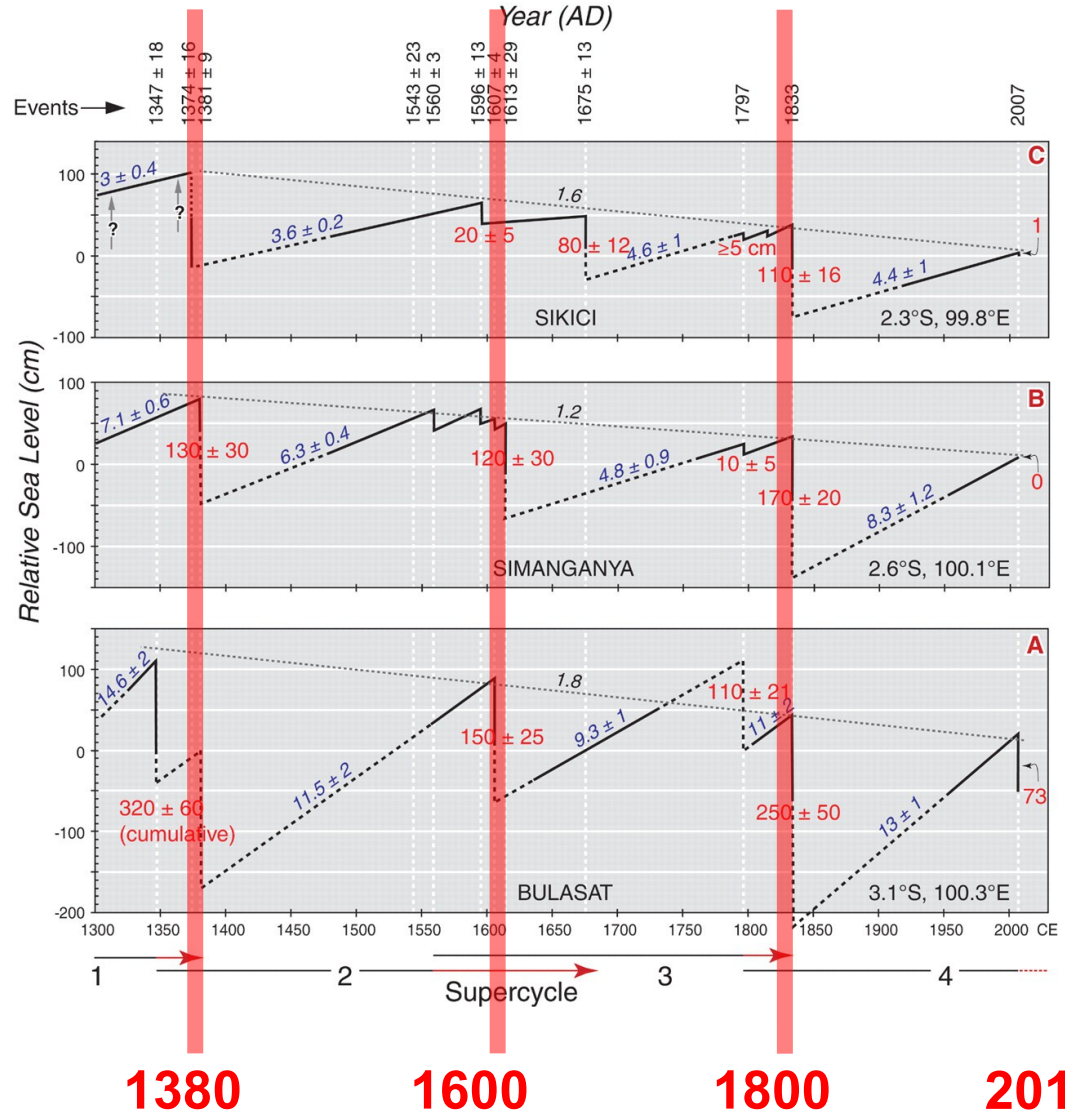
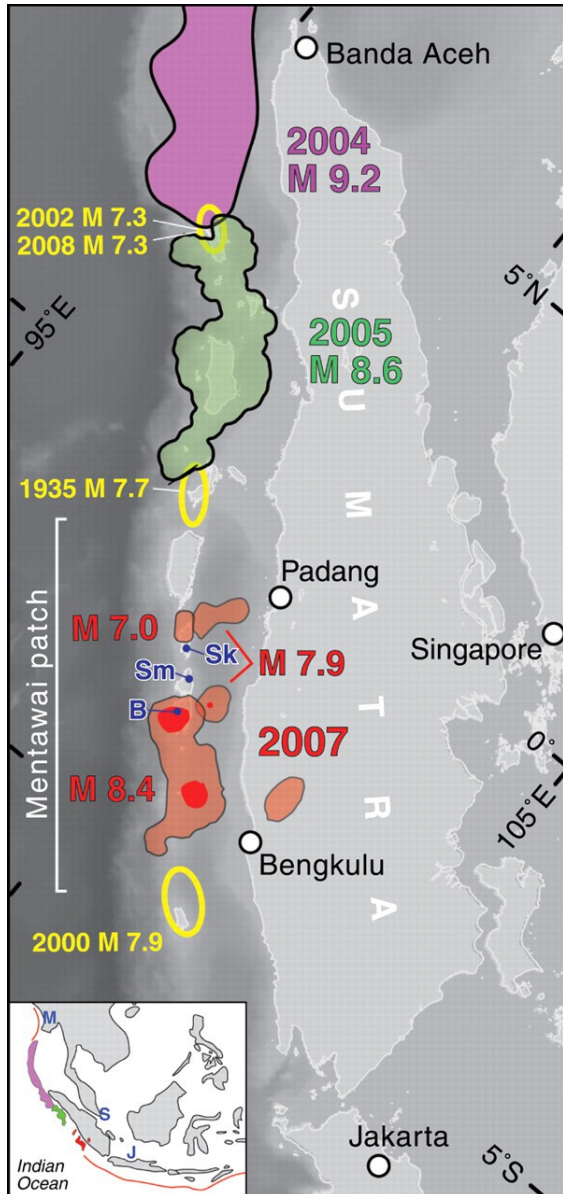


Figure S17. Cross-section of slab Np00A2 and graph of sea level history derived from the cross-section.

Coral reefs records

Sieh et al., 2008



...encore du travail (difficile) en
vue.....



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